

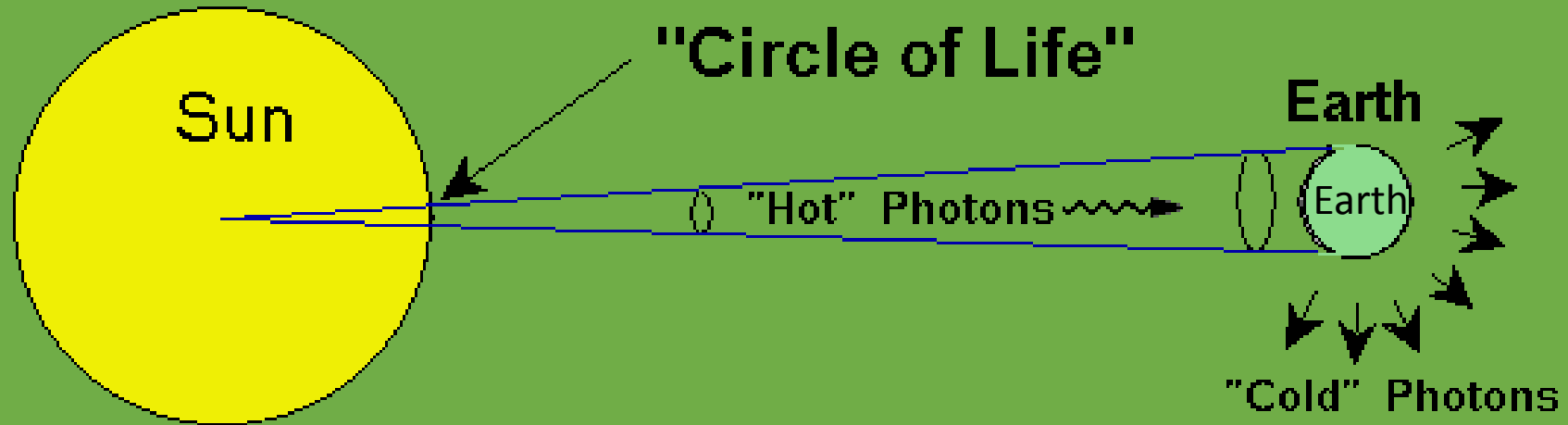
Mediterranean algal mariculture potential under global climate change

Zvy Dubinsky

Faculty of Life Sciences, Bar-Ilan University, Israel.

Solar radiation and the underwater light field

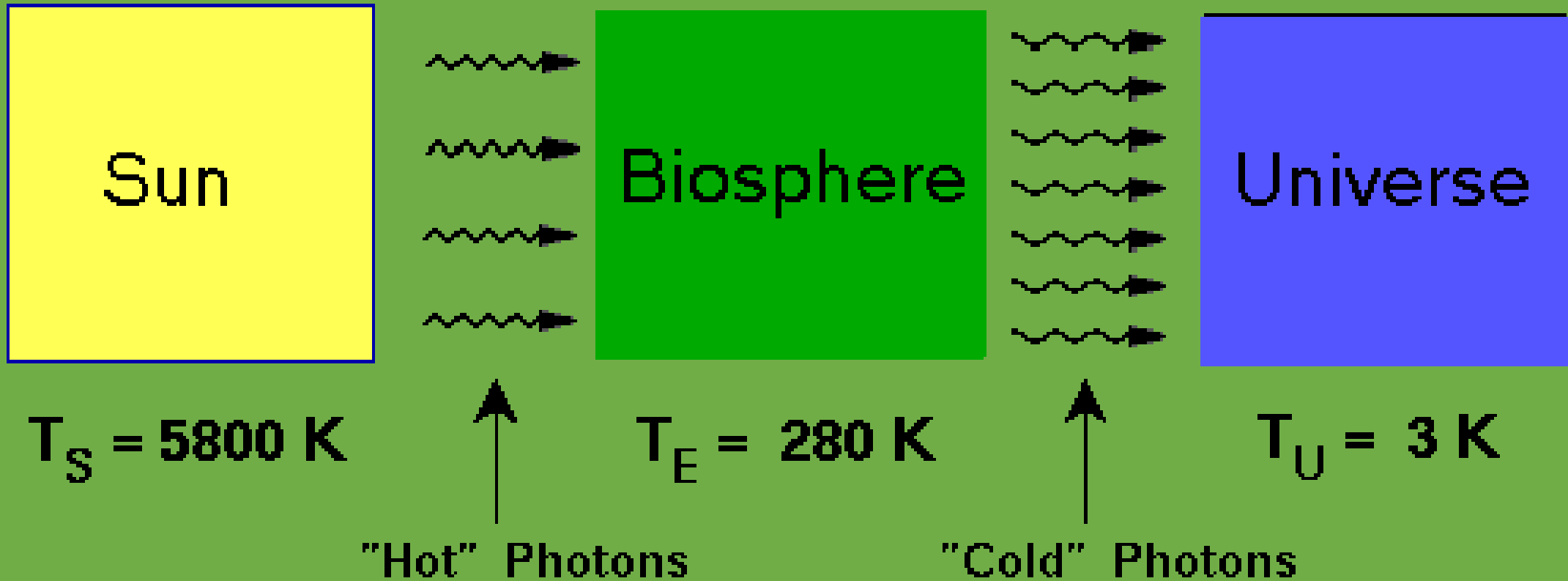
The sun as the source of energy for life on Earth



The energy needed to sustain life on Earth arrives as **sunlight**, through the process of **PHOTOSYNTHESIS**, by which light energy is converted into energy in organic compounds supporting all of life processes. That energy is then converted to **heat**, and dissipated in space.

The flow of energy through the biosphere

Thermodynamics of Life



Radiation and temperature

- The temperature of the sun's core is 20,000,000 degrees, but the radiation of any body depends on its surface temperature, which for the sun is **5760 °K**,

- The Stefan–Boltzmann constant

- **$\sigma = 5.673 \times 10^{-8} \text{ Js}^{-1} \text{ m}^{-2} \text{ K}^{-4}$**

- The radiation emitted by a square meter of the body surface (flux) is

- **$\mathcal{R} = \sigma \text{ K}^4 \text{ Js}^{-1} (\text{W}) \text{ m}^{-2}$**

- $\mathcal{R} = \sigma \times T^4 \text{ Wm}^{-2}$

$$\mathcal{R} = 5.673 \times 10^{-8} \times 5760^4 \text{ Wm}^{-2}$$

- For the entire sun

$$\mathcal{R} = 5.673 \times 10^{-8} \times 5760^4 [\text{Wm}^{-2}] \times 4\pi r^2 [\text{m}^2] \text{ W}$$

- The radius of the sun is $r = 6.965 \times 10^8$

Wein's law

$$\lambda_{\max}[\text{m}] = \frac{289.7 \times 10^{-5} \text{ }^{\circ}\text{K} [\text{m}]}{T[{}^{\circ}\text{K}]}$$

For the sun whose surface is at 5760 °K

$$\lambda_{\max} = 5.18 \times 10^{-7} \text{ m} = 518 \text{ nm}$$

Spectral curves of radiating bodies

Spectral curves for blackbody radiators

↑
Increasing flux

5000 K

4000 K

3500 K

3000 K

2000 K

1×10^{-7}

5×10^{-7}

1×10^{-6}

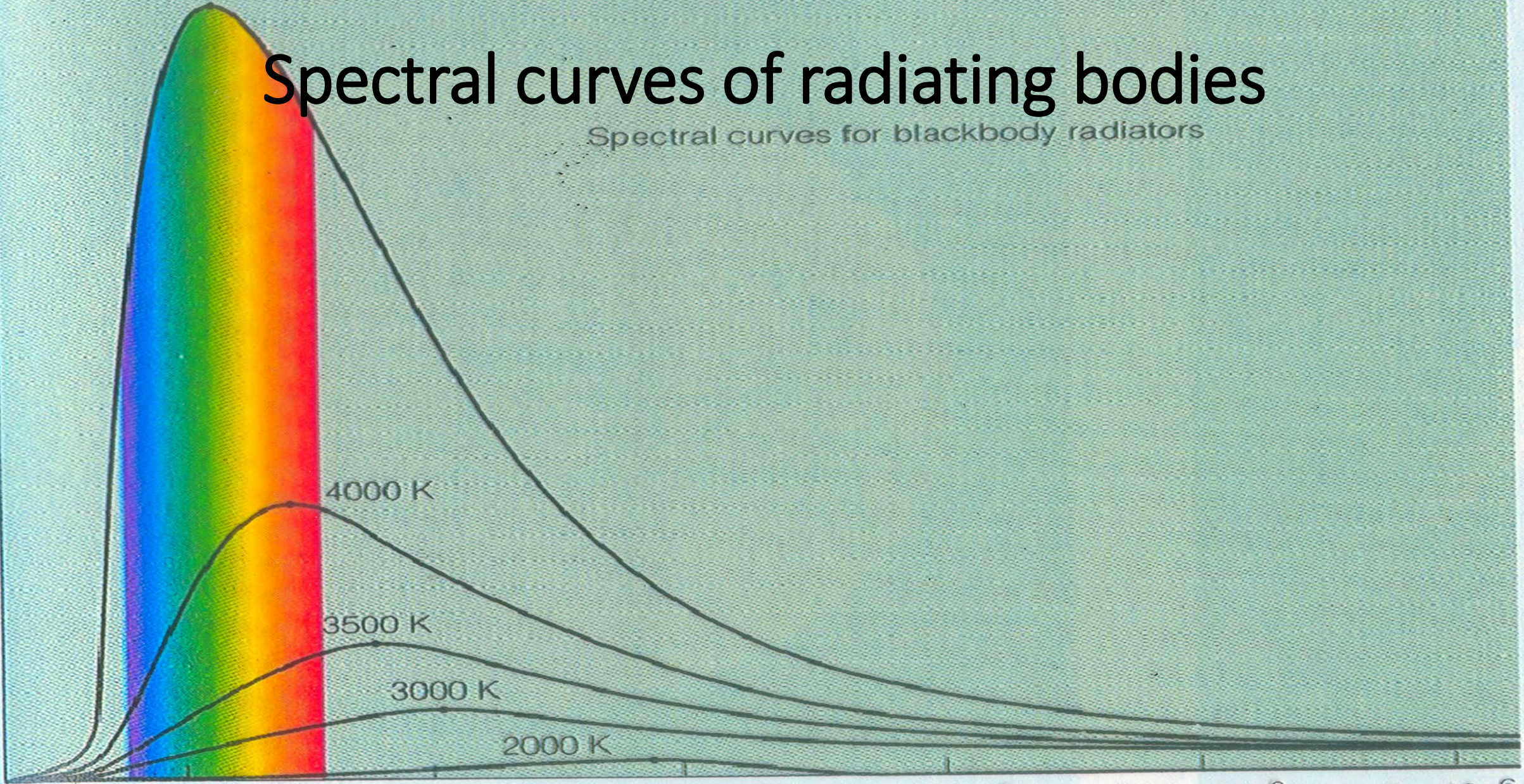
1.5×10^{-6}

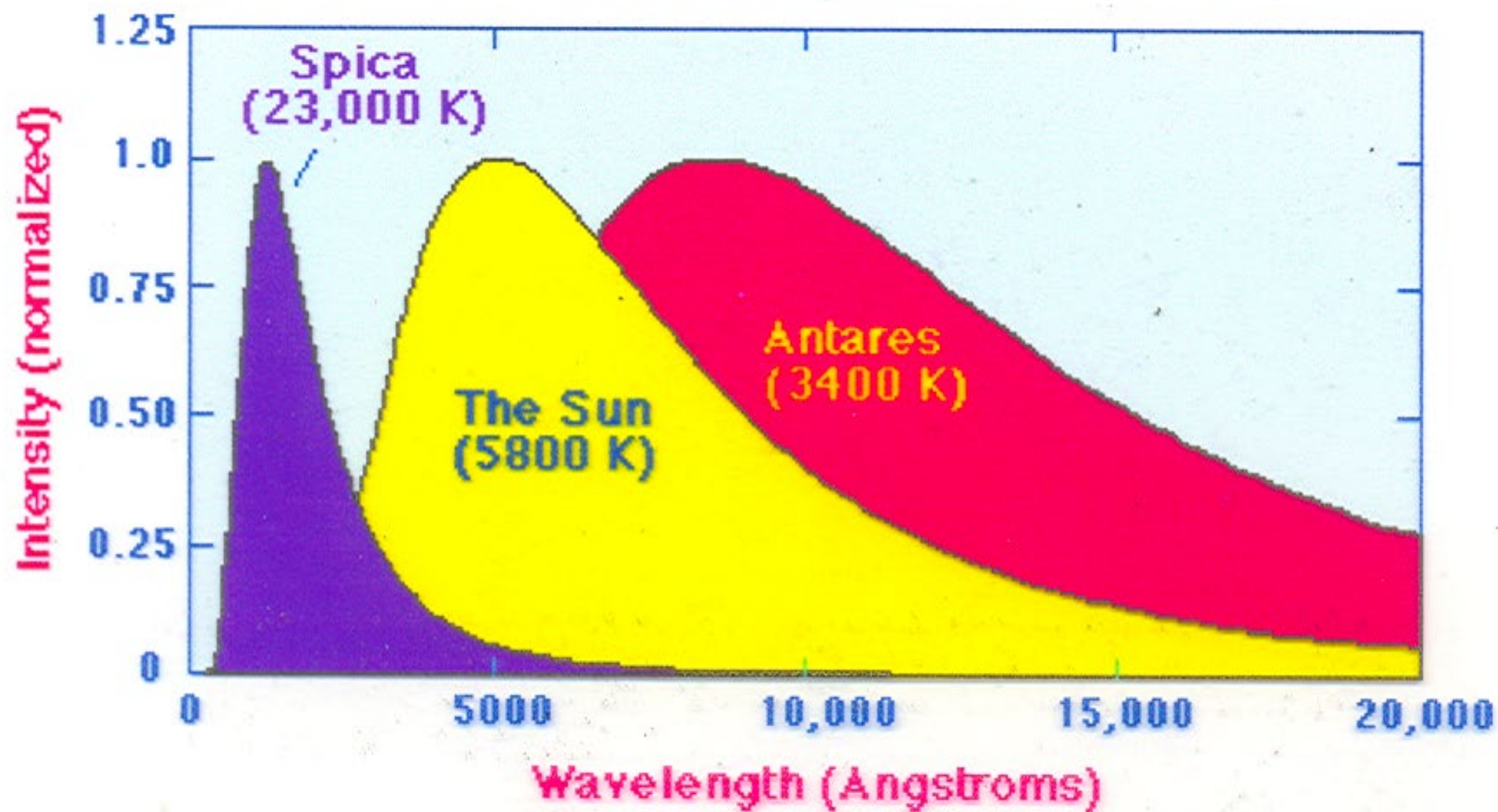
2×10^{-6}

2.5×10^{-6}

3×10^{-6}

Wavelength (m)





Total solar energy
reaching Earth
depends on
Sun's .1

K ,temperature
Sun's surface .2
area $4\pi r^2$
Earth's .3

distance from the
sun, d Earth's
Projected area .4
of Earth's shadow
 πa^2

Earth's surface .5
area $4\pi a^2$

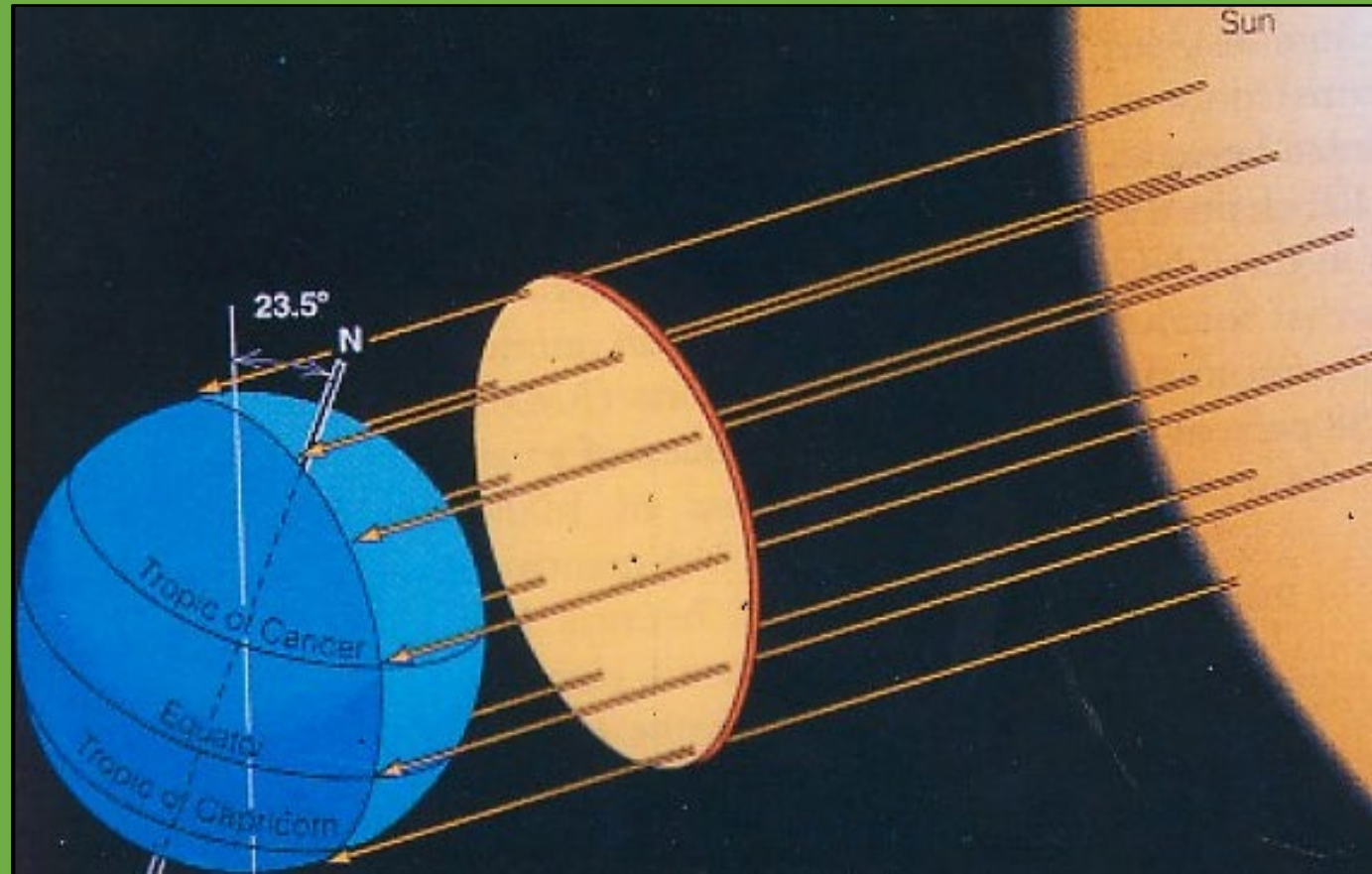
Earth's radius

$$a = 6.37 \times 10^6 \text{m}$$

וכל m^2 של פניו יקבל בממוצע

$$S = \frac{\sigma \text{ } ^\circ\text{K}^4 \times 4\pi r^2 \times \pi a^2}{4\pi d^2 \times 4\pi a^2}$$

$$S = \frac{\sigma \text{ } ^\circ\text{K}^4 \times (r^2/d^2) \times \pi a^2}{4\pi a^2}$$

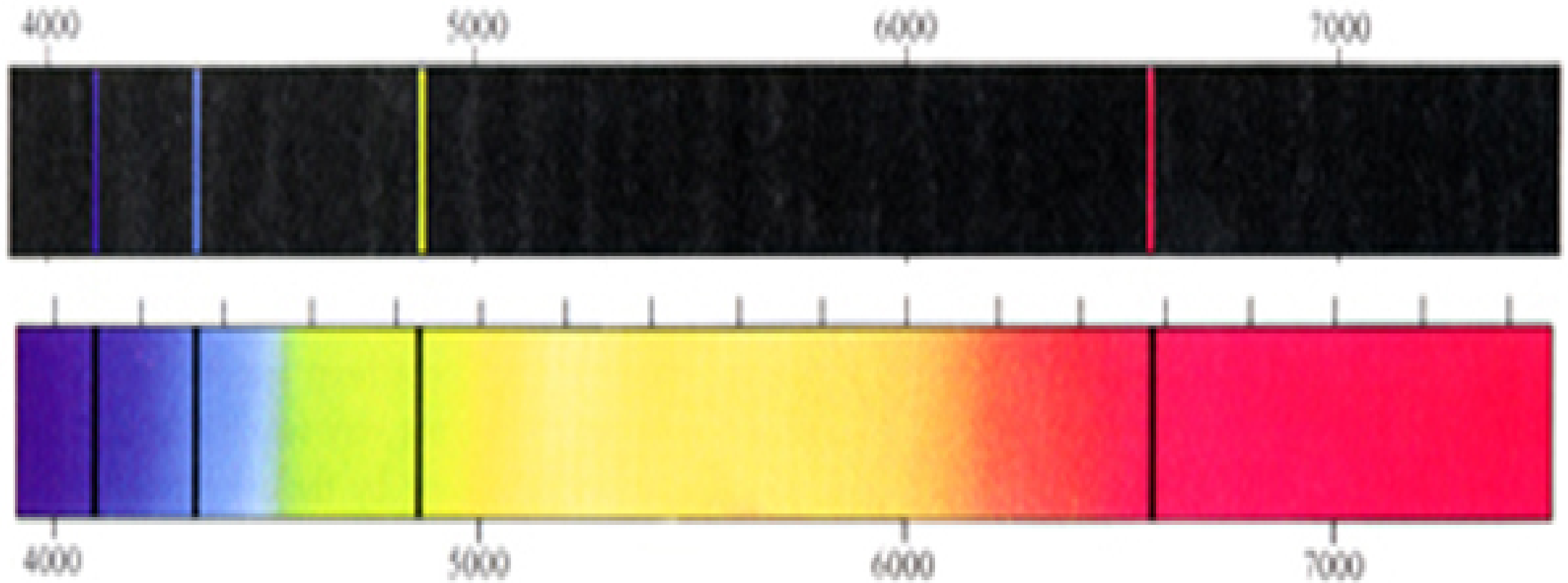


Albedo of the Earth = 0.39

Of that, part, the albedo (between 1 and 0, for Earth 0.4), is reflected back to space.

Thus the surface temperature of any planet assuming no atmosphere and no residual internal heat can be calculated as follows.

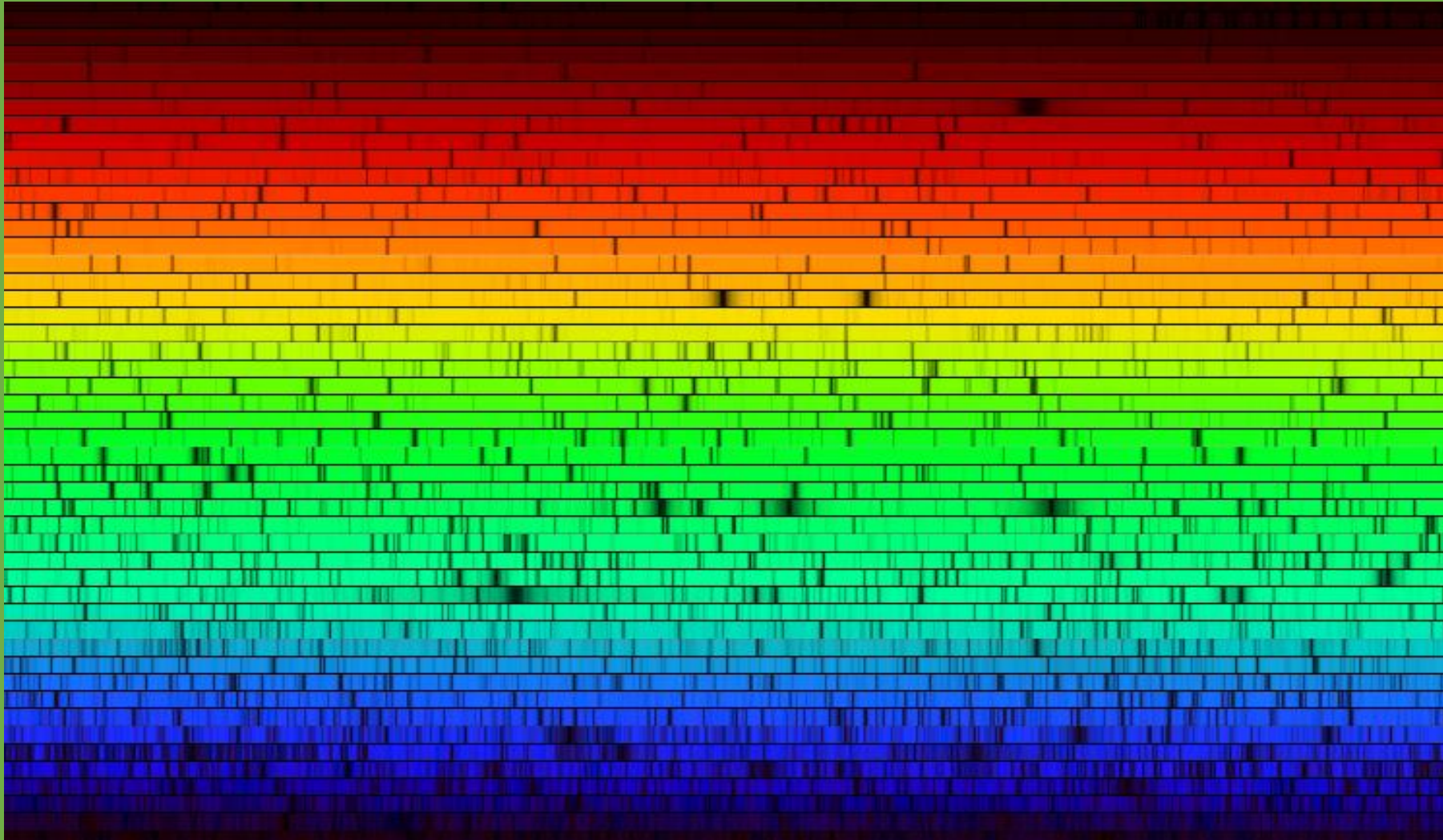
$$\begin{aligned} \text{Energy}_{\text{in}} &= \text{Energy}_{\text{out}} \\ \sigma \text{ } ^\circ\text{K}_{\text{SUN}}^4 \times (r^2/d^2)/4 \times (1 - 0.4) &= \sigma \\ \text{ } ^\circ\text{K}_{\text{surface}}^4 & \\ \text{ } ^\circ\text{K}_{\text{surface}}^4 &= 0.6 \text{ } ^\circ\text{K}_{\text{SUN}}^4 \times (r^2/d^2)/4 \end{aligned}$$



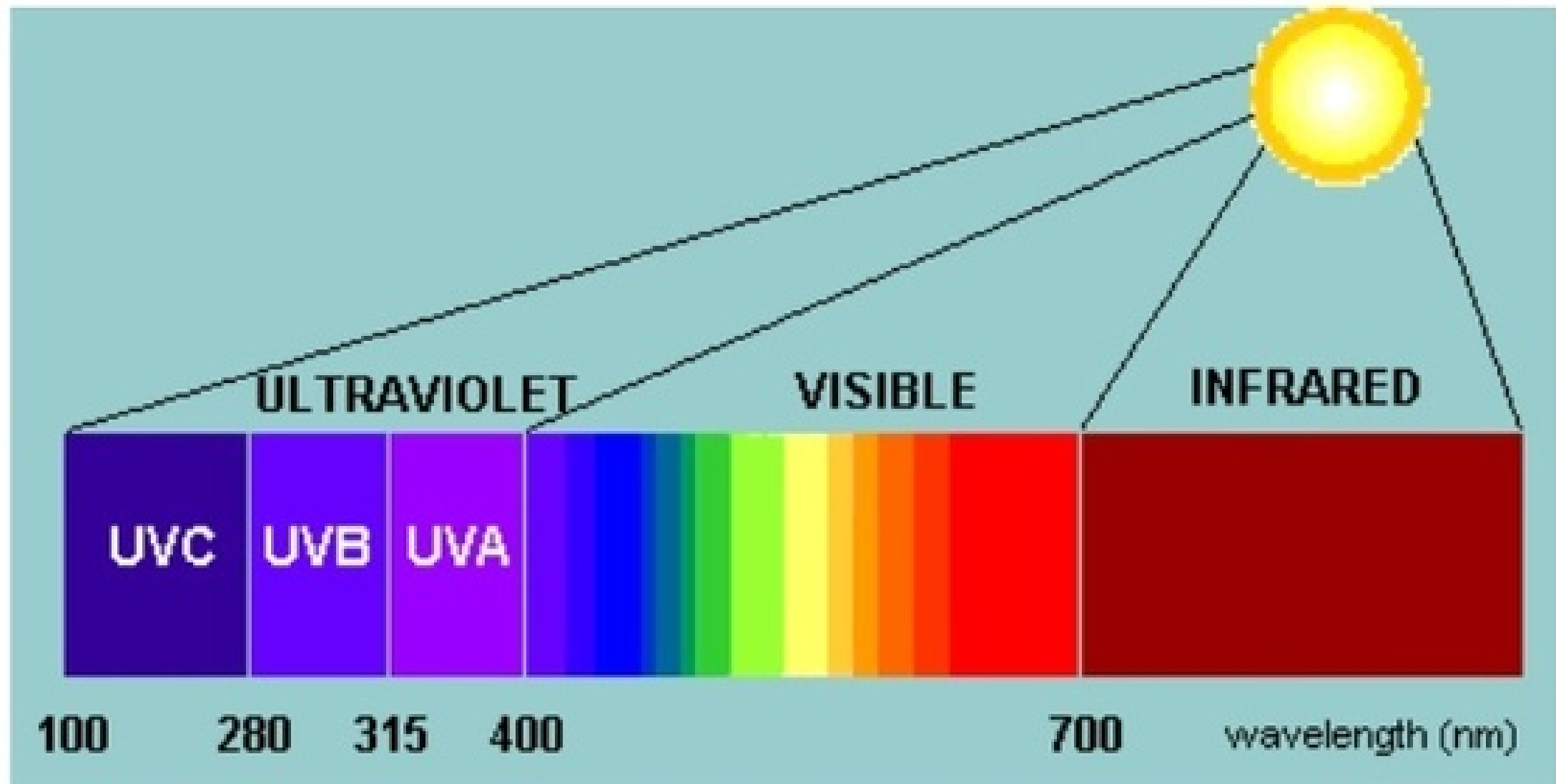
The Hydrogen Spectrum
Above: Emission
Below: Absorption

- Of the red shift we know from the position of the hydrogen absorption and emission lines

Star spectra contain information of the elements in it



- **UV radiation includes UV-A**, the least dangerous form of UV radiation, with a wavelength range between 315nm to 400nm, **UV-B** with a wavelength range between 280nm to 315nm, and **UV-C** which is the most dangerous between 100nm to 280nm. UV-C is unable to reach Earth's surface due to stratospheric ozone's ability to absorb it. (Last, 2006)



Sun

UV-a

UV-b

UV-c

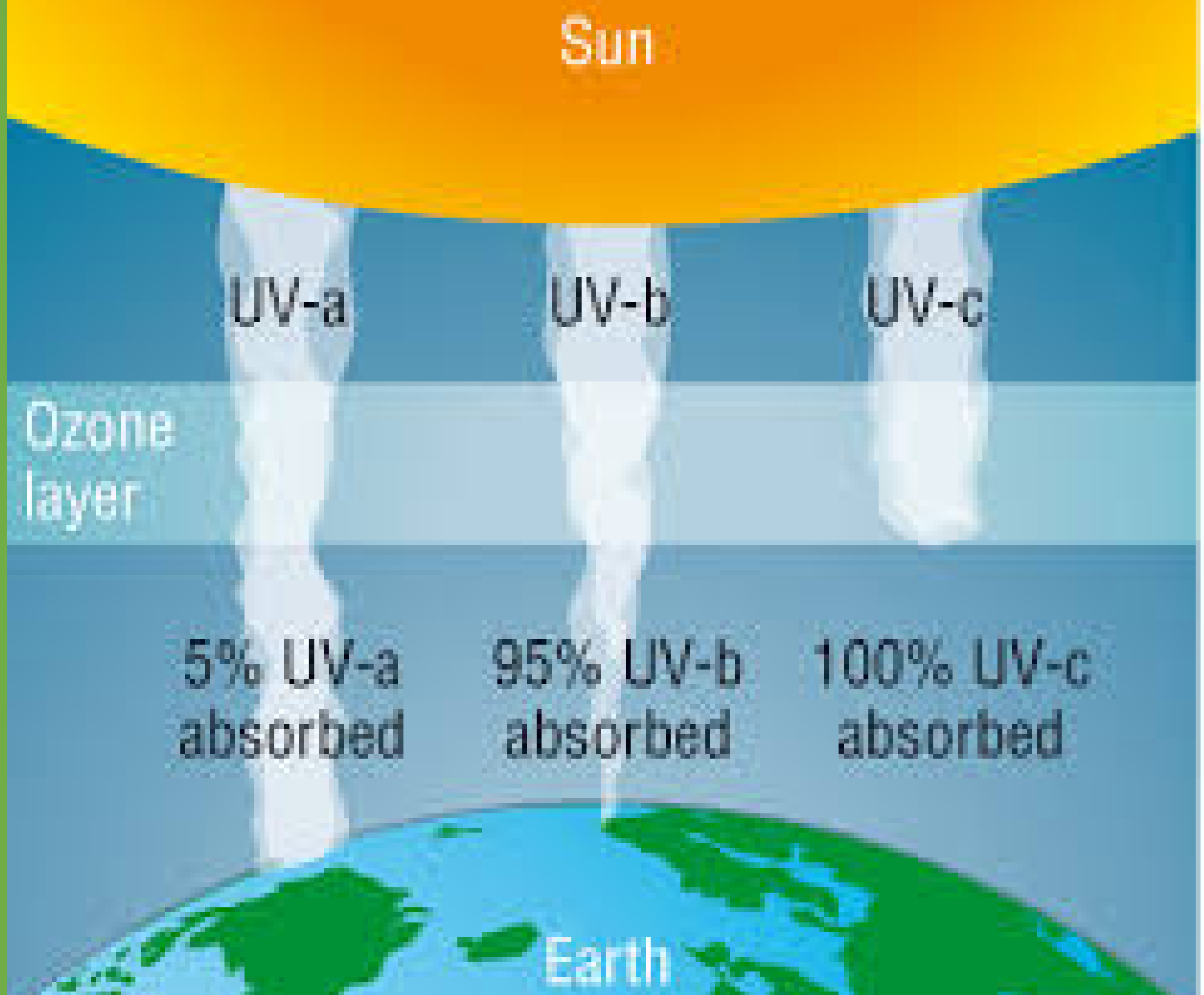
Ozone layer

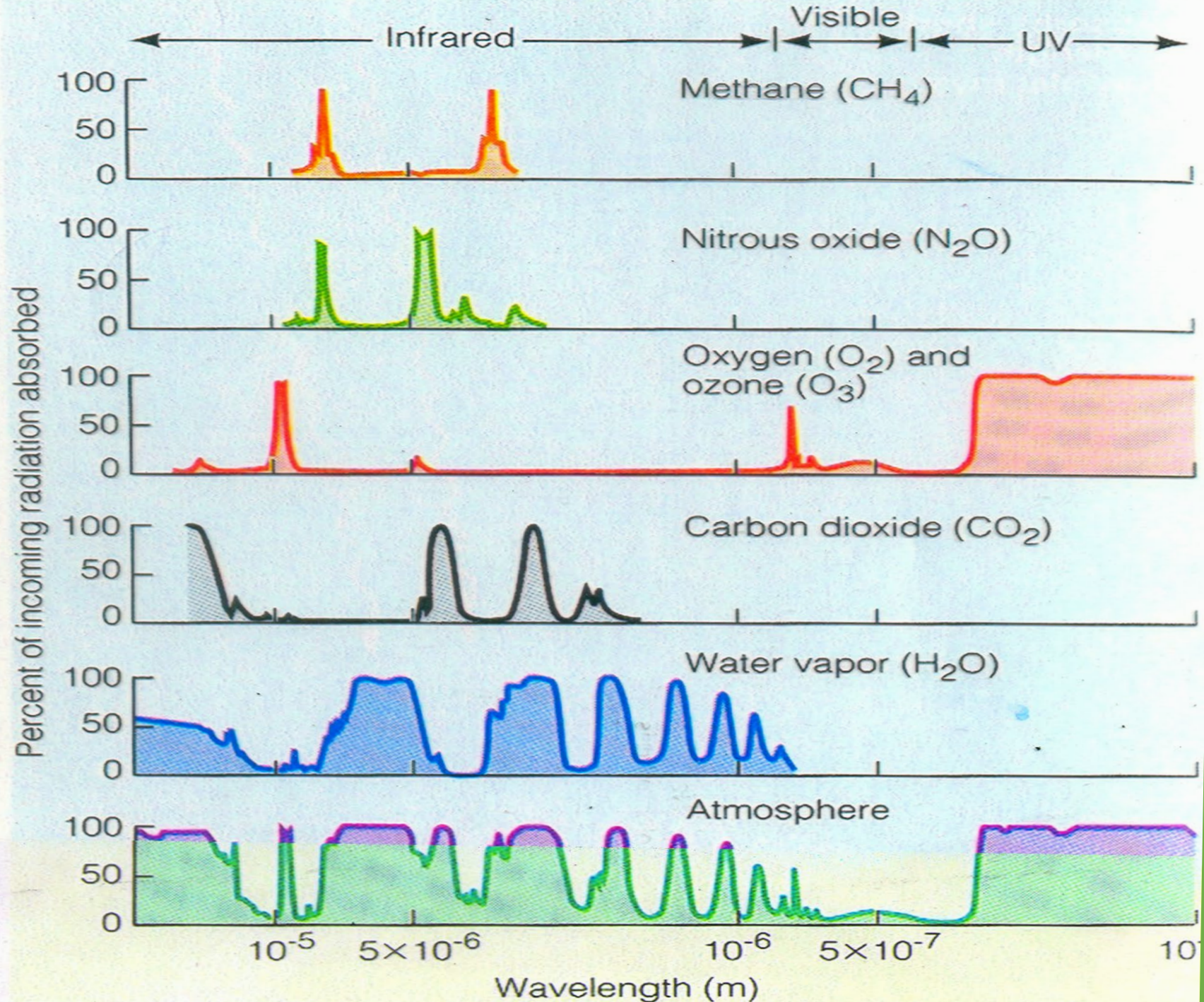
5% UV-a absorbed

95% UV-b absorbed

100% UV-c absorbed

Earth





Increasing flux ↑

Solar spectrum outside atmosphere

Solar spectrum at sea level

Visible

Infrared

Ultraviolet

Wavelength (m)

2×10^{-7}

6×10^{-7}

1×10^{-6}

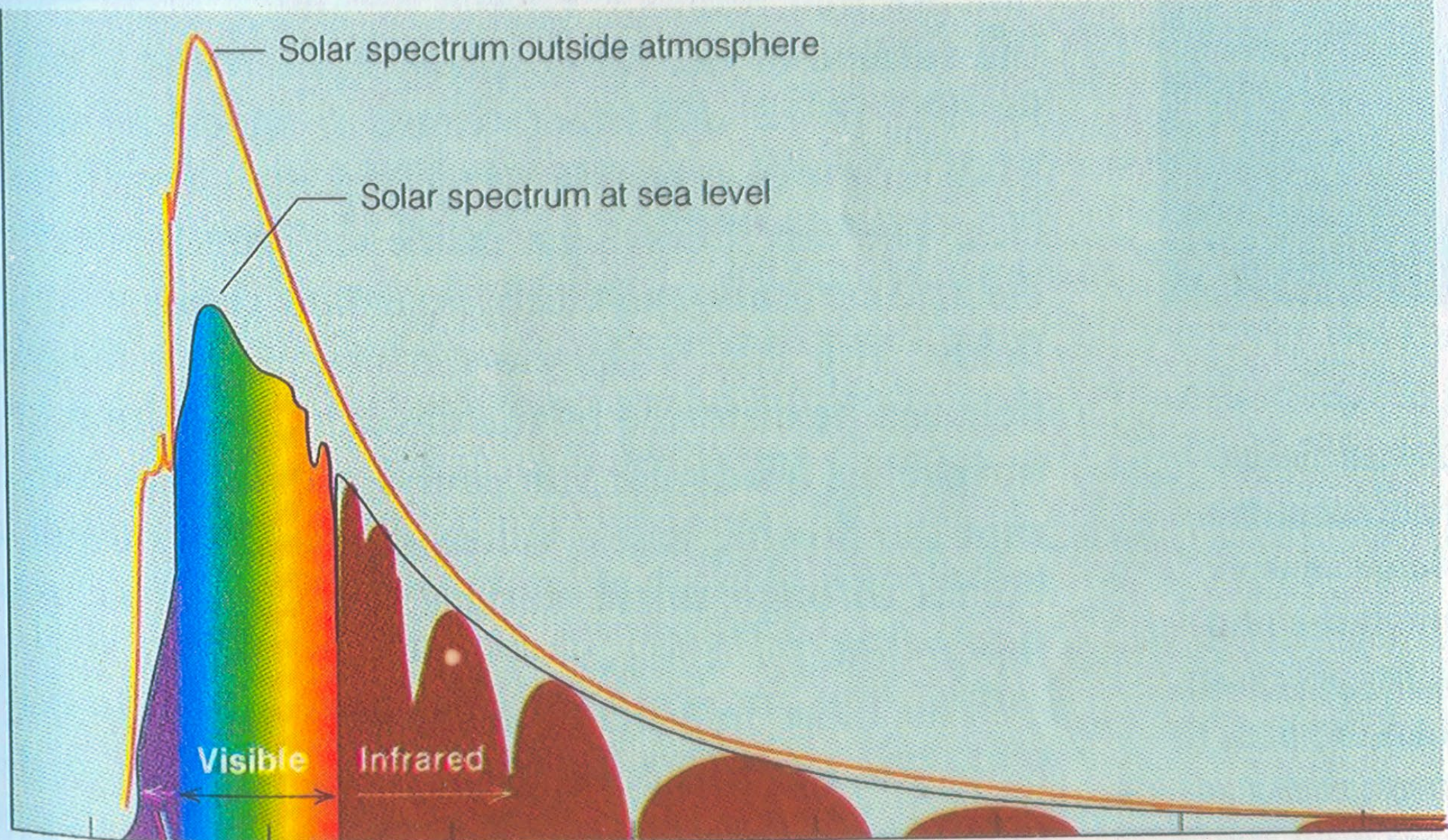
1.4×10^{-6}

1.8×10^{-6}

2.2×10^{-6}

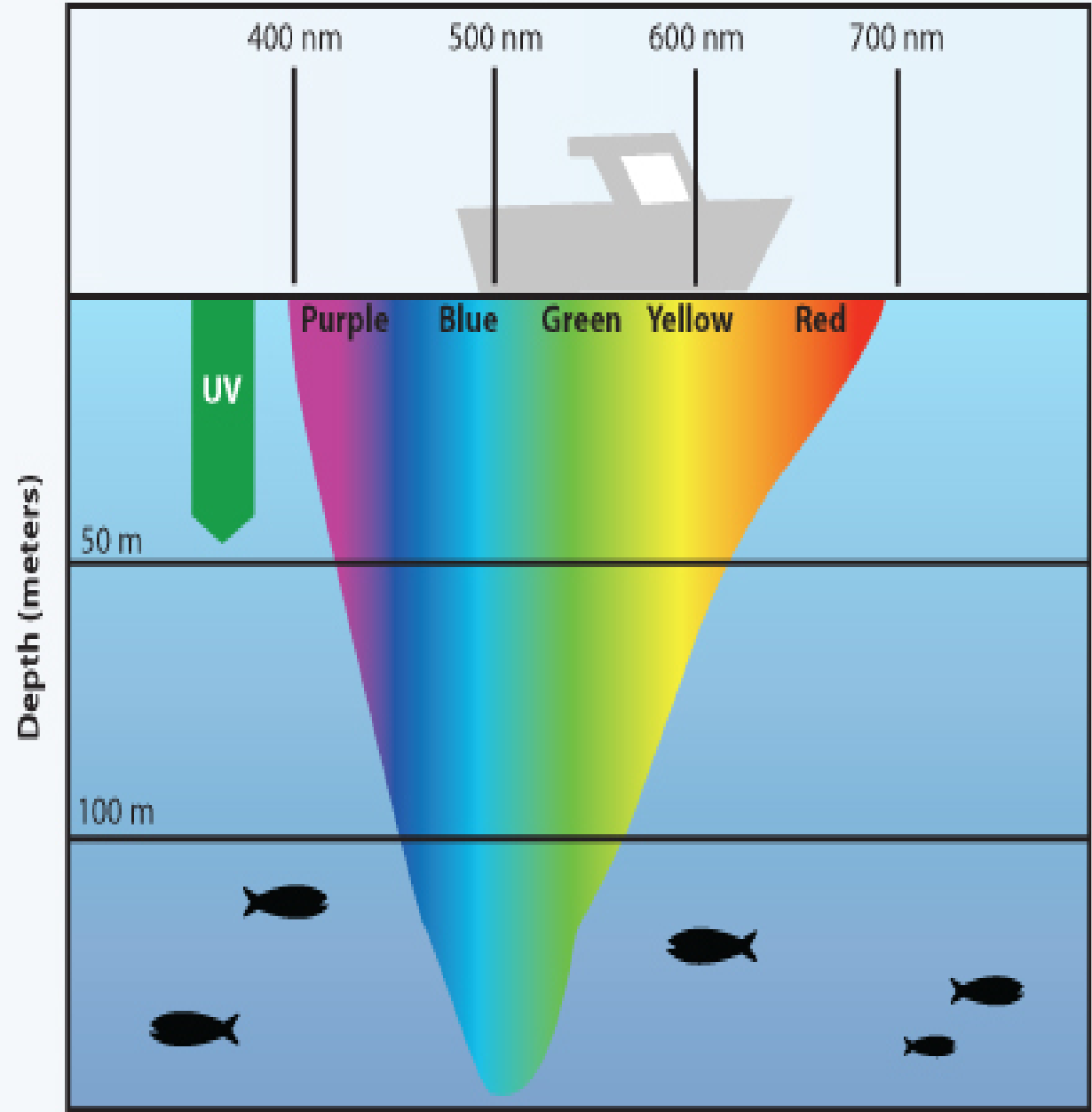
2.6×10^{-6}

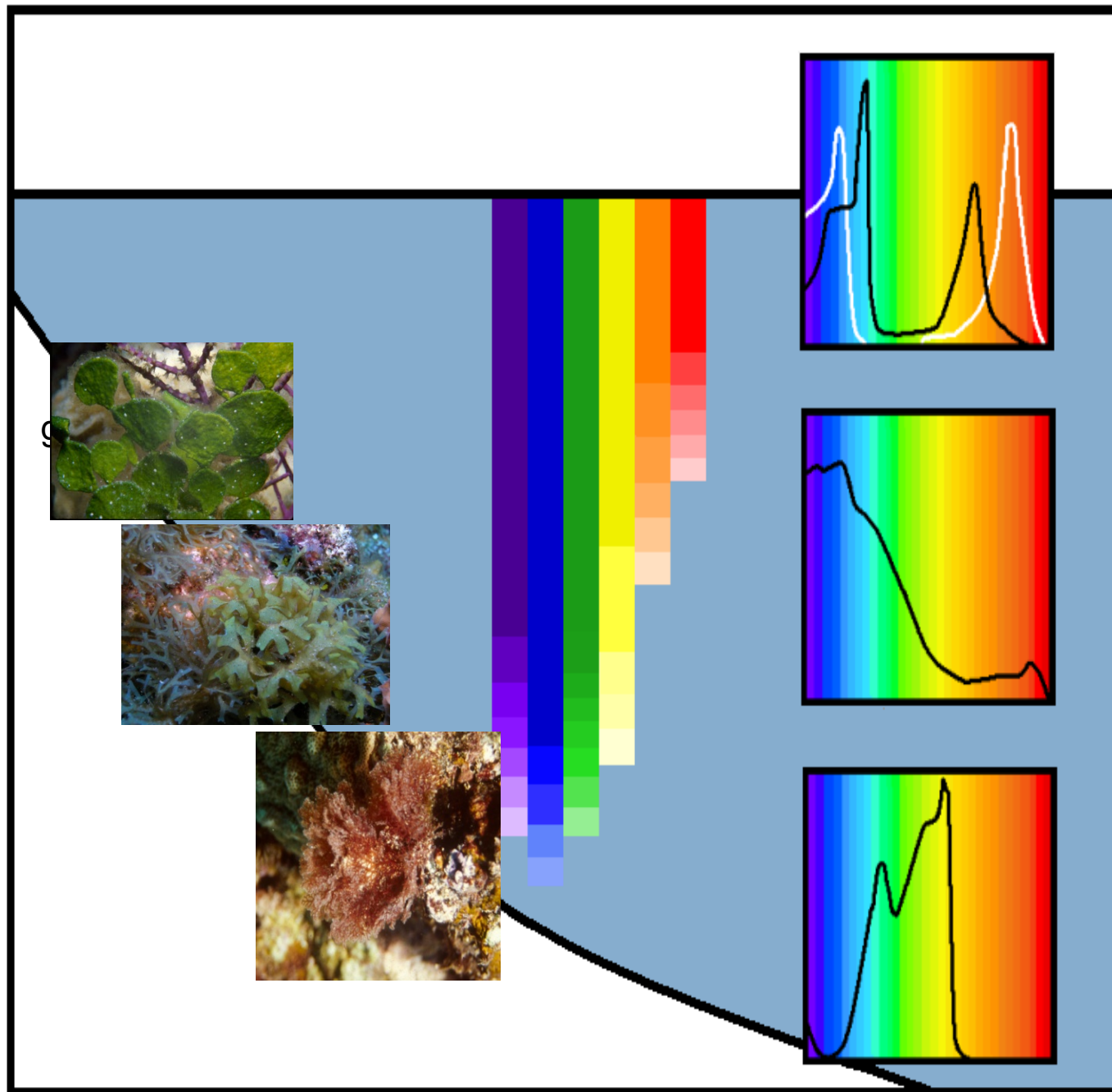
3×10^{-6}



Light Penetration in Lake Superior (Open Water, Clear Day)

Wavelength (nanometers)



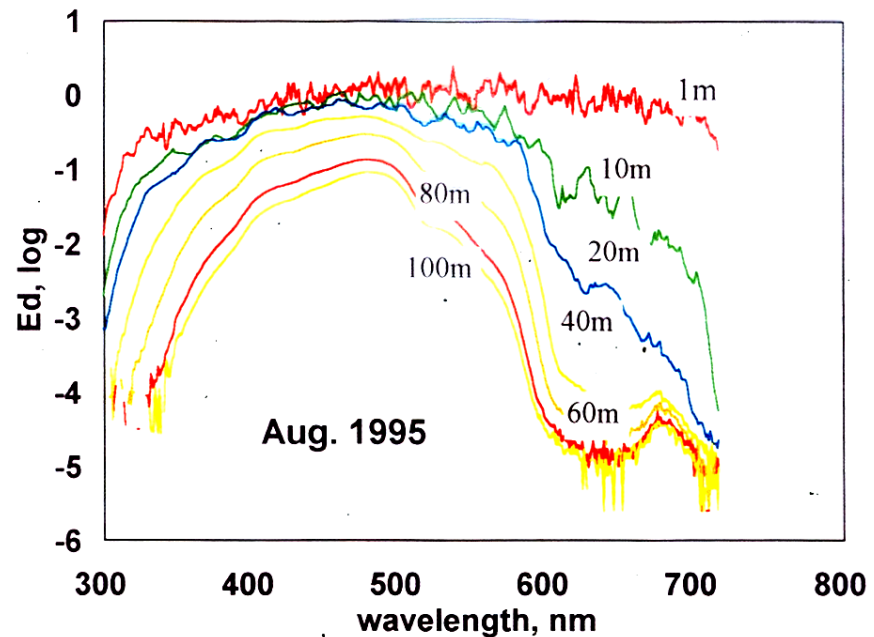
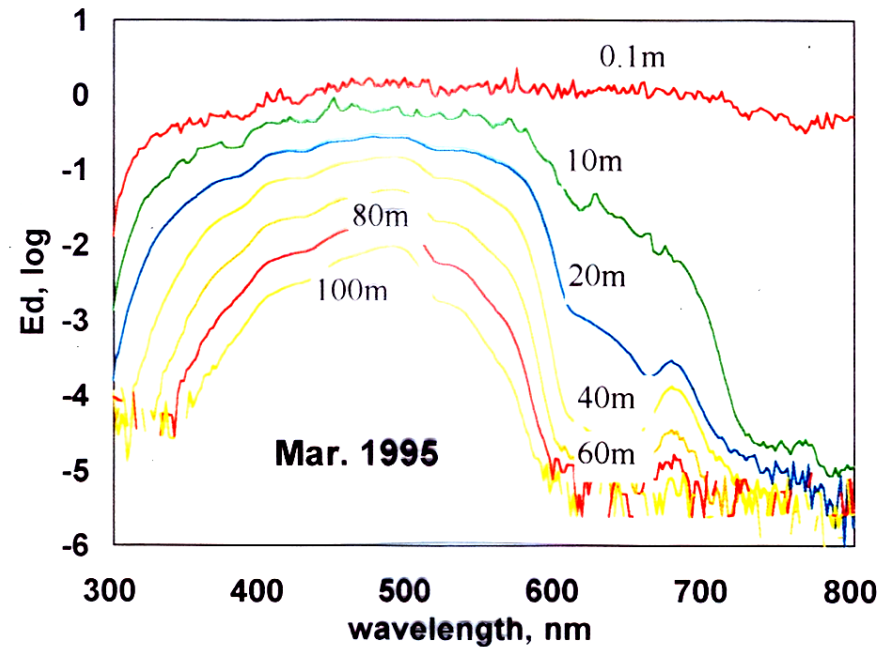


All algae have chlorophyll a (white line in top graph)

Green algae also have chlorophyll b and are generally limited to shallow turbid water.

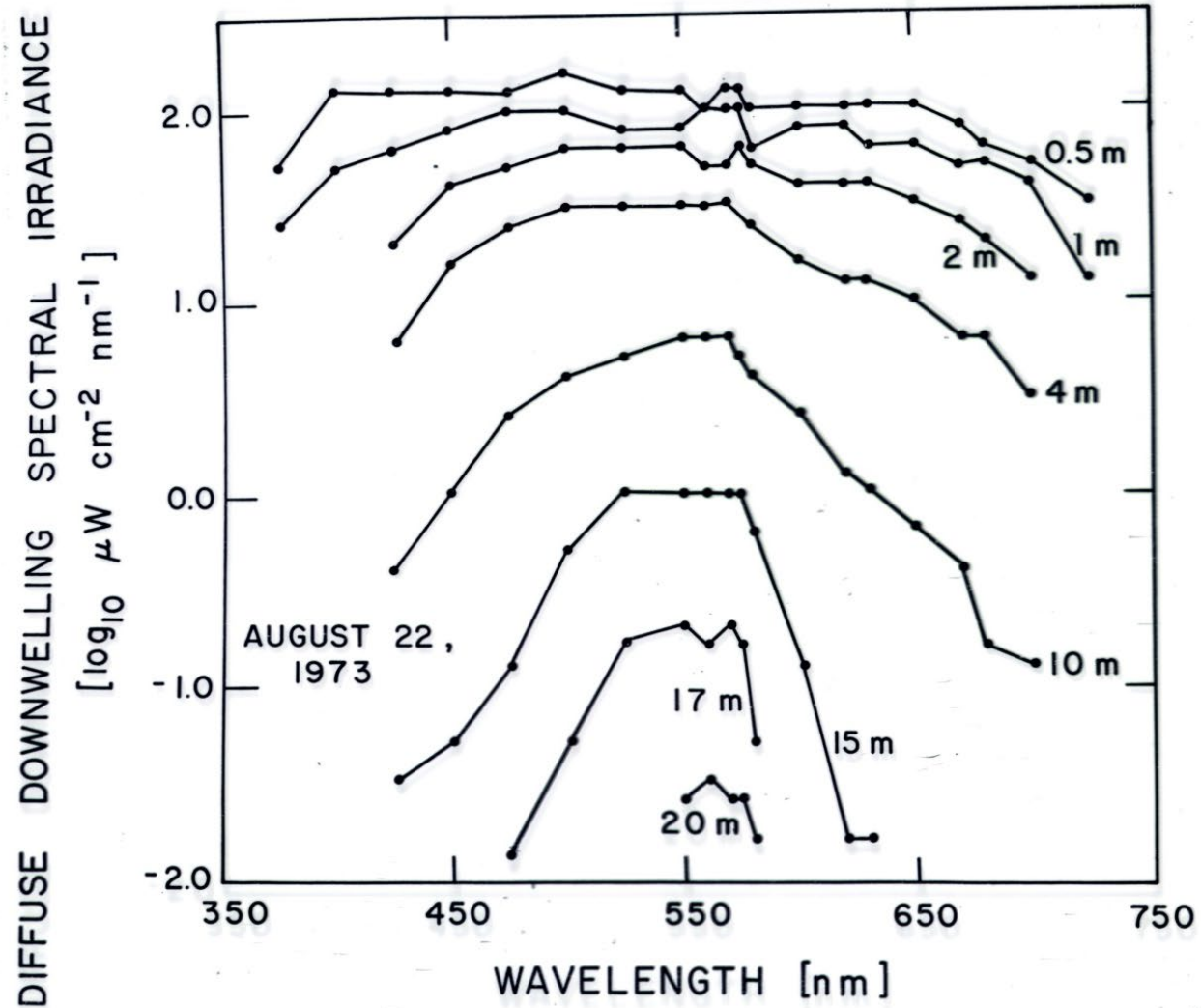
Brown algae have auxiliary pigments such as fucoxanthin that expand absorption into the green wavelengths and allow them to proliferate at intermediate depths

Red algae can survive at the deepest depths as auxiliary pigments such as phycoerythrin absorb blue and green light

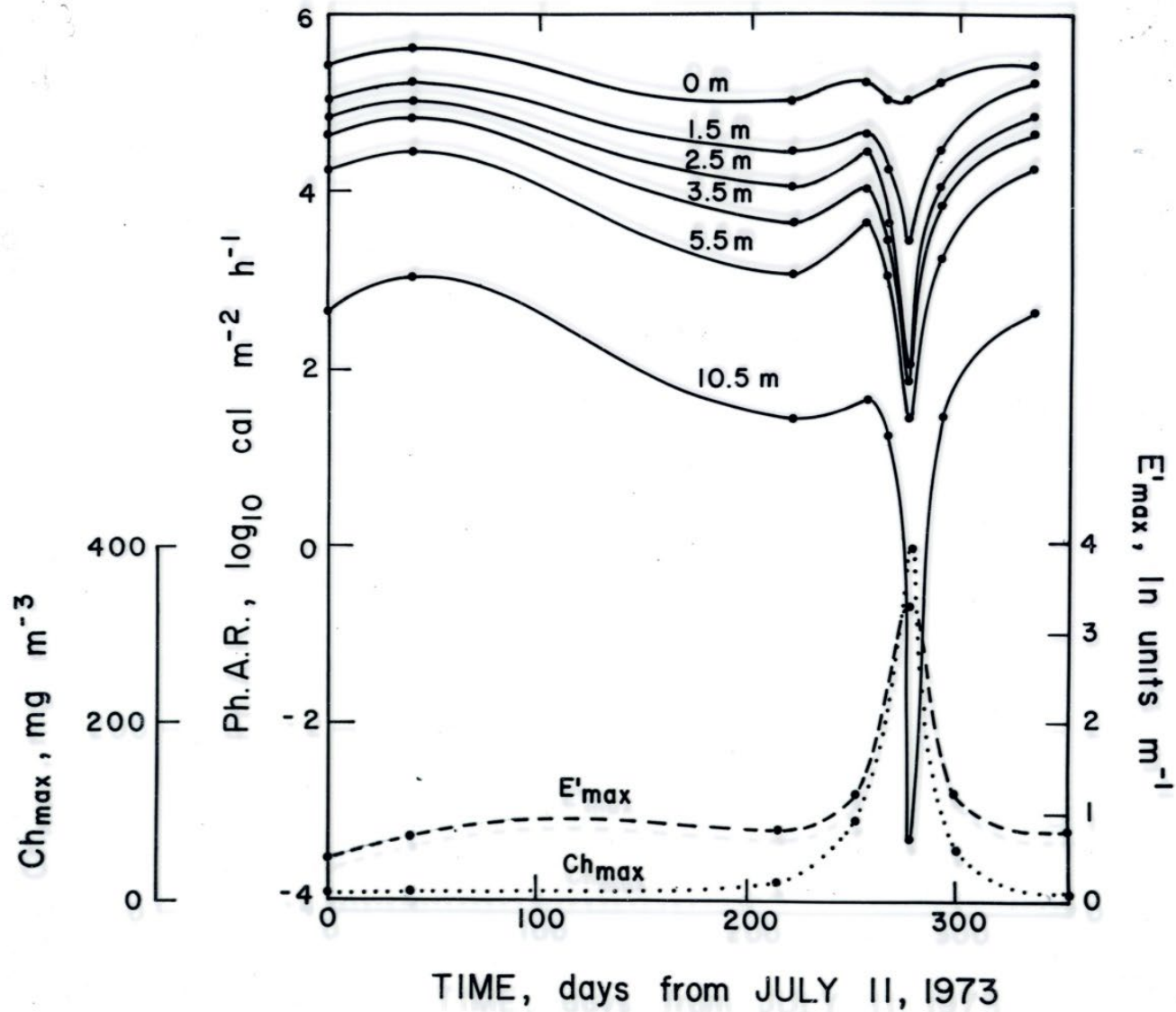


Spectral distribution of underwater light in spring and summer in the oligotrophic Gulf of Eilat (Red Sea)

Spectral distribution of underwater light in spring and summer in the eutrophic Lake Kinneret

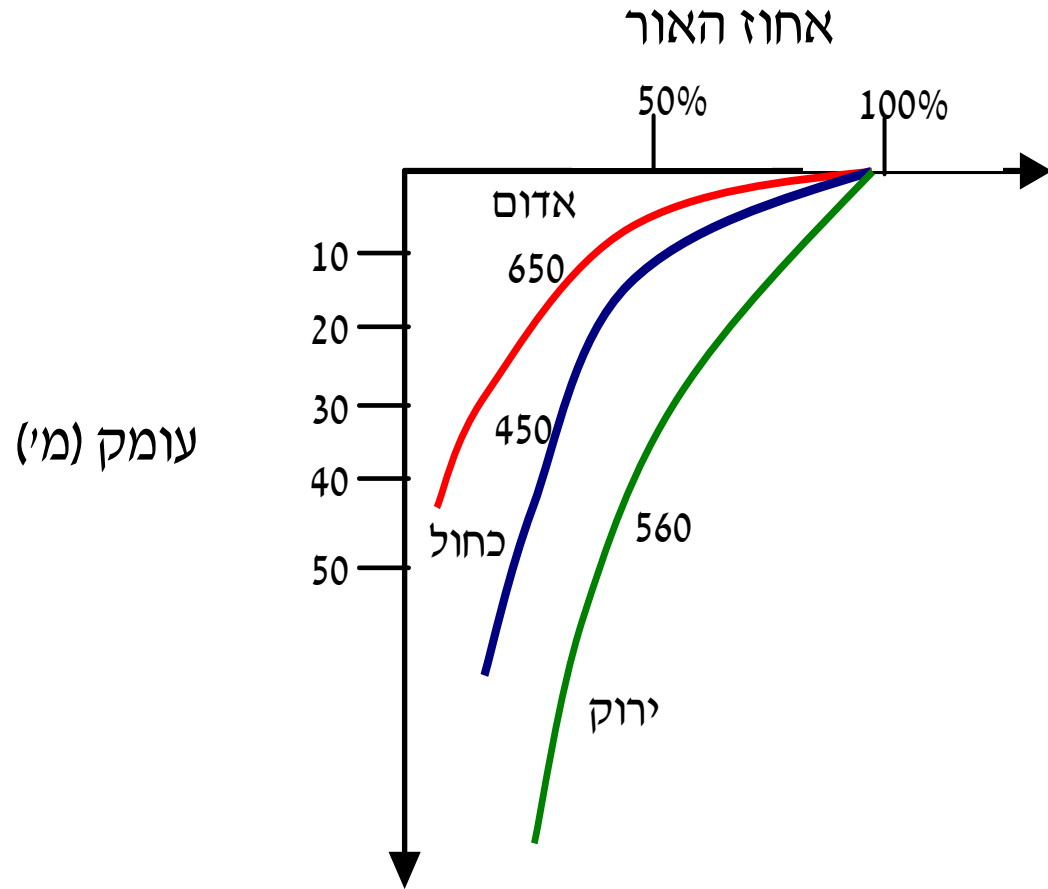


Light and chlorophyll in Lake Kinneret

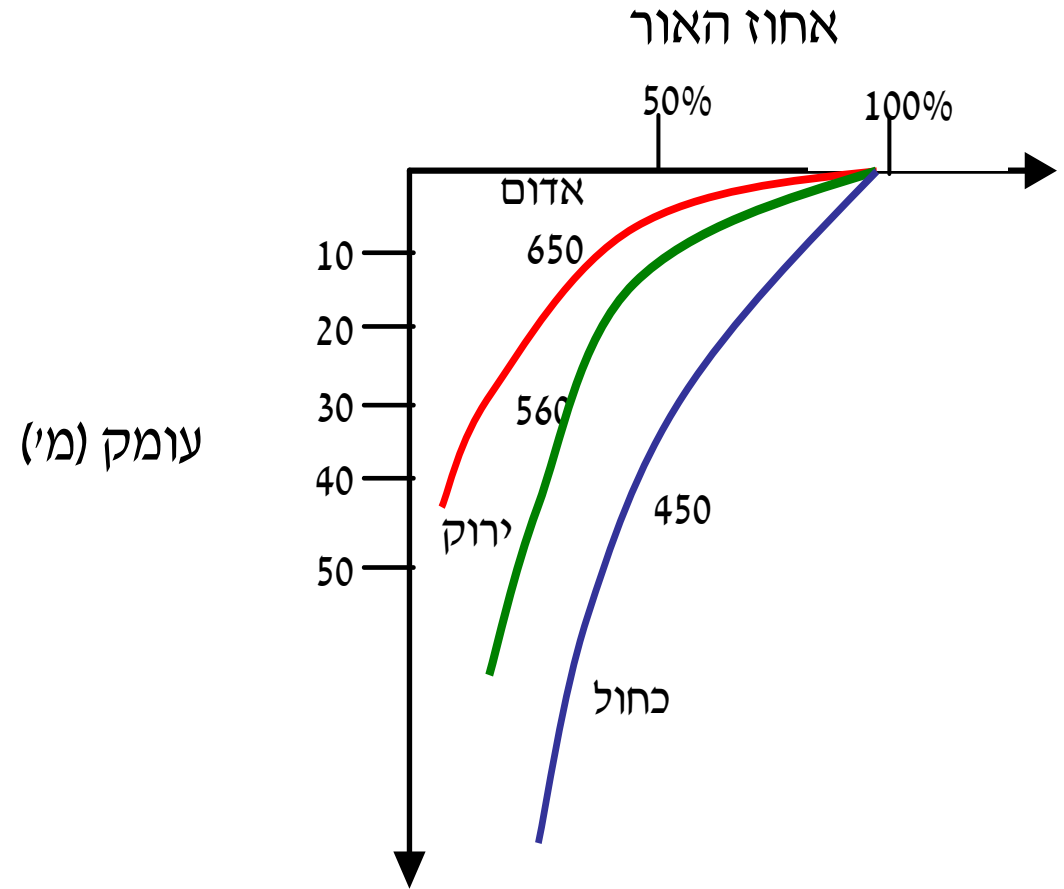


Each wavelength is absorbed differently with depth

אאוטרופים



אוליגוטרופים



Light in water

- Light is absorbed with depth:

- $E_z = E_0 e^{-k_d z}$

- E_z is the light intensity at z meters
- E_0 is the light intensity at 0 meters
- k_d is the light attenuation coefficient

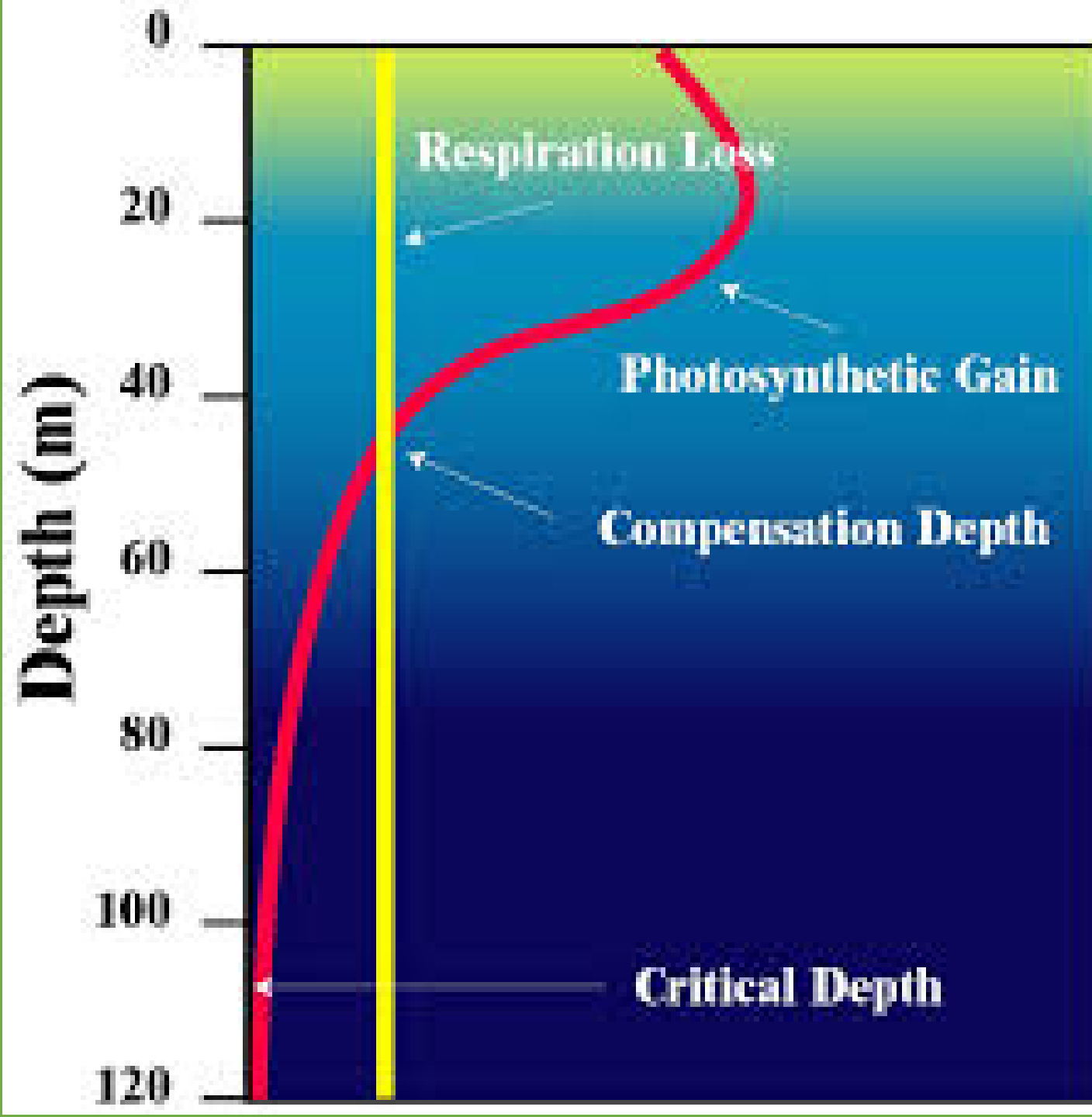
- $E_z/E_0 = e^{-k_d z}$

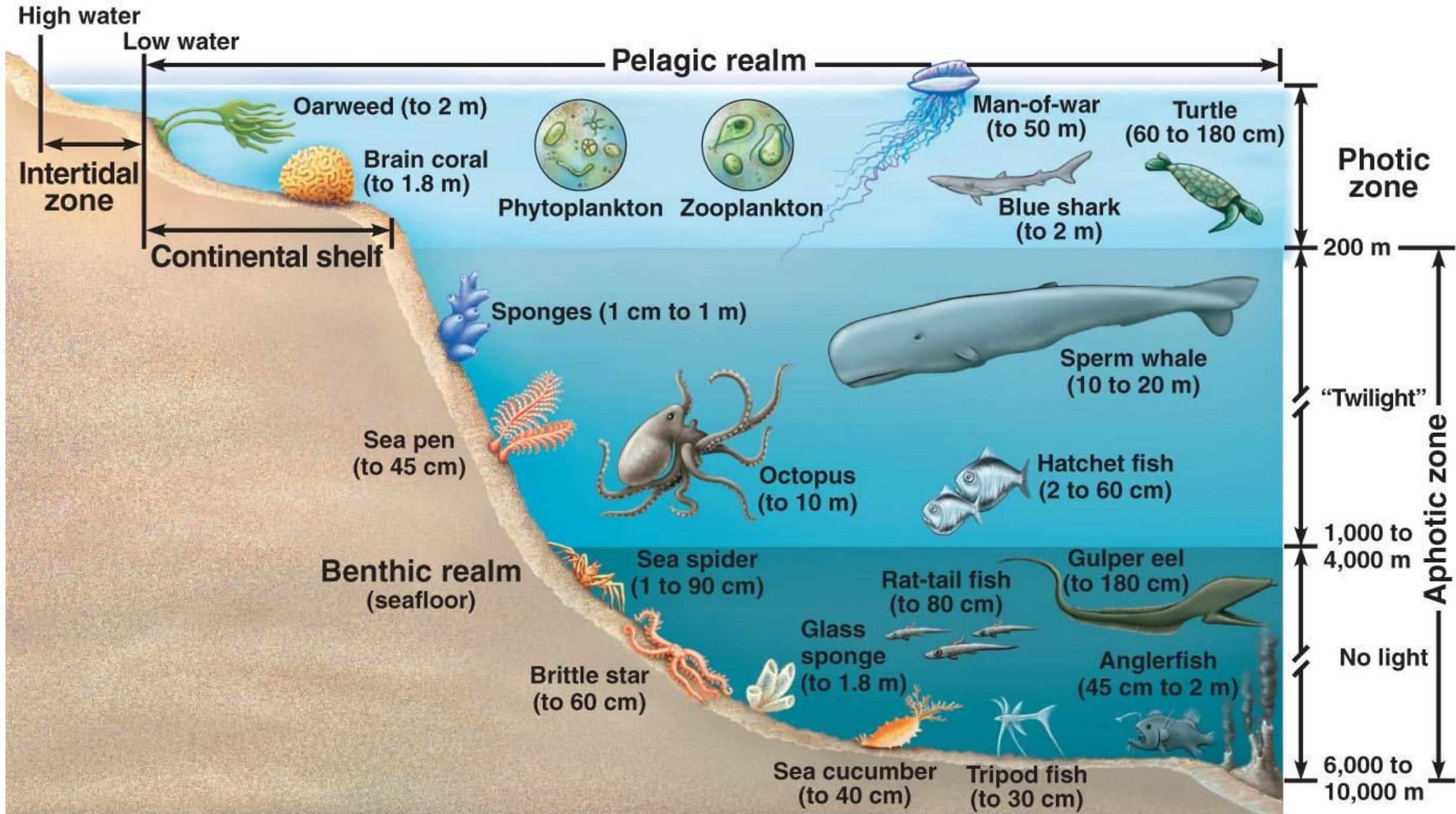
- $\ln(E_z/E_0) = -k_d z$

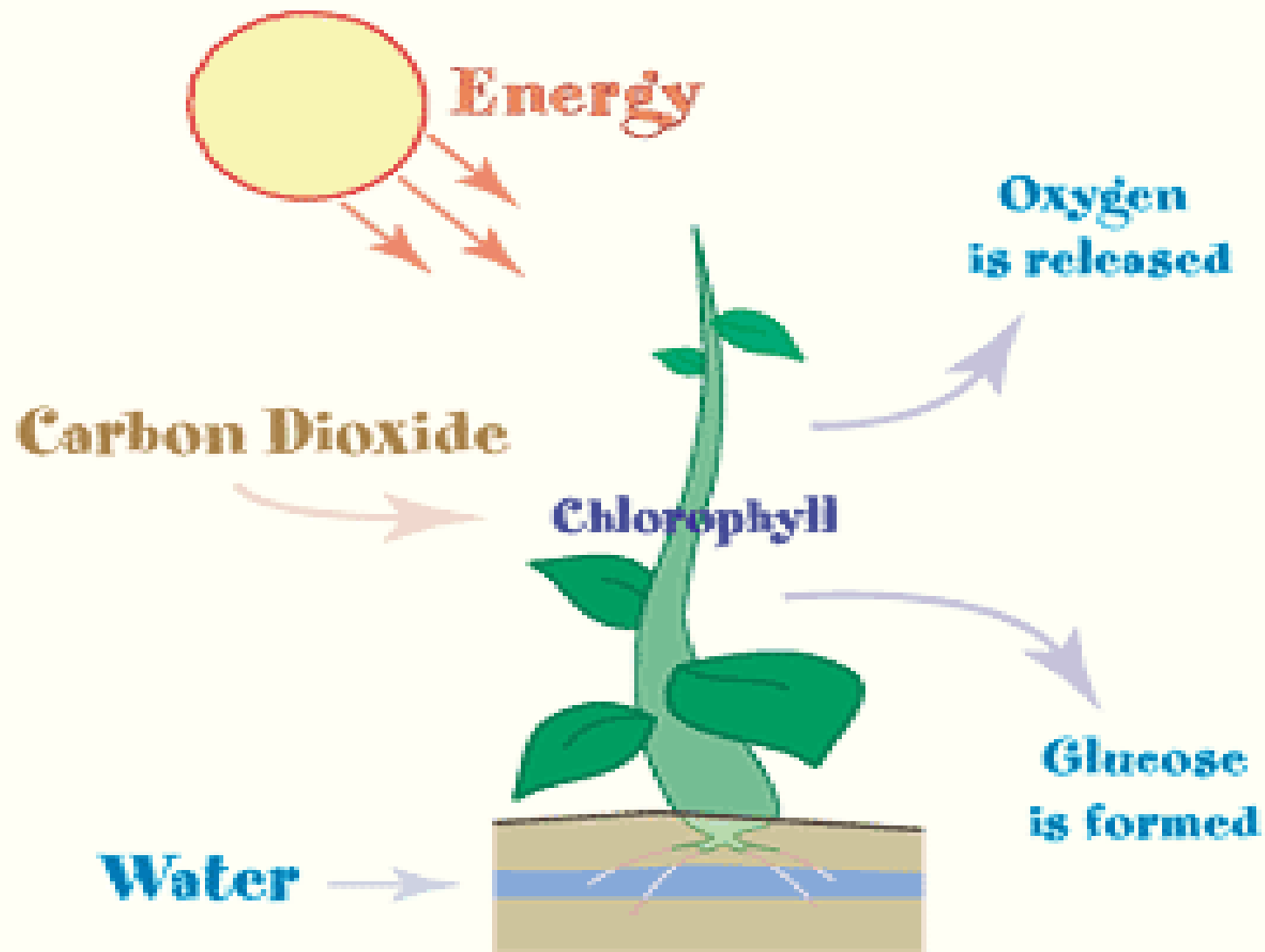
- $k_d = -\ln(E_z/E_0)/z$ [m]

The euphotic depth z_{eu}

- z_{eu} Is the depth at which light is 1% of its surface intensity
 - $E_z = 1\% E_0$
 - $E_z = 0.01 E_0$
 - $0.01 E_0 = E_0 e^{-k_d z_{eu}}$
 - $0.01 = e^{-k_d z_{eu}}$
 - $\ln 0.01 = -k_d z_{eu}$
 - $z_{eu} [m] = -\ln 0.01 / k_d$







Photosynthesis

Photosynthesis



- $P_G - R = P_N$

- $P_G = P_N + R$

הפוטוסינתזה



- $P_G - R = P_N$

- נשימה R

- פוטוסינתזה ברוטו (כולל הפסדי נשימה)

- P_G

- P_N פוטוסינתזה נטו

Photosynthesis



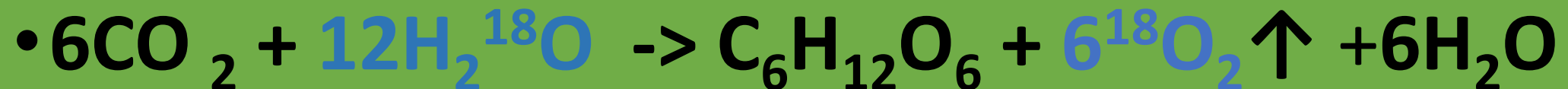
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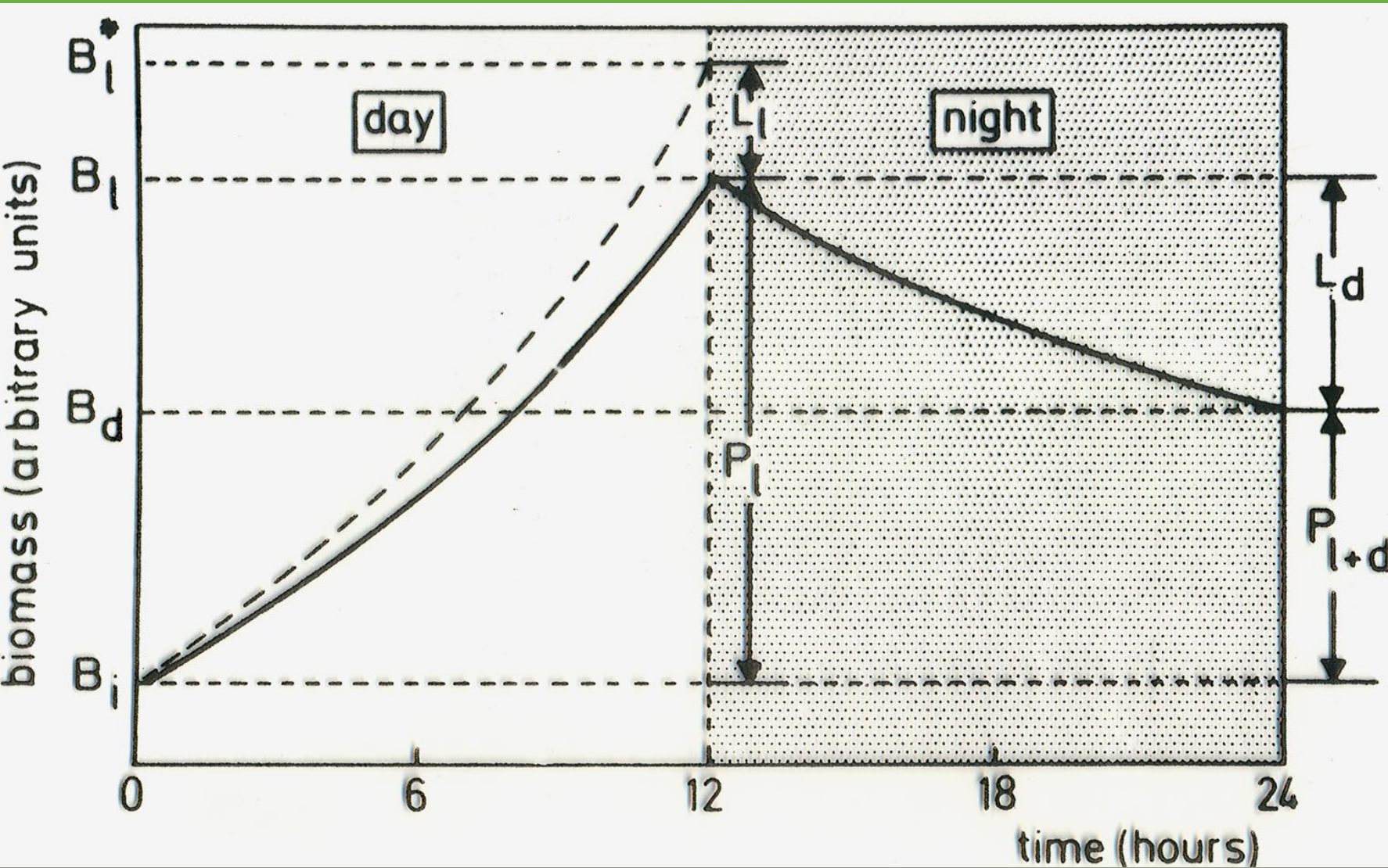


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Antarctic photosynthesis and daylength

Dashed line
represents gross
photosynthesis



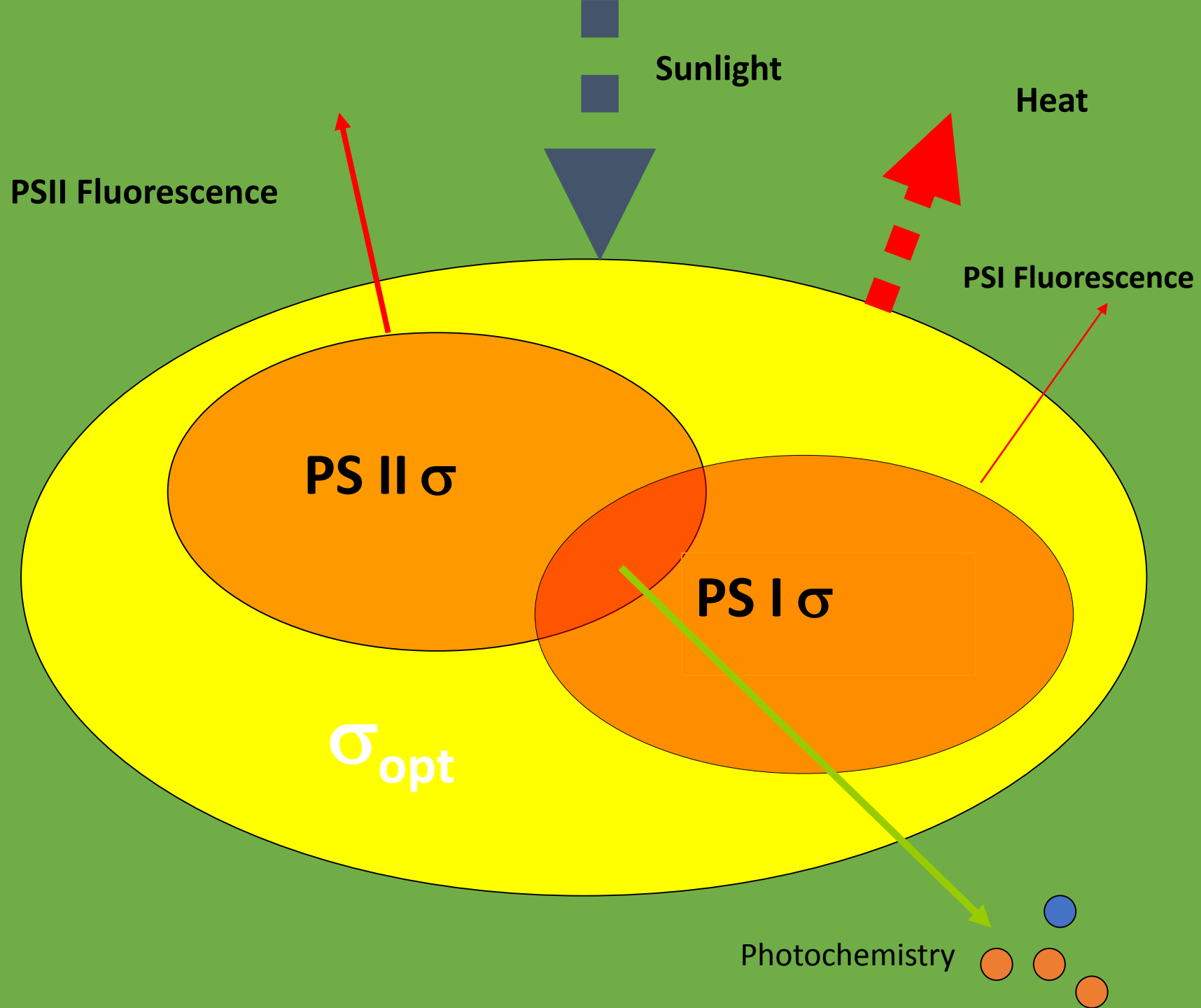
measuring photosynthesis



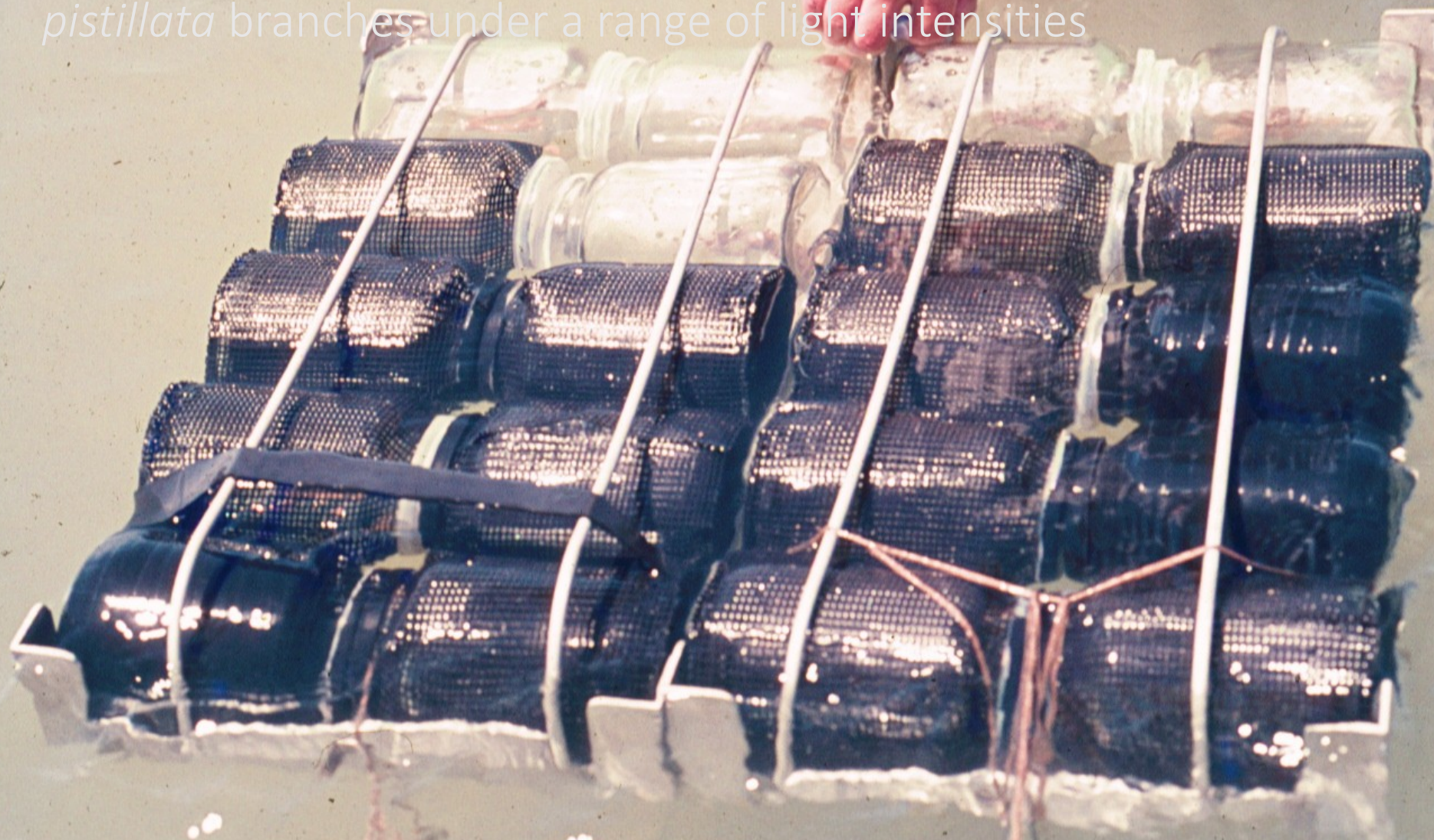
fluorescence

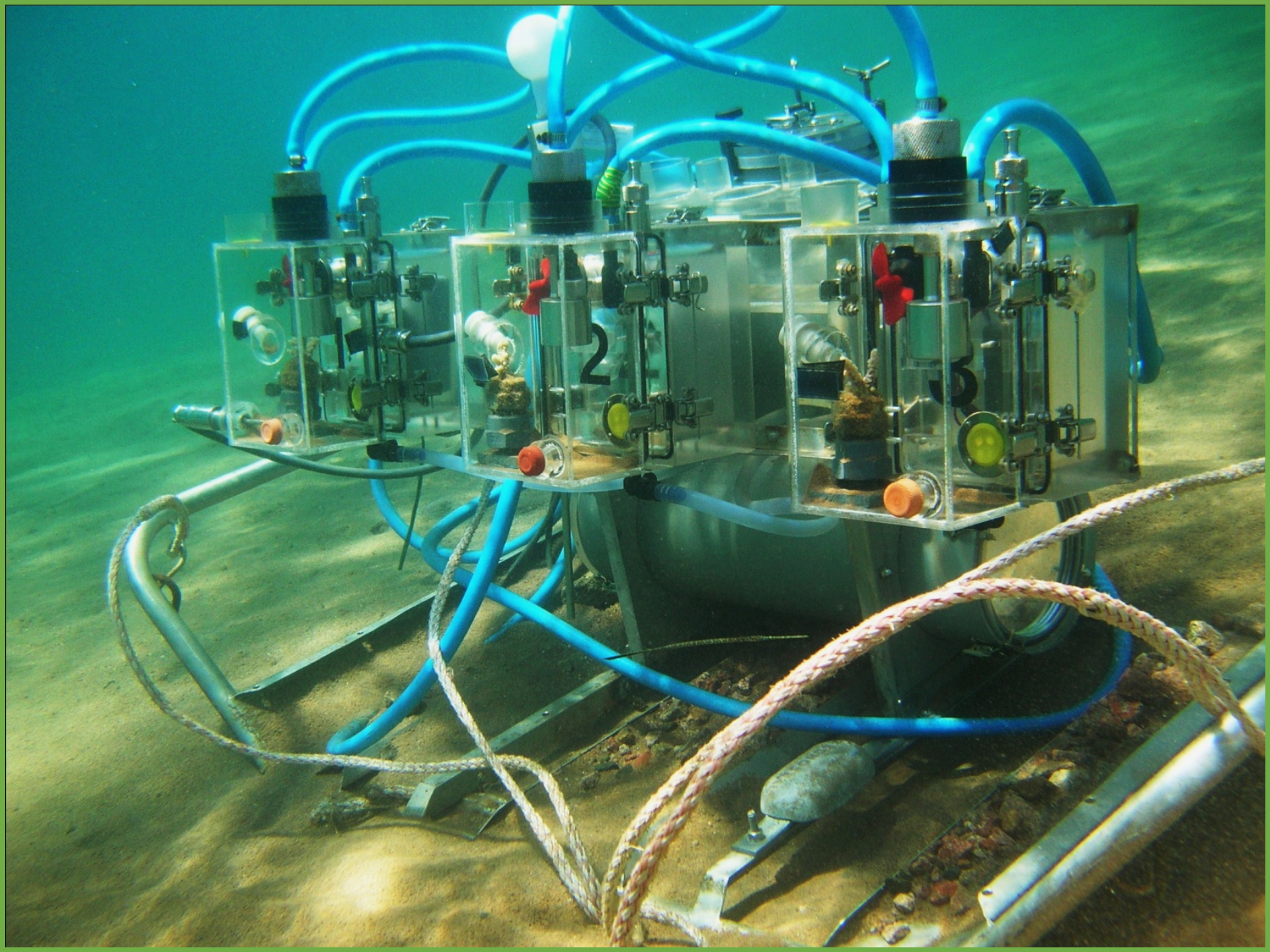
photoacoustics

$$P_G - R = P_N$$

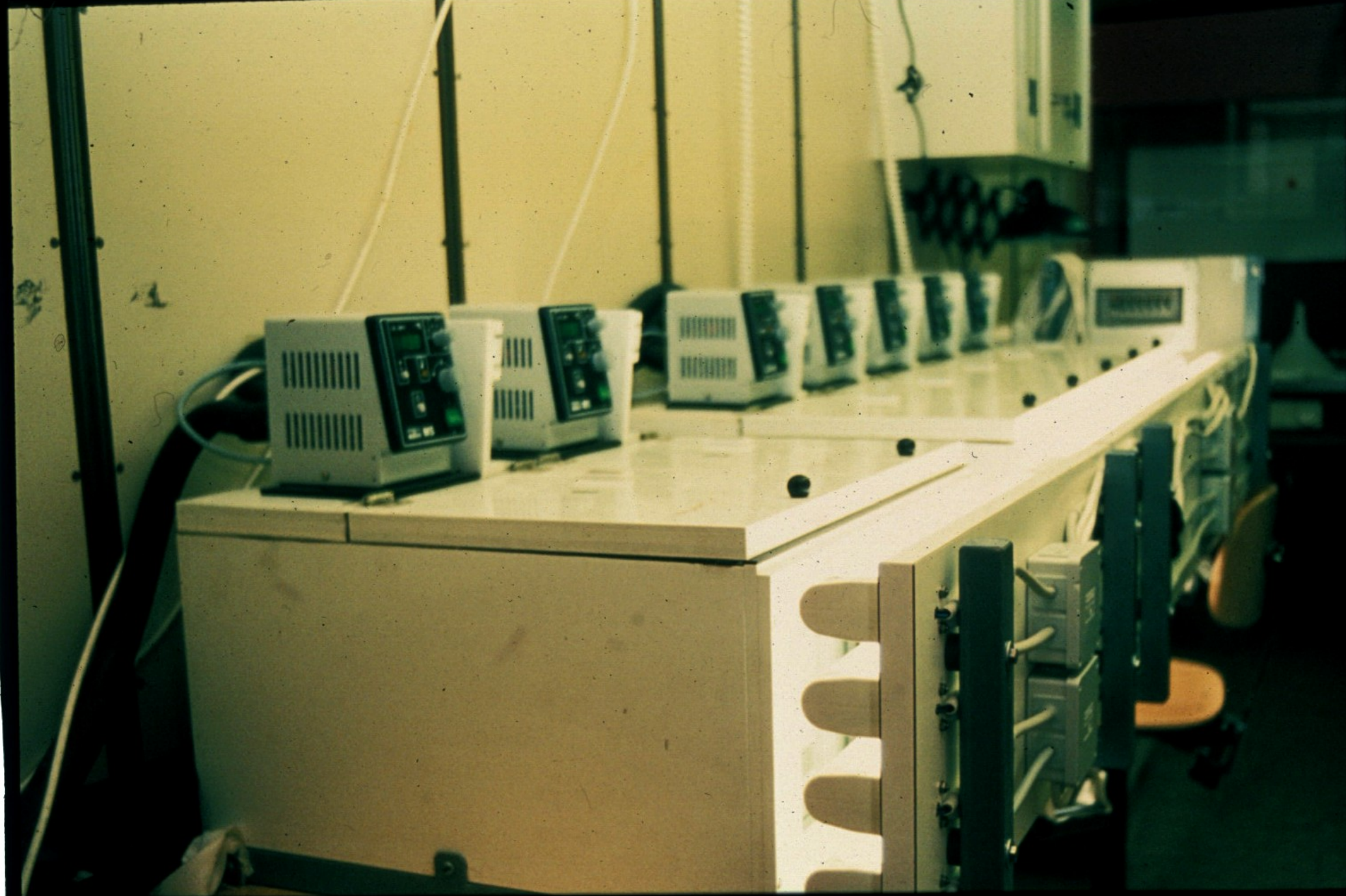


Radioactive carbon labelling of high and low light *Stylophora pistillata* branches under a range of light intensities











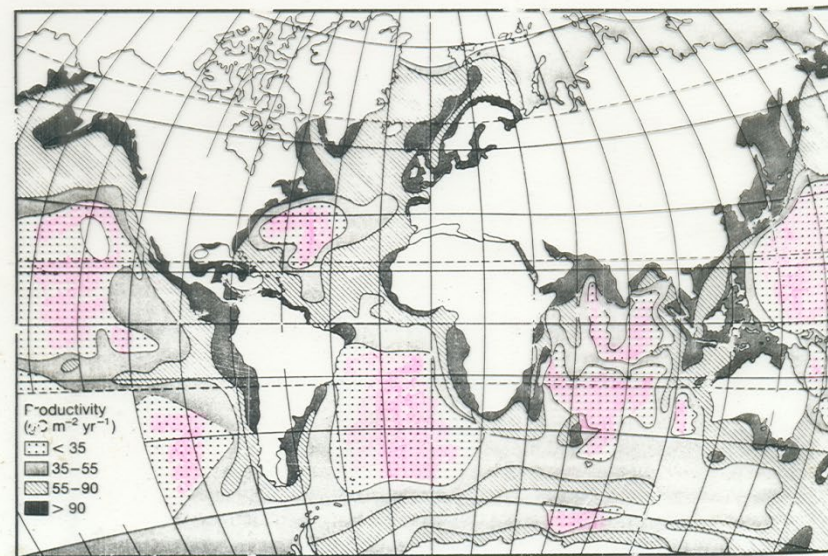
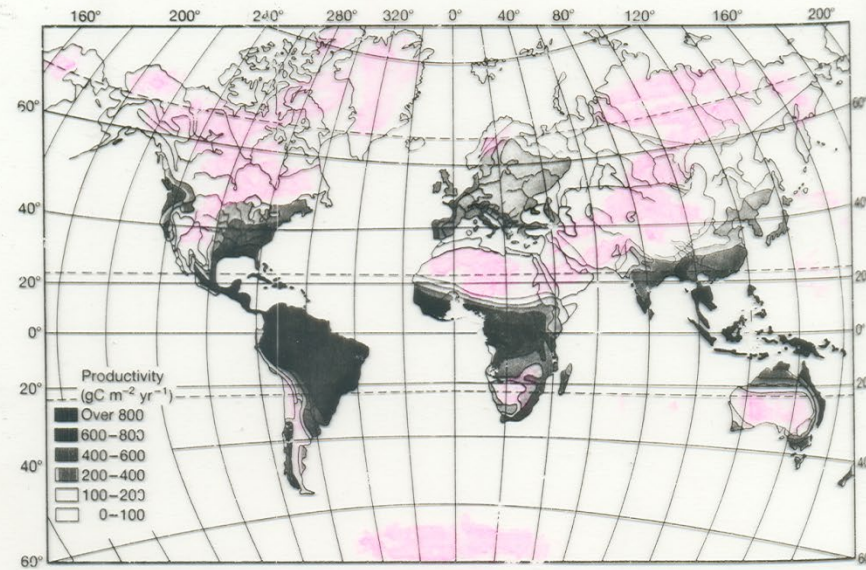


Figure 17.1. Top: Worldwide pattern of net primary productivity on land (from Reichle, 1970).

Bottom: Worldwide pattern of net primary productivity in the oceans (from Koblentz-Mishke *et al.*, 1970).

Table 18.1 Annual net primary productivity (NPP) and standing crop biomass estimates for contrasting communities of the world. (After Whittaker, 1975.)

Ecosystem type	Area (10^6 km^2)	NPP, per unit area (g m^{-2} or t km^{-2}) y^{-1}		World NPP (10^9 t)	Biomass per unit area (kg m^{-2})		World biomass (10^9 t)
		Normal range	Mean		Normal range	Mean	
Tropical rainforest	17.0	1000-3500	2200	37.4	6-80	45	765
Tropical seasonal forest	7.5	1000-2500	1600	12.0	6-60	35	260
Temperate evergreen forest	5.0	600-2500	1300	6.5	6-200	35	175
Temperate deciduous forest	7.0	600-2500	1200	8.4	6-60	30	210
Boreal forest	12.0	400-2000	800	9.6	6-40	20	240
Woodland and shrubland	8.5	250-1200	700	6.0	2-20	6	50
Savannah	15.0	200-2000	900	13.5	0.2-15	4	60
Temperate grassland	9.0	200-1500	600	5.4	0.2-5	1.6	14
Tundra and alpine	8.0	10-400	140	1.1	0.1-3	0.6	5
Desert and semi-desert shrub	18.0	10-250	90	1.6	0.1-4	0.7	13
Extreme desert, rock, sand and ice	24.0	0-10	3	0.07	0-0.2	0.02	0.5
Cultivated land	14.0	100-3500	650	9.1	0.4-12	1	14
Swamp and marsh	2.0	800-3500	2000	4.0	3-50	15	30
Lake and stream	2.0	100-1500	250	0.5	0-0.1	0.02	0.05
Total continental	149		773	115		12.3	1837
Open ocean	332.0	2-400	125	41.5	0-0.005	0.003	1.0
Upwelling zones	0.4	400-1000	500	0.2	0.005-0.1	0.02	0.008
Continental shelf	26.6	200-600	360	9.6	0.001-0.04	0.01	0.27
Algal beds and reefs	0.6	500-4000	2500	1.6	0.04-4	2	1.2
Estuaries	1.4	200-3500	1500	2.1	0.01-6	1	1.4
Total marine	361		152	55.0		0.01	3.9
Full total	510		333	170		3.6	1841

Primary productivity in tropical rainforest and upwelling marine zones

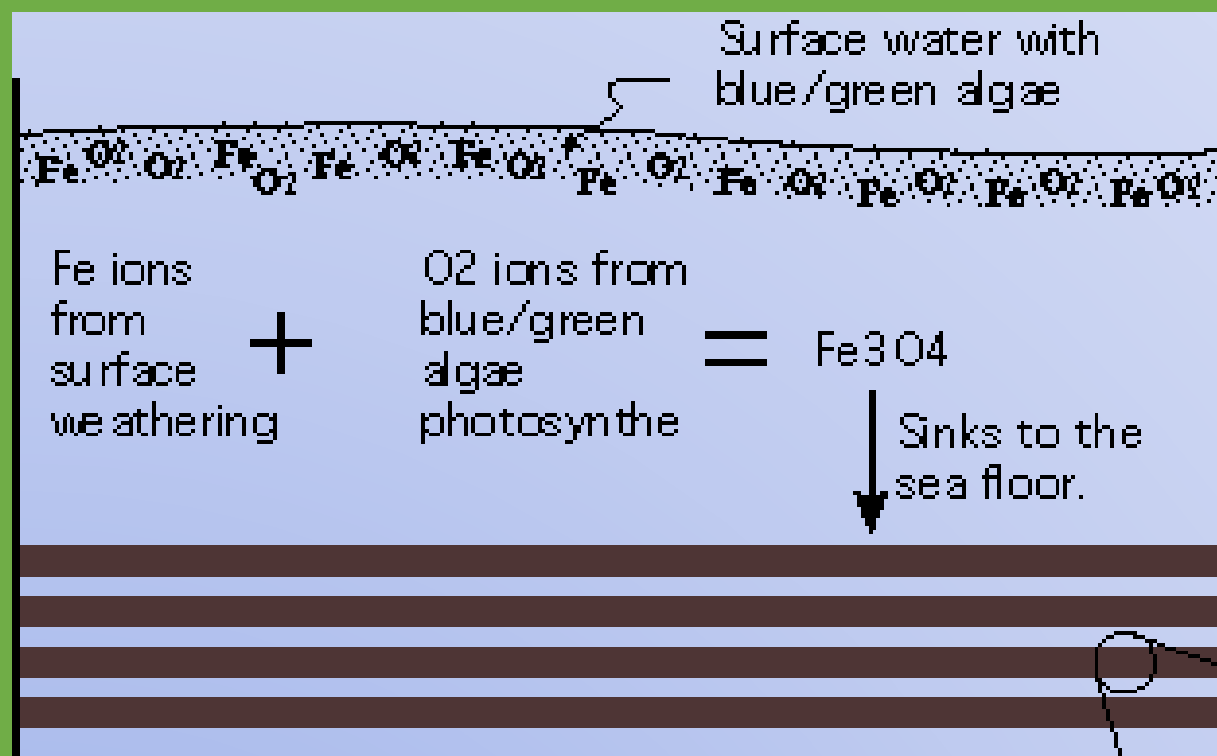
$$\text{gm}^{-2} \text{y}^{-1} / 45000 \text{ gm}^{-2} \sim 1/20 \text{ y}^{-1} \quad 2200$$

$$45000 \text{ gm}^{-2} / 2200 \text{ gm}^{-2} \text{ y}^{-1} \sim 20 \text{ y}$$

$$500 \text{ gm}^{-2} \text{ y}^{-1} / 20 \text{ gm}^{-2} \sim 25 \text{ y}^{-1}$$

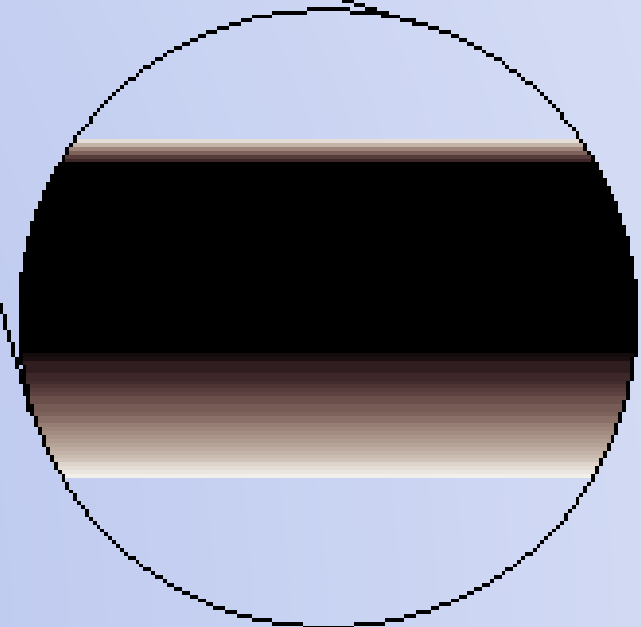
$$20 \text{ gm}^{-2} / 500 \text{ gm}^{-2} \text{ y}^{-1} \sim 1/25 \text{ y}$$

$$20:1/25=500!!!$$



After combining the Fe and O₂ ions into Magnetite (Fe₃O₄), the mineral grains sink to the sea floor, where they accumulate into iron-rich and iron-poor layers.

In an ideal setting, you would expect the magnetite-rich layers to exhibit a reversed graded bedding. Looking from the bottom up, this would involve a slow transition into the magnetite-rich layers, representing slowly increasing O₂ levels in the upper sea water in response to the increasing population of blue/green algae. The upper contact of each magnetite-rich layer would be relatively abrupt, reflecting the sudden extinction of the population due to O₂ poisoning, and the resulting loss of available O₂ in the water to combine with the iron ions.



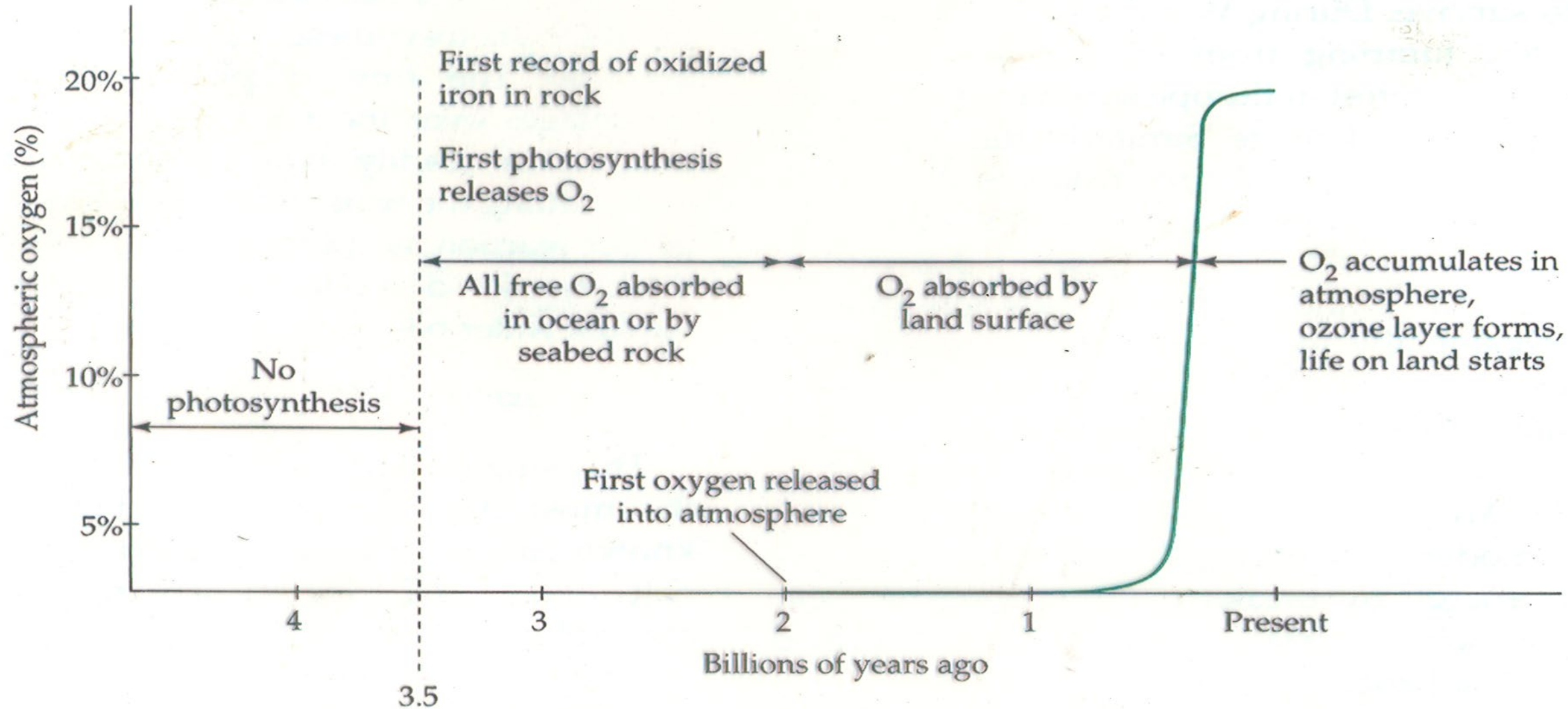
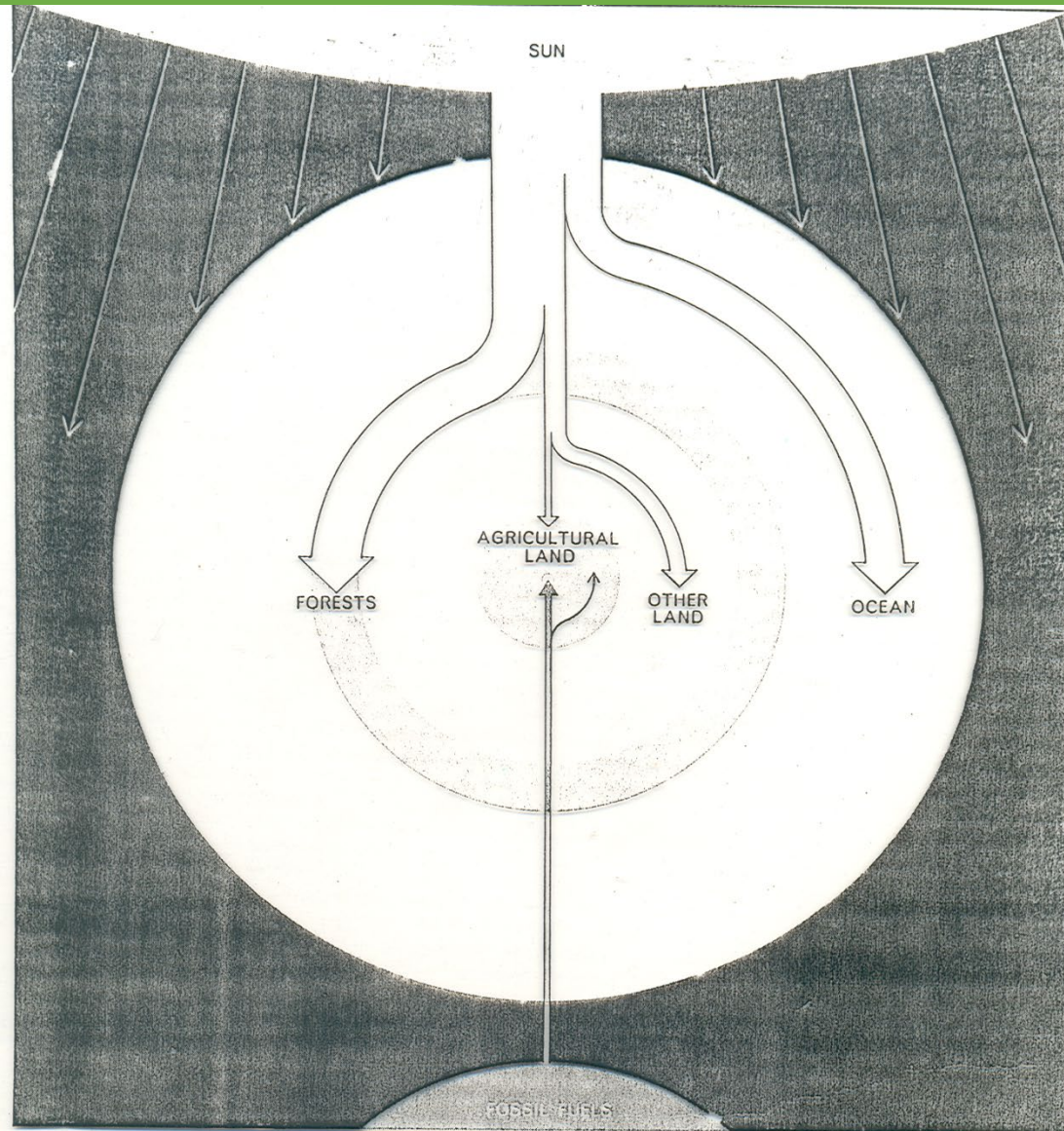


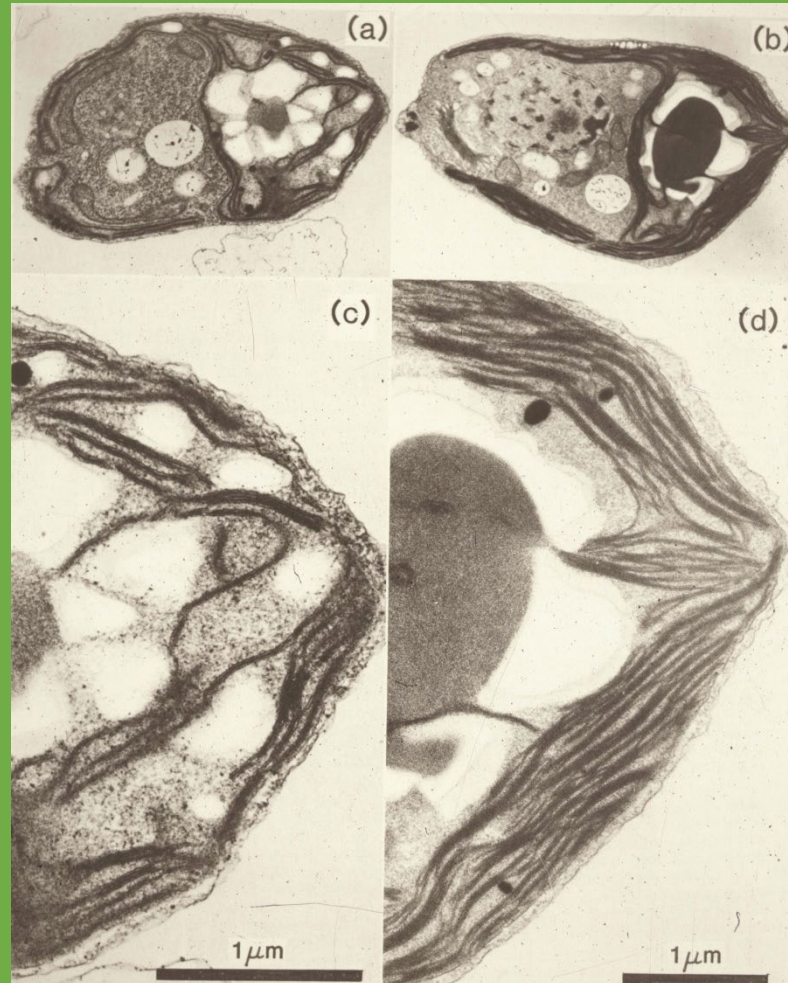
Figure 1.6 The accumulation of atmospheric oxygen through geologic time. The first oxygen produced was absorbed first by oxygen-hungry minerals that produced a banded-iron formation and later by “Red-Beds” (terrestrial rocks that absorbed atmospheric oxygen). Once these minerals had been saturated, oxygen could accumulate in the atmosphere. Modern oxygen levels were attained about 400 million years ago. (After M. Schidlowski, 1980.)



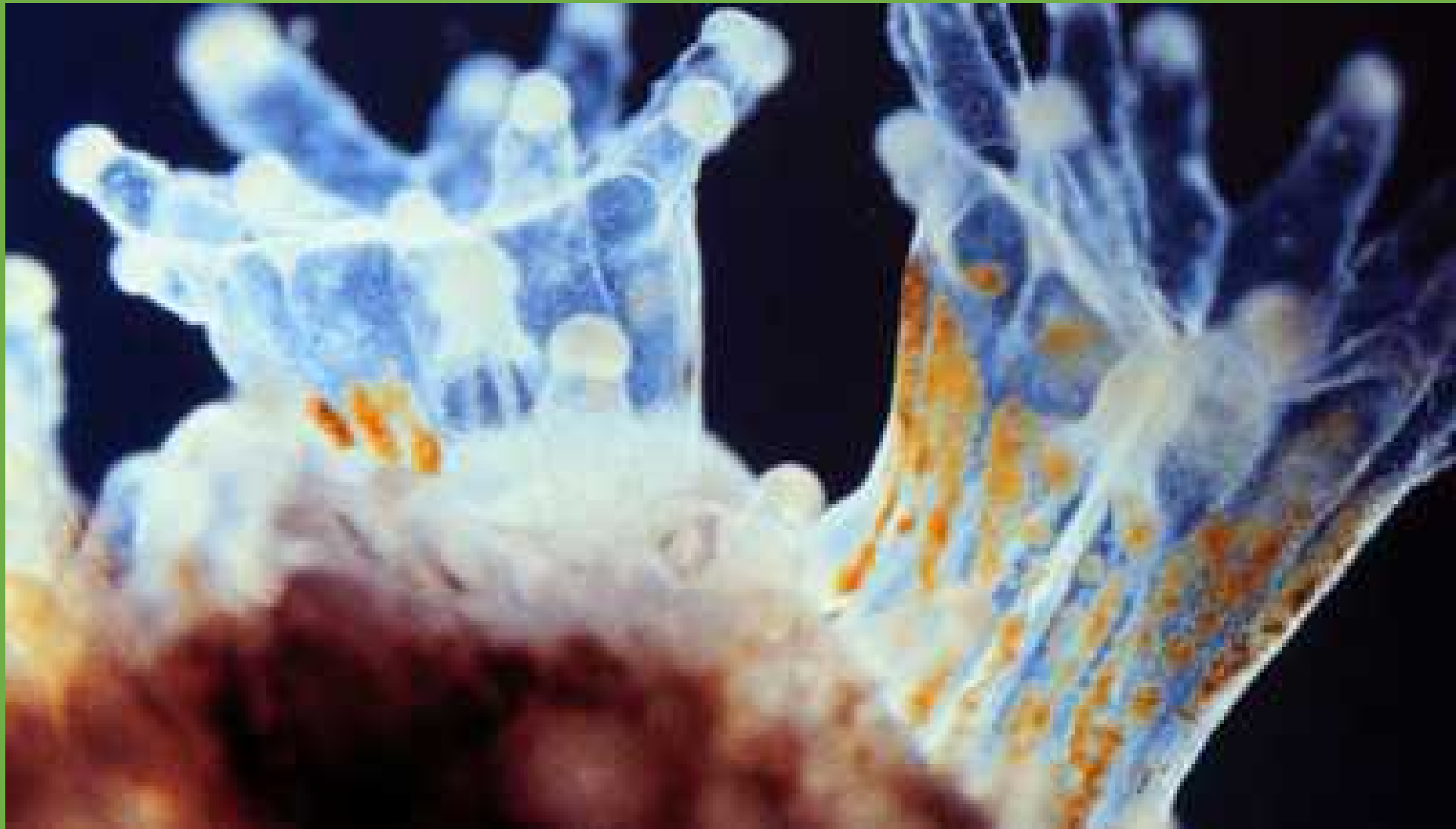
ENERGY FIXED by the earth's primary producers is equivalent to about 162 billion metric tons of dry organic matter a year, according to Robert H. Whittaker and Gene E. Likens of Cornell University. About 5 percent of the energy is fixed by agricultural ecosystems and is utilized directly by man, one species among millions. Man also draws annually on fossil fuel reserves for about the same amount of energy. In this anthropocentric view of the

biosphere the area of the concentric rings is proportional to the major ecosystems' share of the surface area of the earth (indicated in millions of square kilometers). The width of the arrows is proportional to the amount of energy fixed in each ecosystem and contributed by fossil fuels (indicated in billions of metric tons of dry matter per year). The intensity of the color in each ring suggests the productivity (production per unit area) of each ecosystem.

(b&d) מאוקלמת לאור גבוה *Dunaliella salina* פוטואקלימציה של פוטוסינטזה:
(a & c) ונמוד



פוליפים של אלמוג ובתוכם אצות שיתופיות



זואוקסנטלות
באלמוגים שונים





Low light

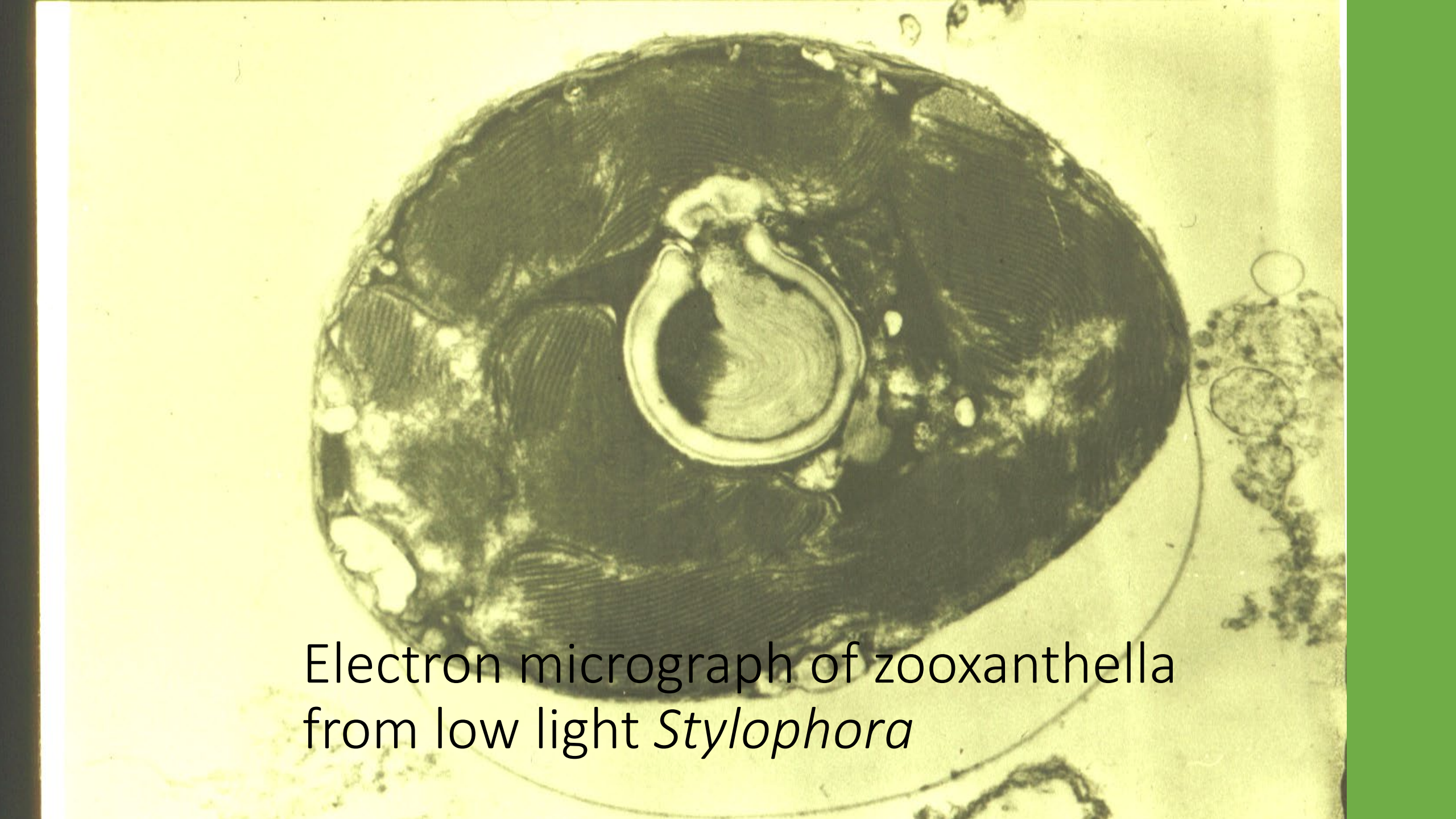


High light

Stylophora pistillata

High and low light acclimation in *Stylophora pistillata* zooxanthellae

Parameter Treatment	$\mu\text{g ch } a$ cm^{-2}	10^6 cells cm^{-2}	$\text{pg ch } a$ cell
HL	3.6 ± 1.1	1.7 ± 0.3	2.2 ± 0.3
LL	14.2 ± 4	1.6 ± 0.1	8.3 ± 0.5
LL/HL	3.9	1.1	3.7

An electron micrograph showing a cross-section of a zooxanthella cell. The cell is roughly circular and contains a large, central, electron-dense structure with a distinct, concentric, layered appearance, likely representing a chloroplast or a specialized organelle. The surrounding cytoplasm is filled with various organelles, including smaller, electron-dense structures and a network of membranes. The overall appearance is that of a highly organized, specialized cell. The image is presented in a yellowish-green color scheme.

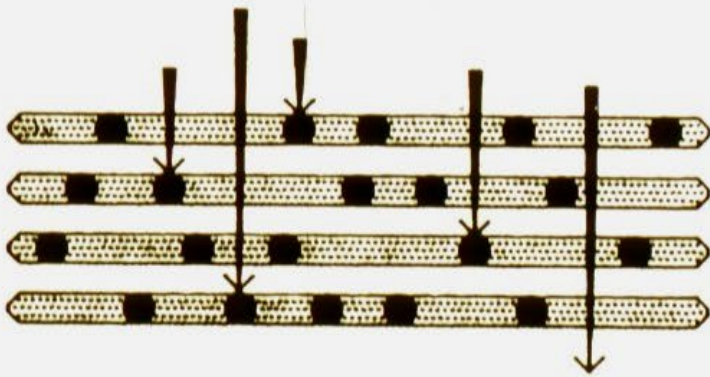
Electron micrograph of zooxanthella
from low light *Stylophora*

Electron micrograph of zooxanthella from high light *Stylophora*

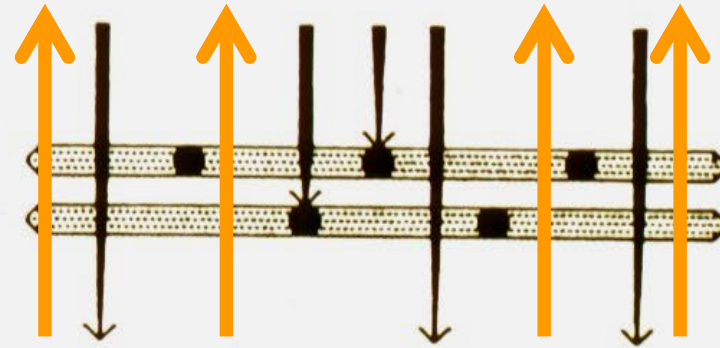


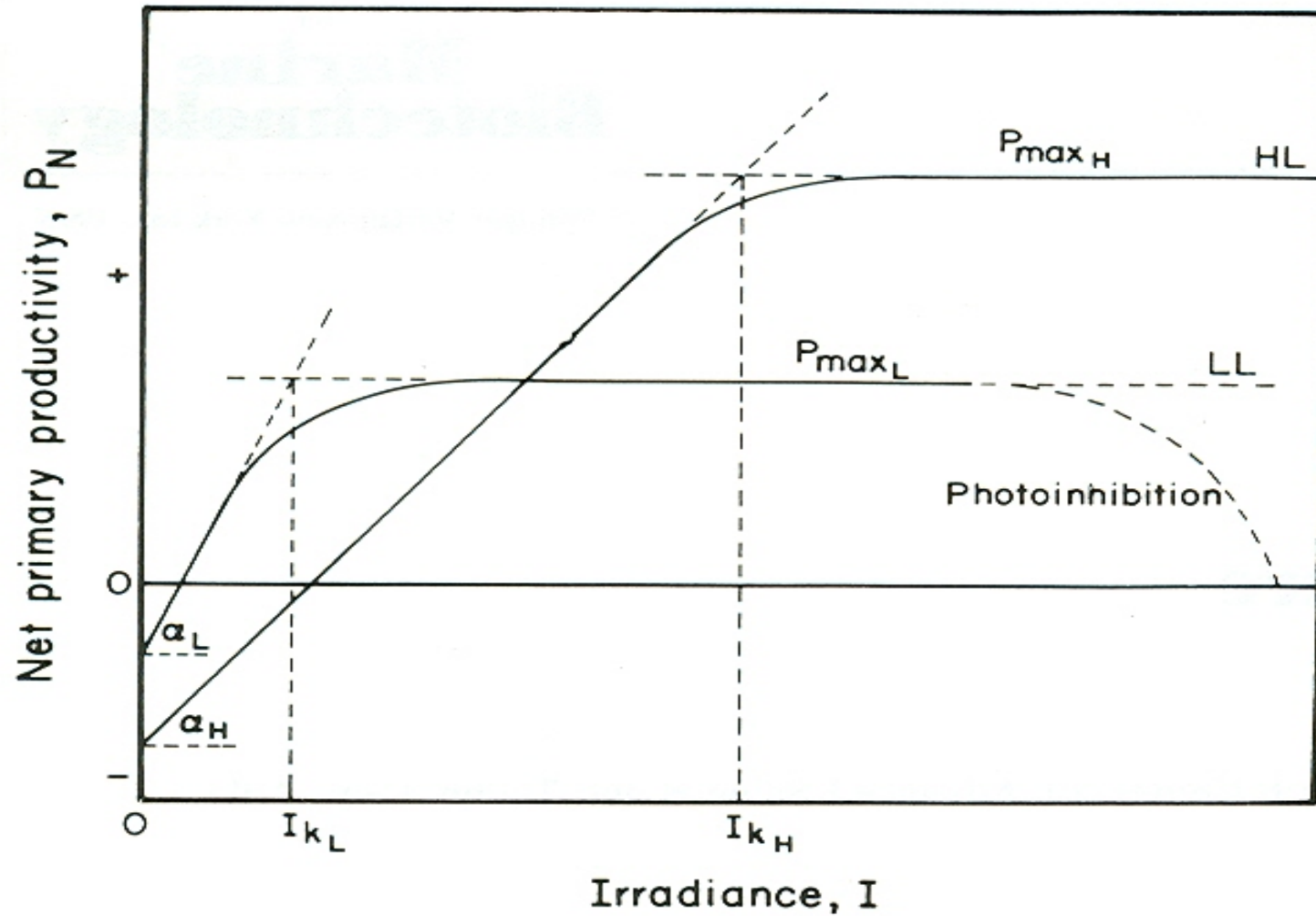
Photoacclimation strategy

Low Light

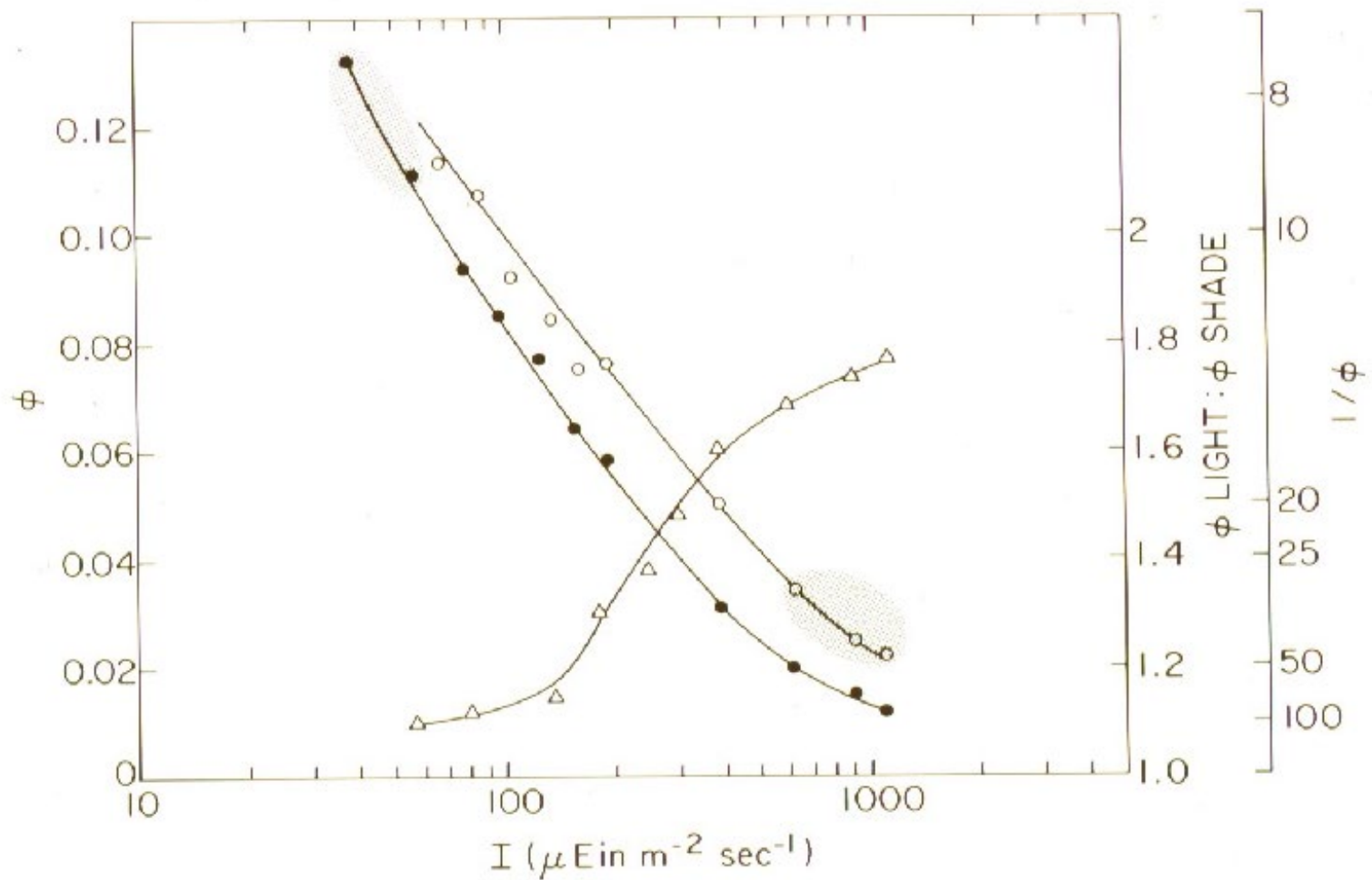


High Light



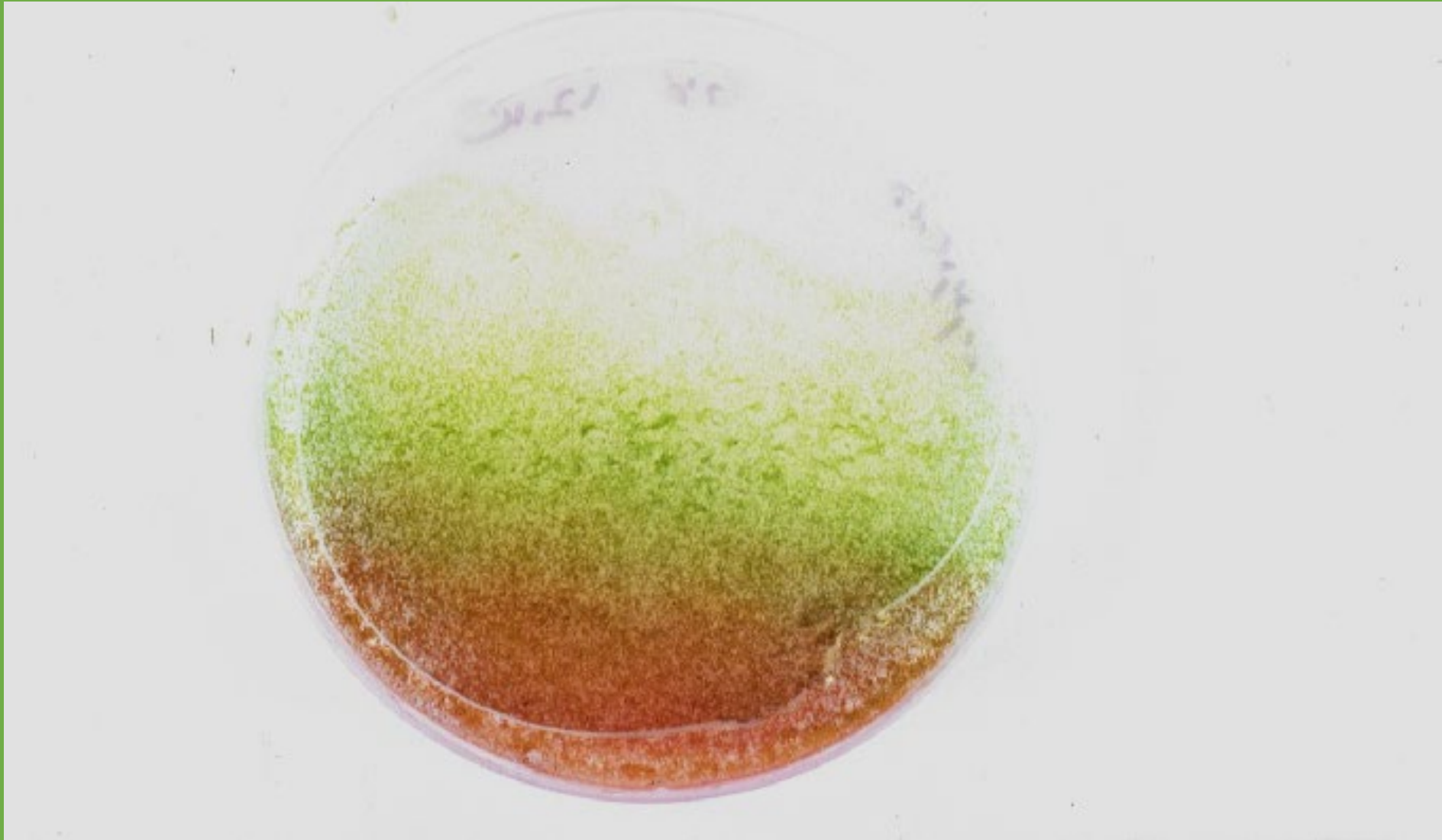


לאצות (LL) לאור גבוה ו- (HL) פוטוסינתזה ועצמת אור מאוקלמות לאור חלש



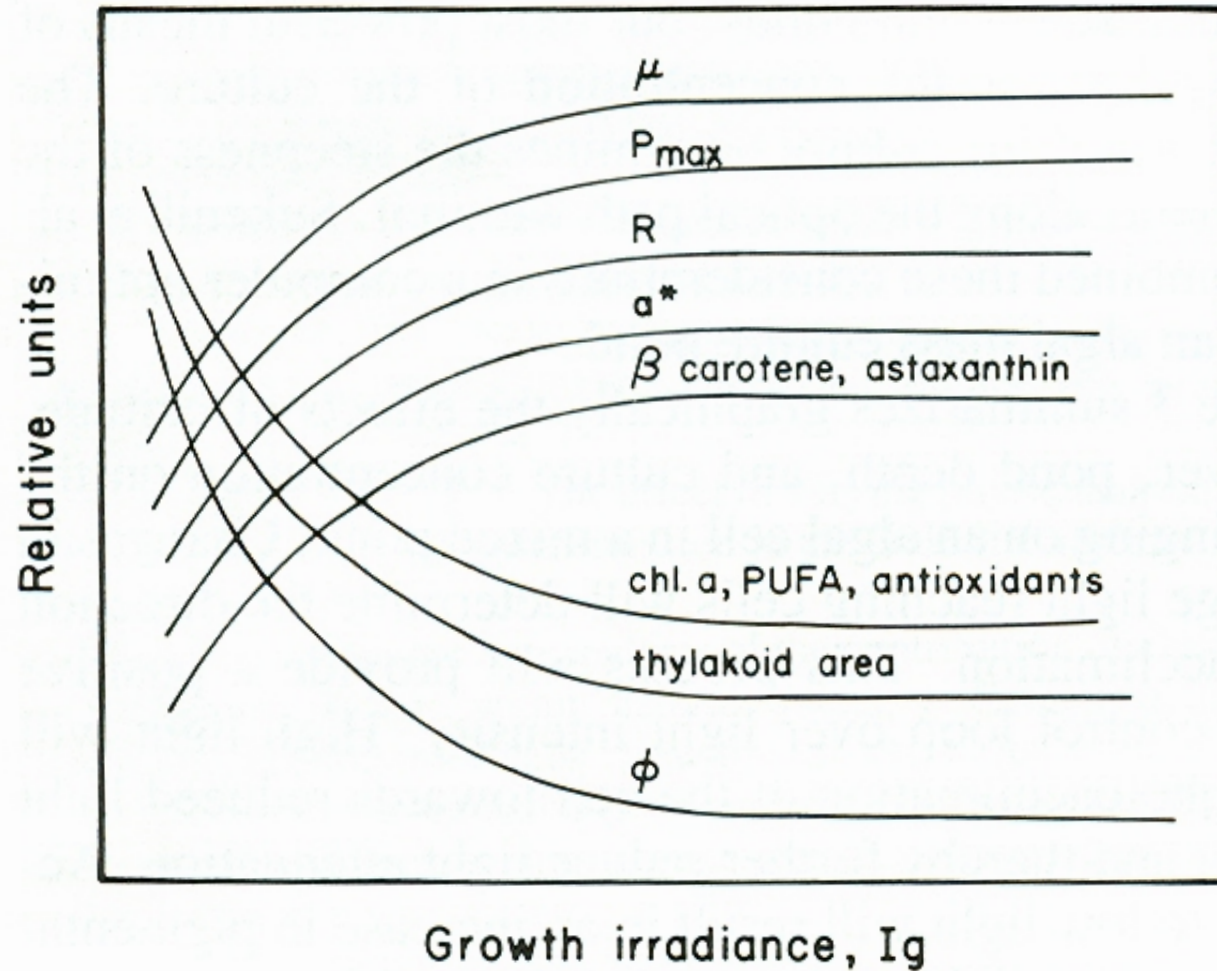
(בזאוקסנטלות ϕ נצולת פוטוסינטטית)
 מאלמוגי אור גבוה ונמוך

Haematococcus pluvialis תרבית של
חזק למטה, אור חלש למעלה,

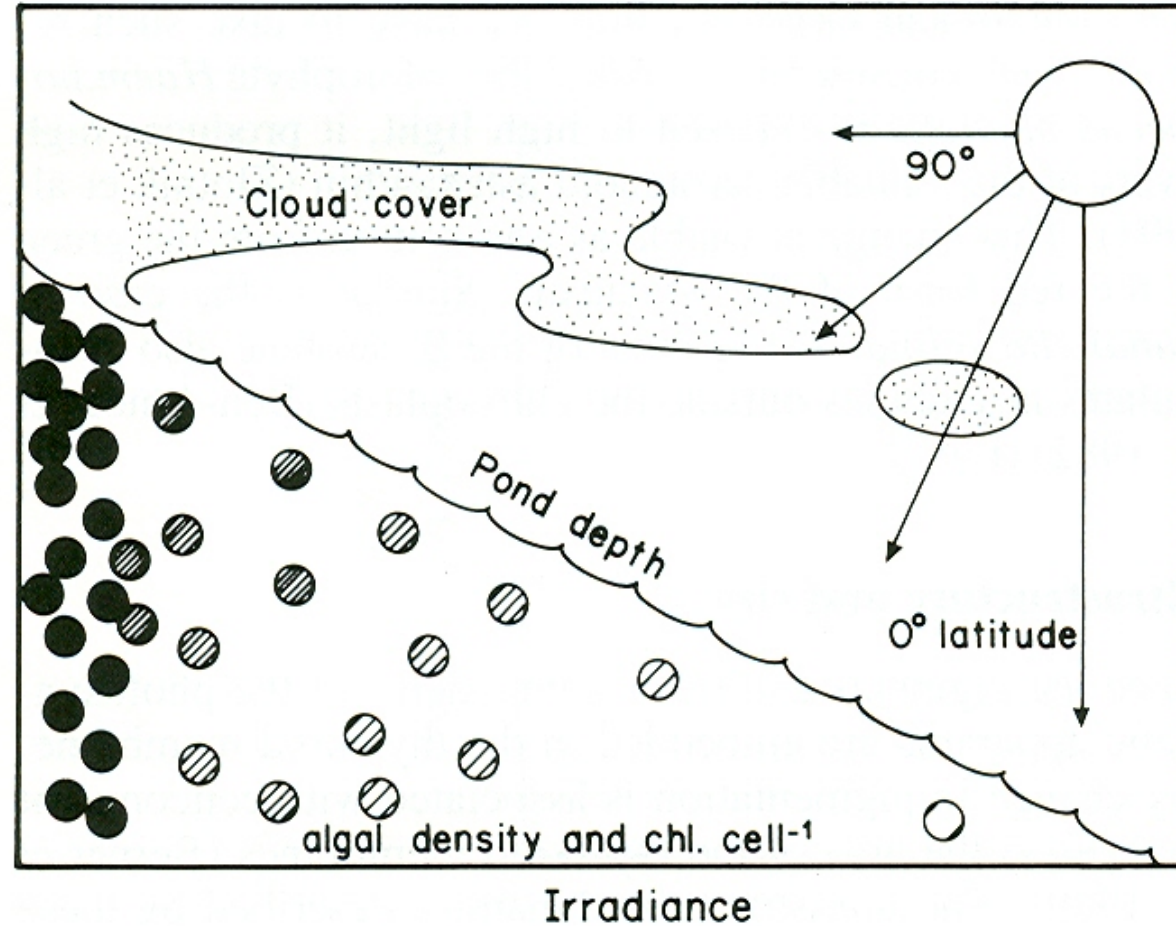


Carotenoids הקרוטנואידים

- מתחלקים ל"קוצרי אור"
- Fucoxanthine
- Peridinine
- "למגיני אור" photoprotective
- B carotene
- Astaxanthine
- Xanthophyll פיגמנטים של מעגל ה-



השפעת עוצמת אור גידול של אצות על קצבי
הפוטוסינתזה והגדילה שלהן

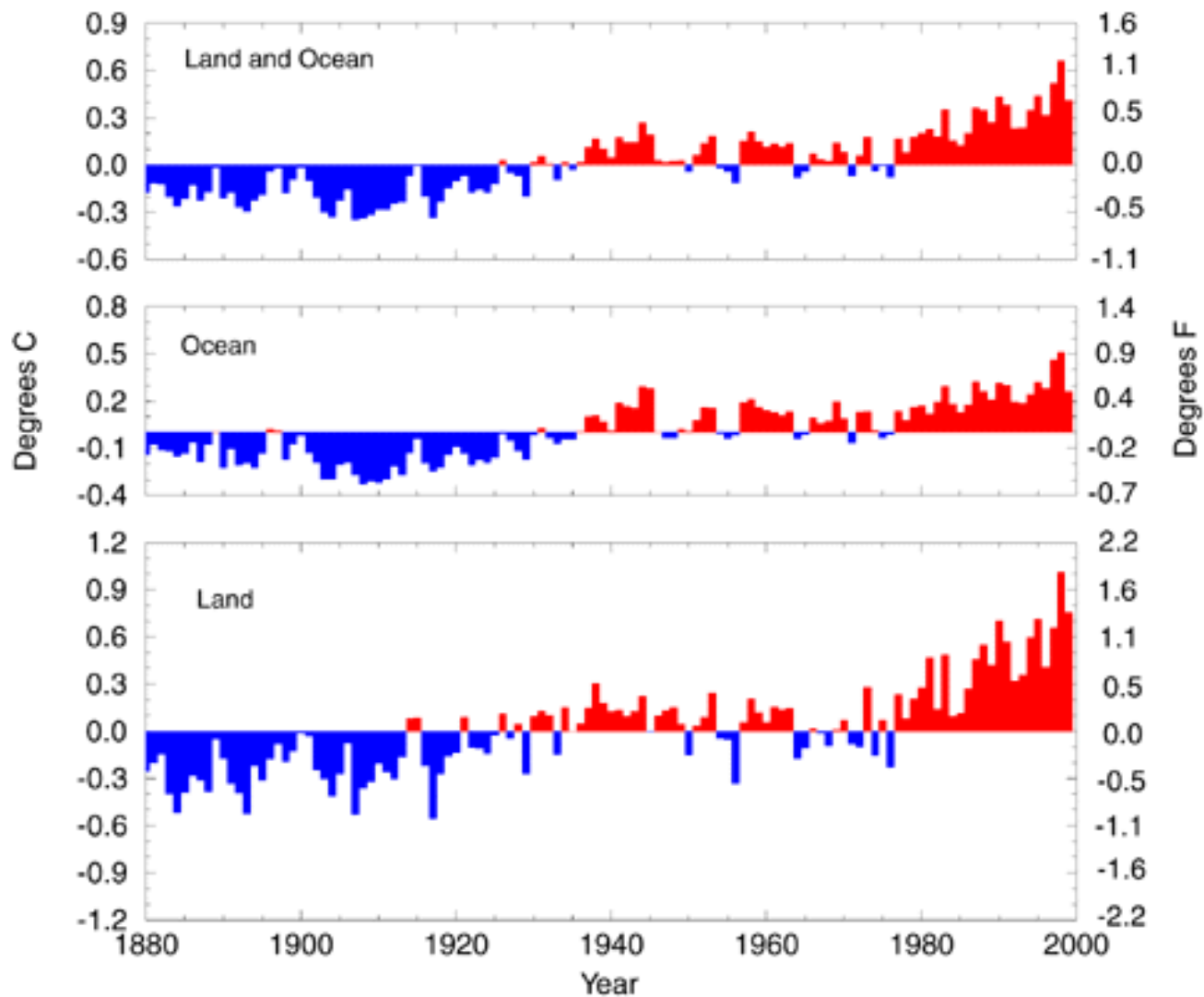


Schematic representation of the effects of latitude, cloud cover, pond depth and cellular pigment content on the time integrated average photon flux



Annual Global Surface Mean Temperature Anomalies

National Climatic Data Center/NESDIS/NOAA



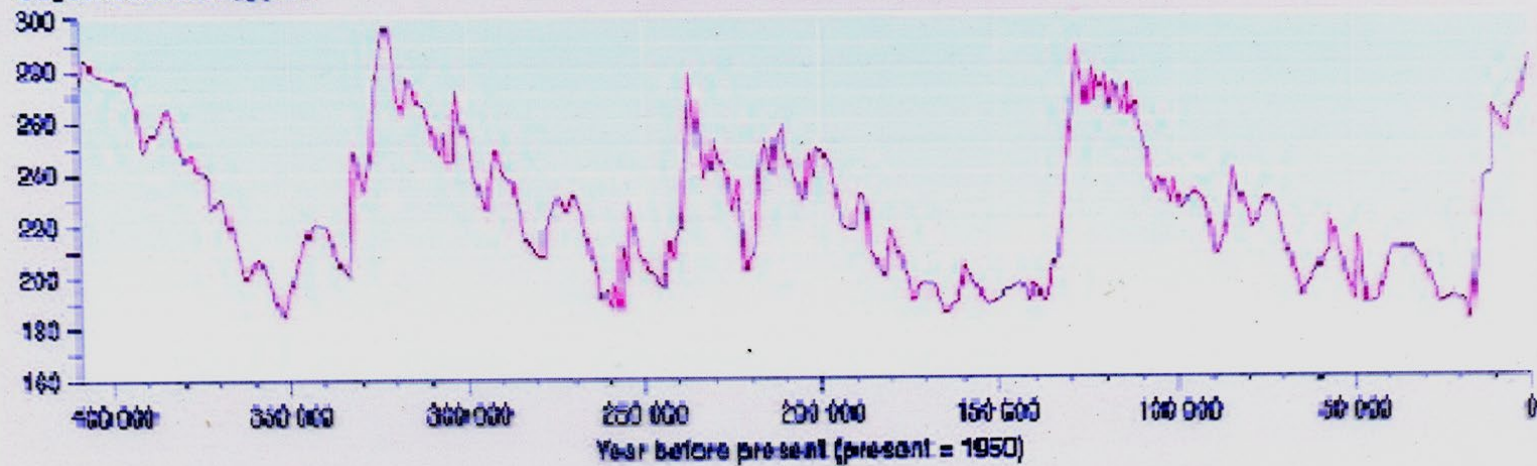




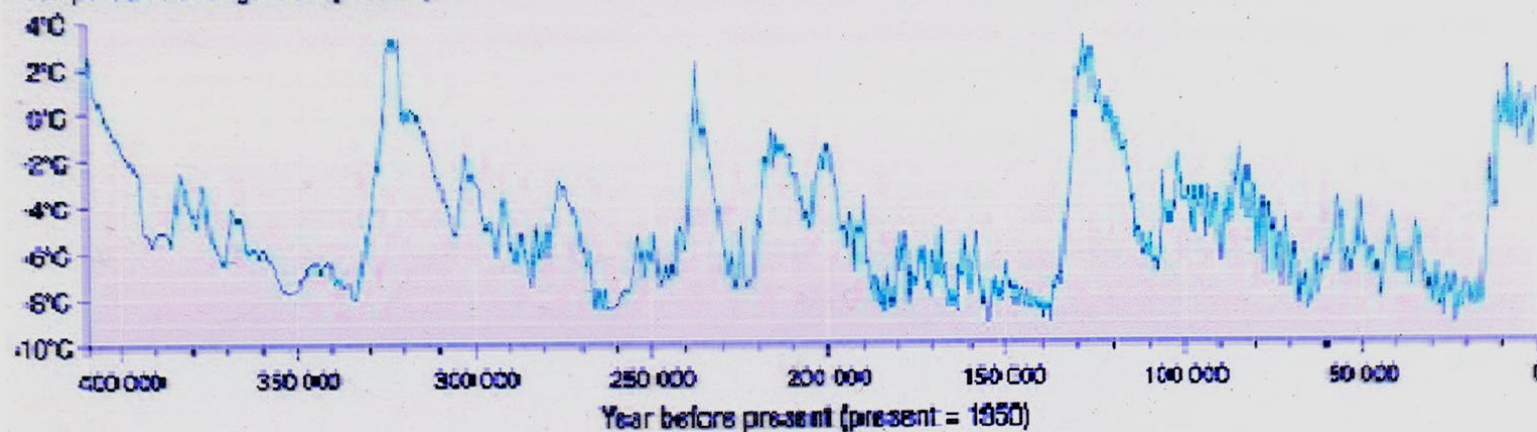


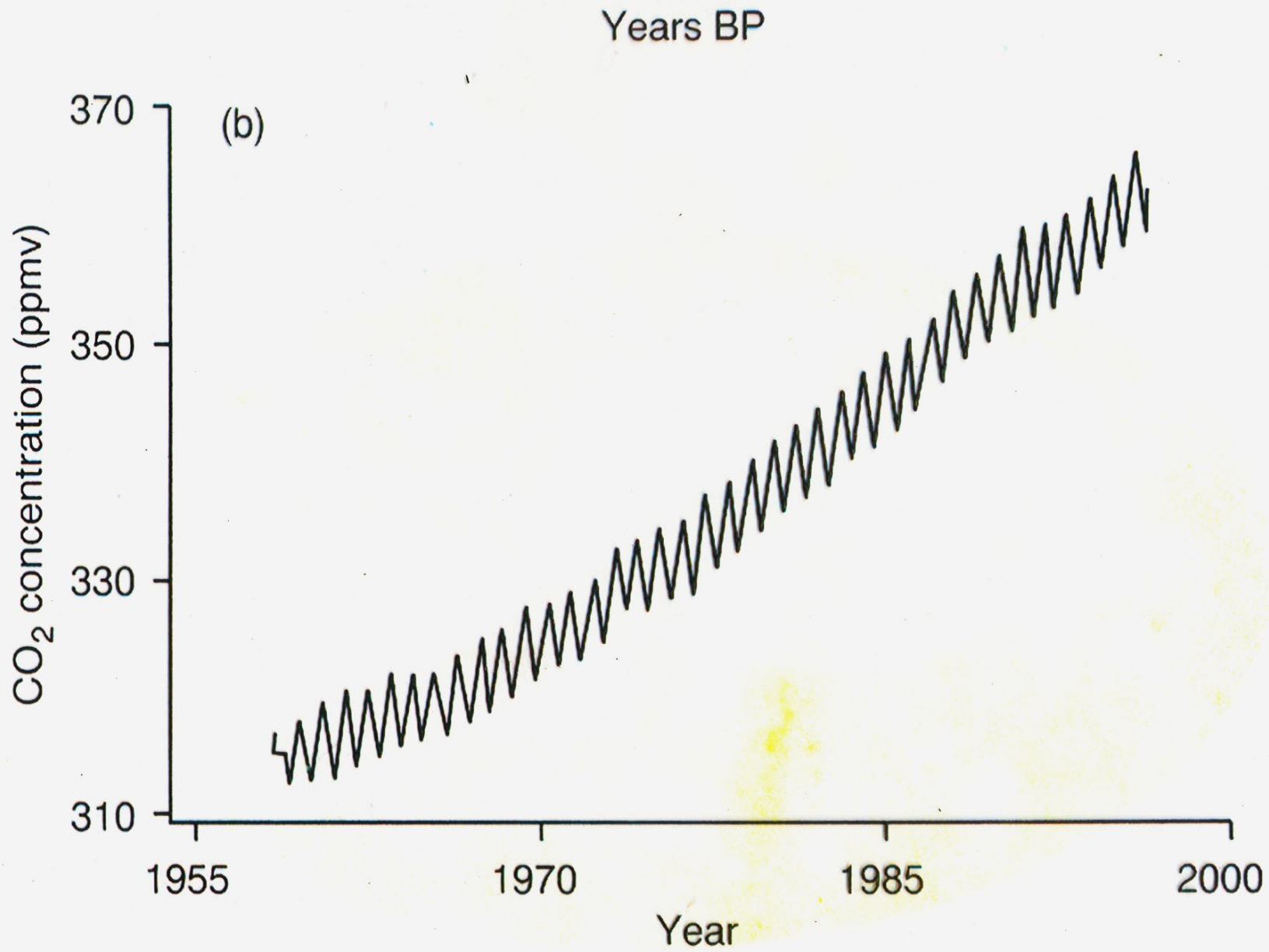
Temperature and CO₂ concentration in the atmosphere over the past 400 000 years (from the Vostok ice core)

CO₂ concentration, ppmv

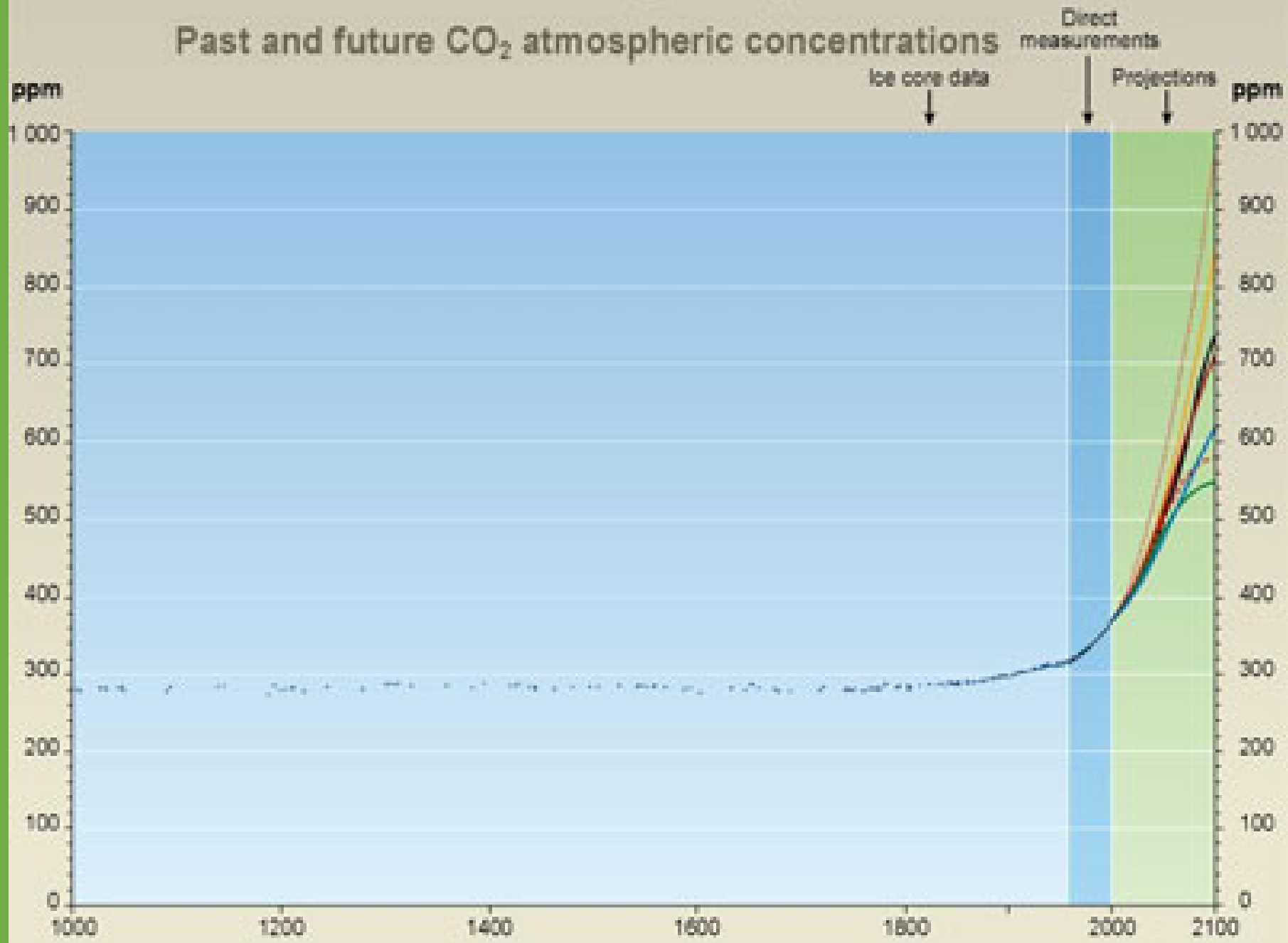


Temperature change from present, °C





Past and future CO₂ atmospheric concentrations



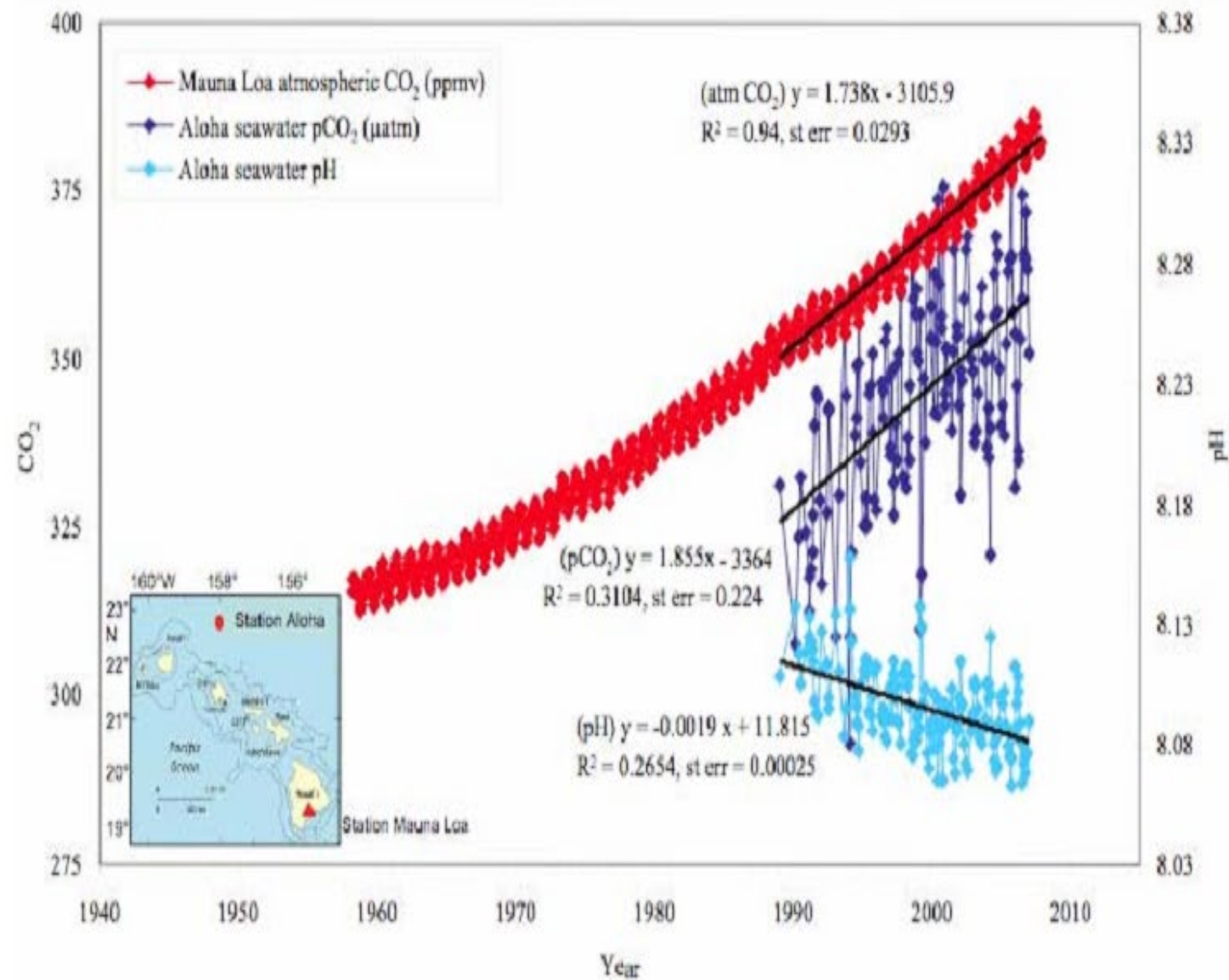
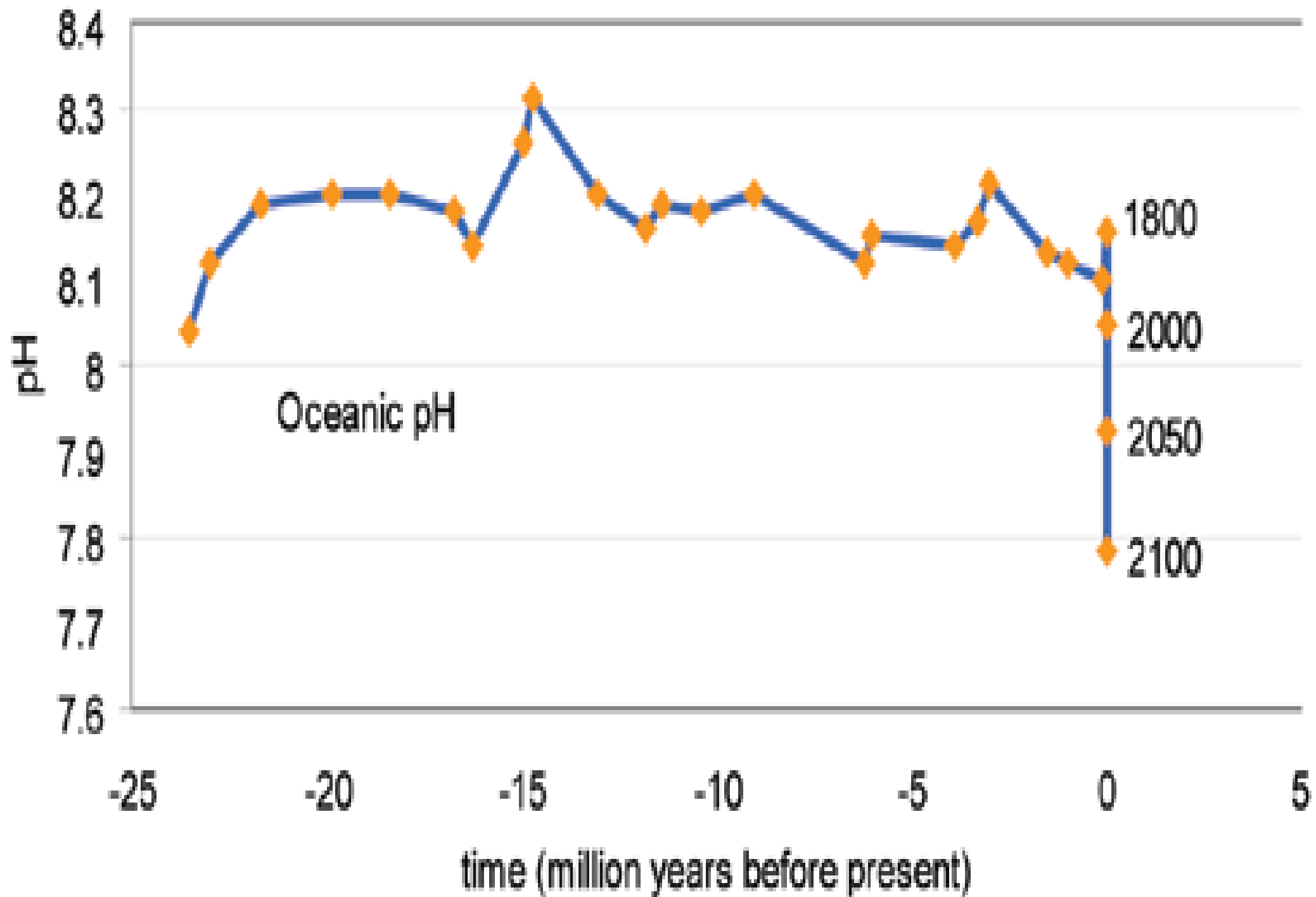
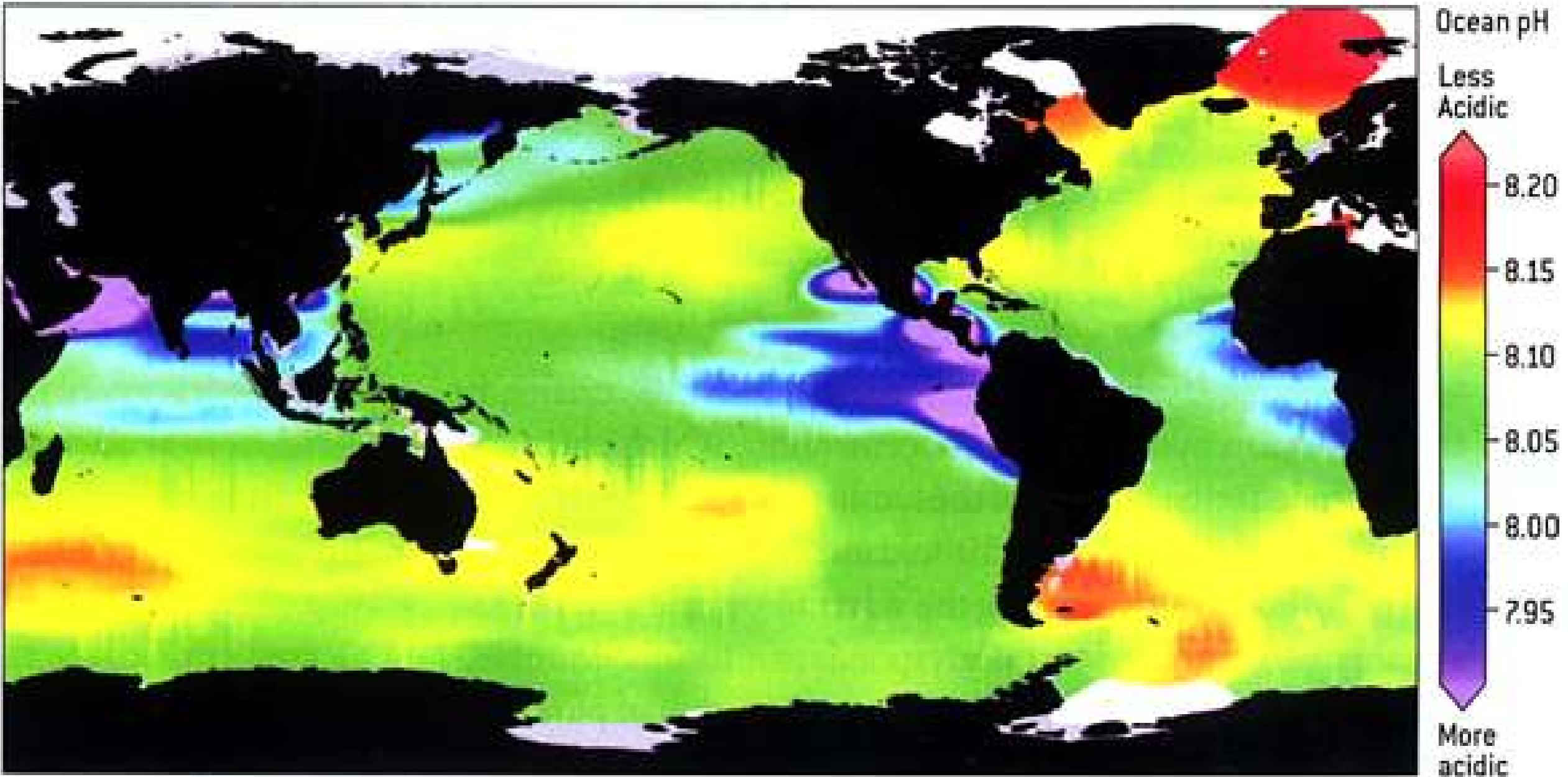


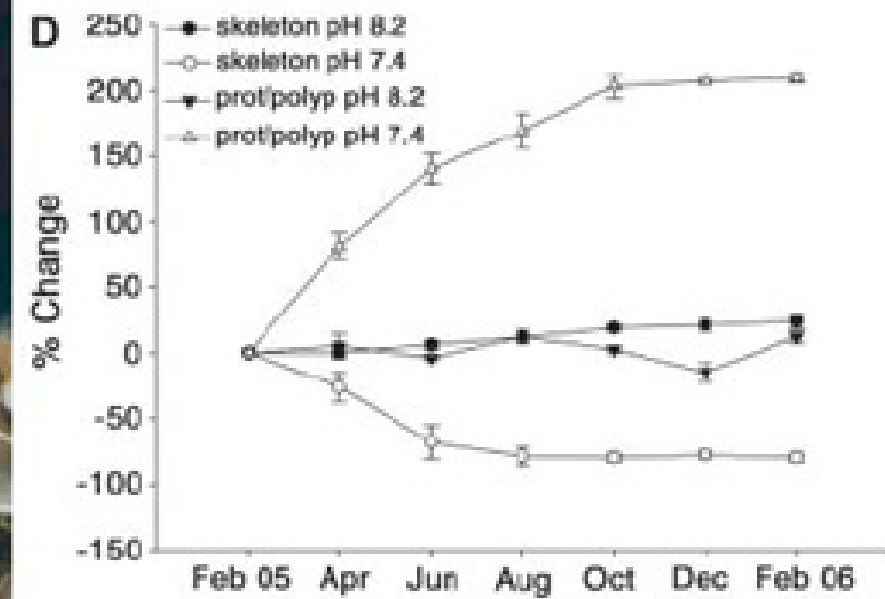
Fig. 1 Time series of atmospheric CO₂ at Mauna Loa (ppmv) and surface ocean pH and pCO₂ (μatm) at Ocean Station Aloha in the subtropical North Pacific Ocean. Note that the increase in oceanic CO₂ over the last 17 years is consistent with the atmospheric increase within the statistical limits of the measurements. Mauna Loa data: Dr. Pieter Tans, NOAA/ESRL (<http://www.esrl.noaa.gov/gmd/ccgg/trends/>); HOTS/Aloha data: Dr. David Karl, University of Hawaii (<http://hahana.soest.hawaii.edu>).

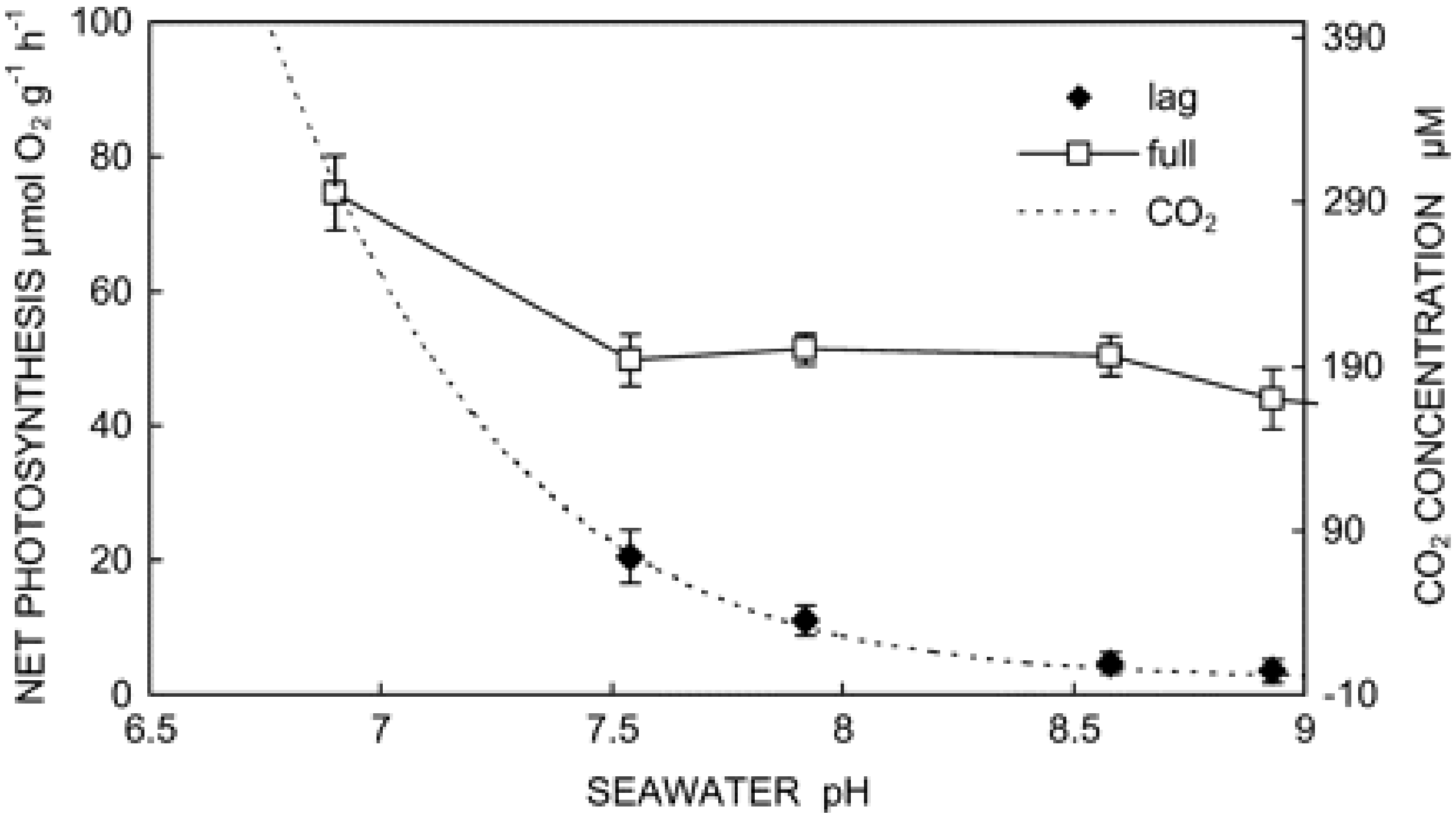




Ocean pH

Source: Scott C Doney, SciAm March 2006



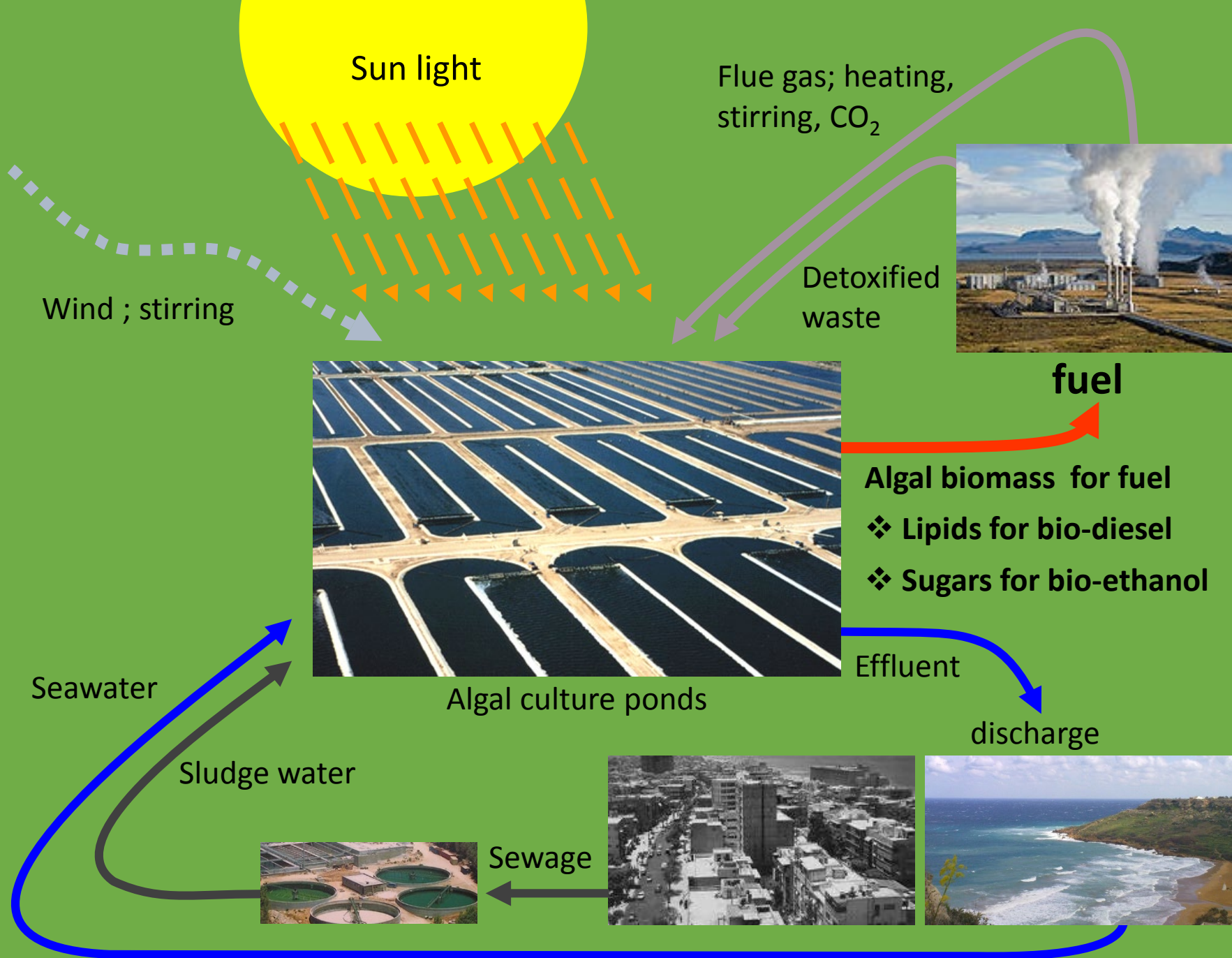


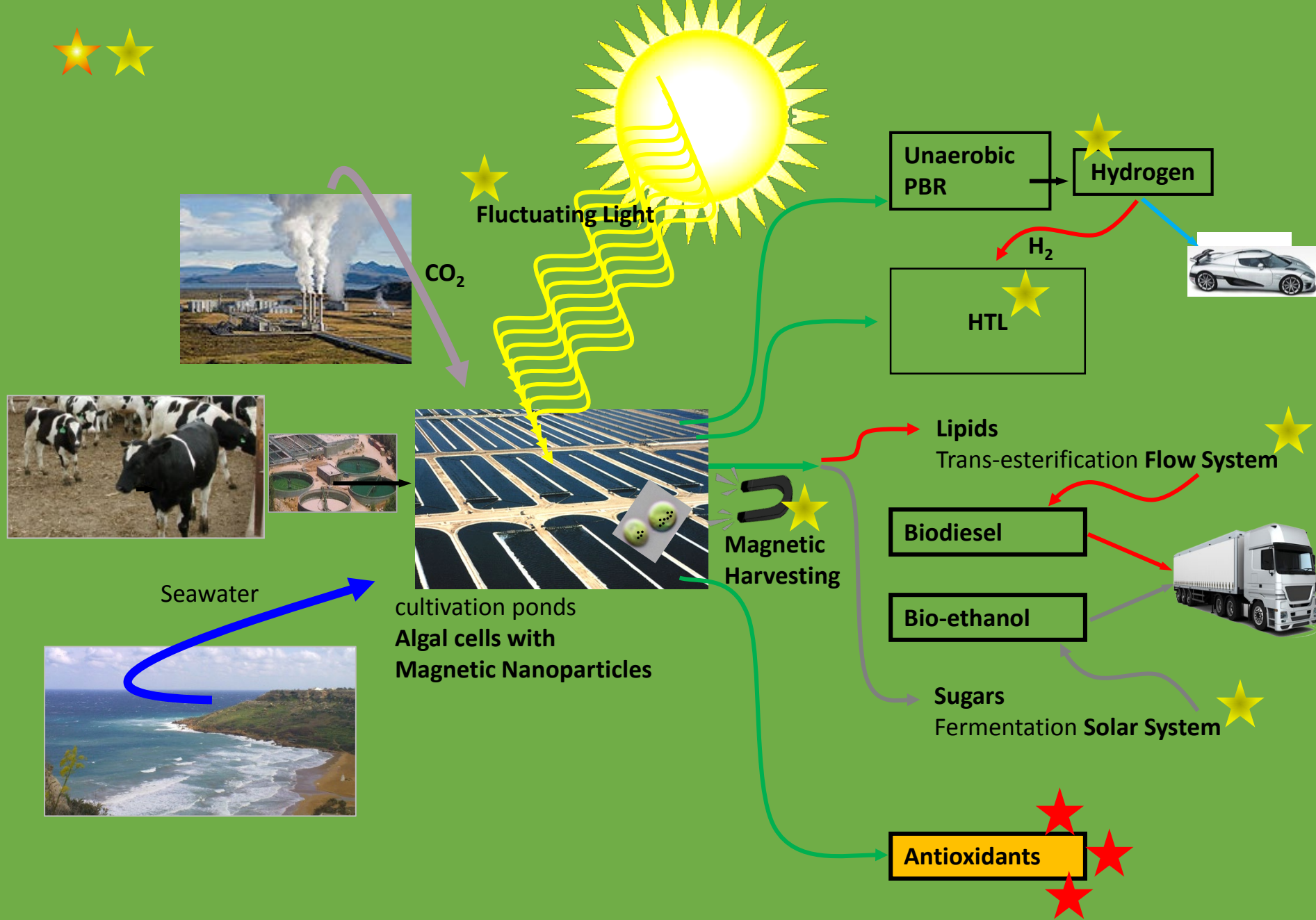
Why algae
Quantum yield
Areal efficiency
=income

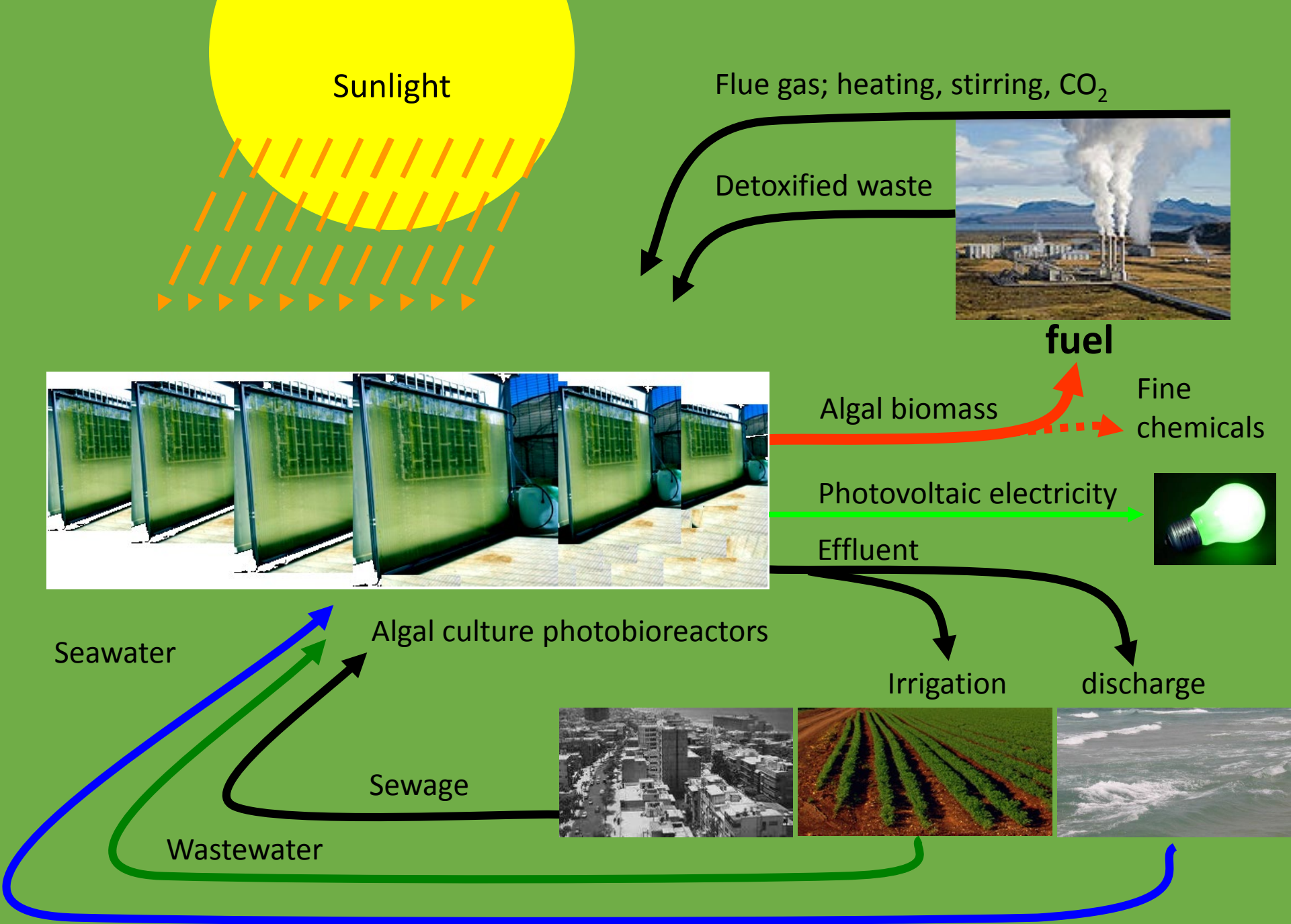


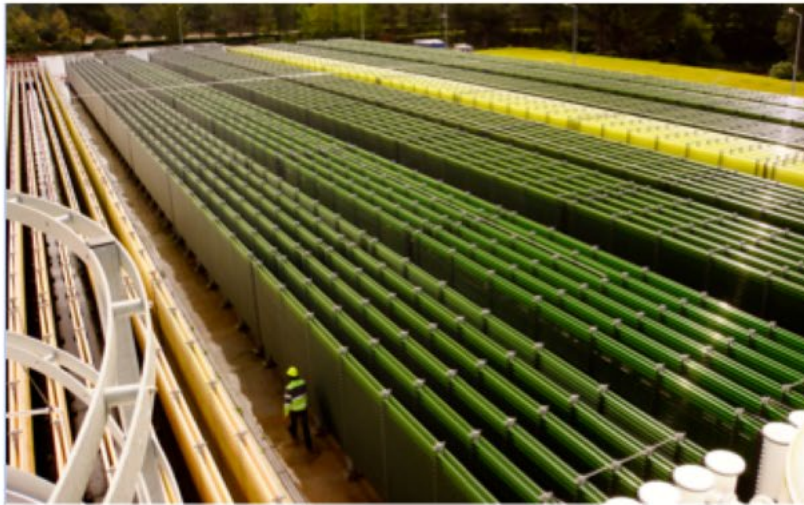












AlgaFarm designed, built, operated by A4F
Tubular Photobioreactors > 1.300 m³

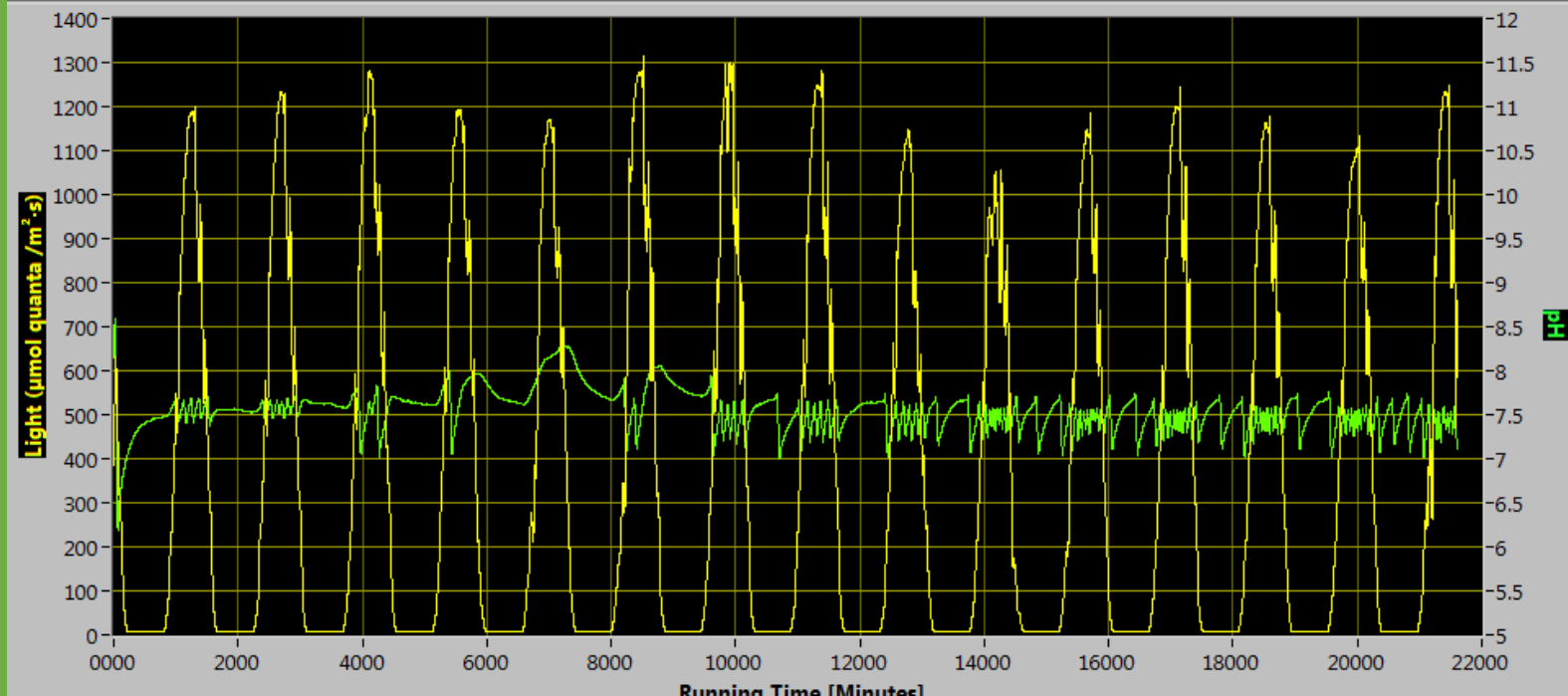
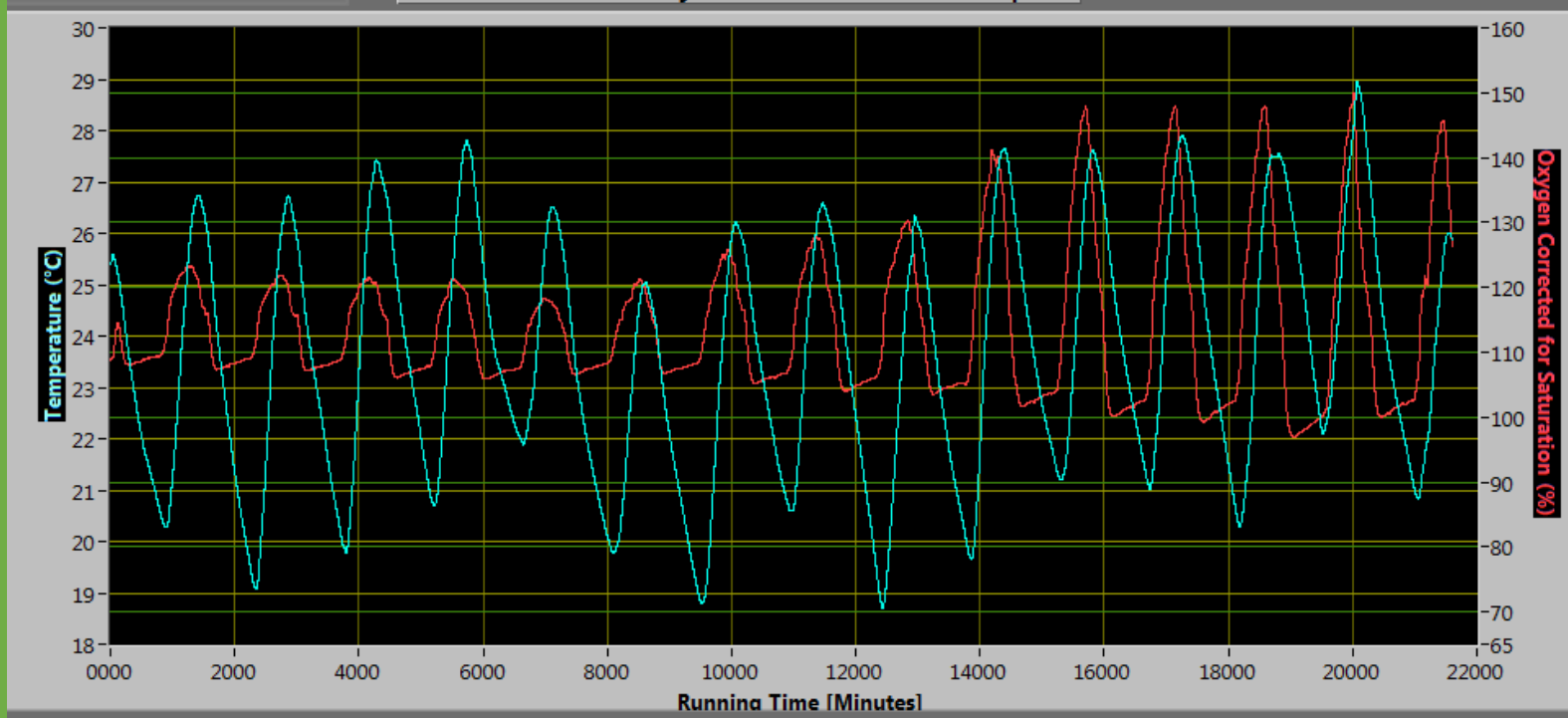


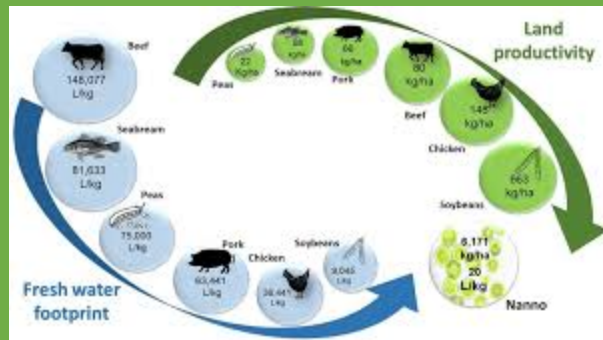
BIOFAT designed, built, operated by A4F
Cascade Raceways > 3.000 m² - 900 m³

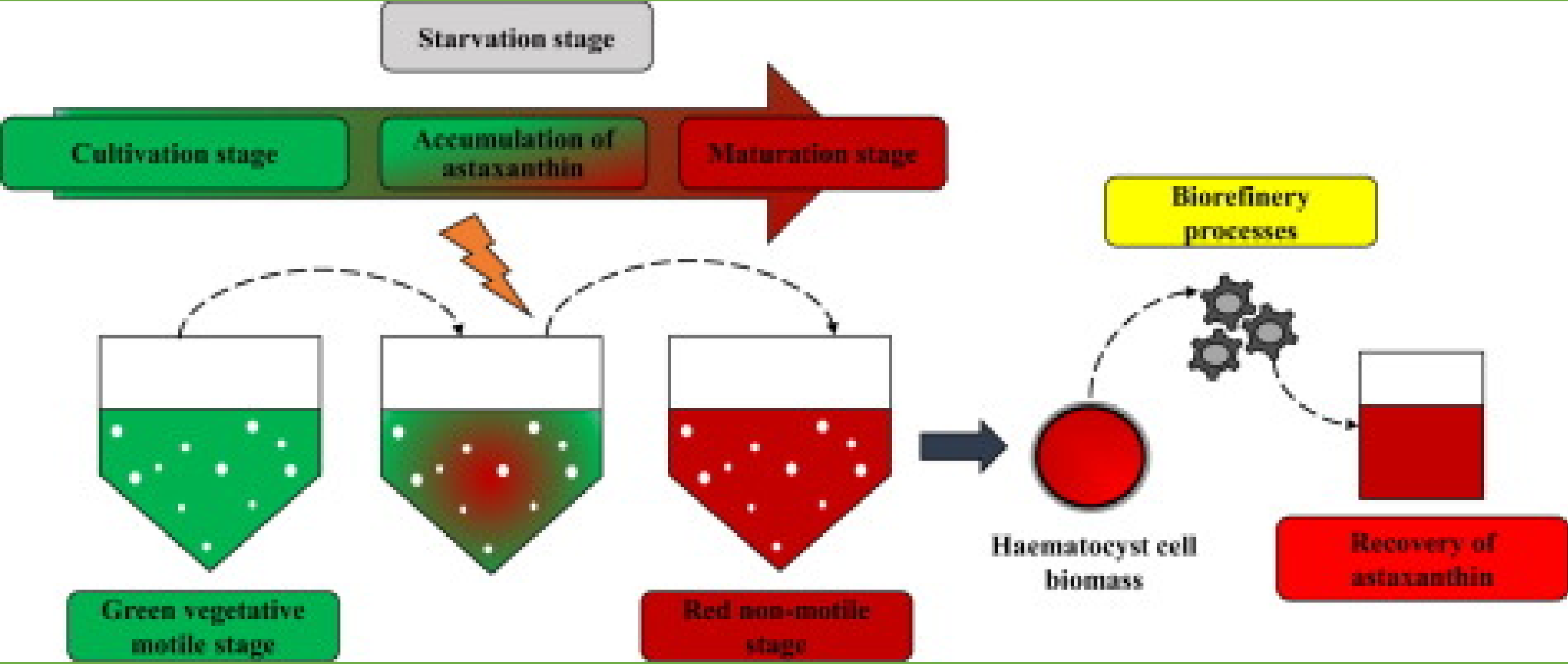
Algal biotechnology unit at Bar Ilan University

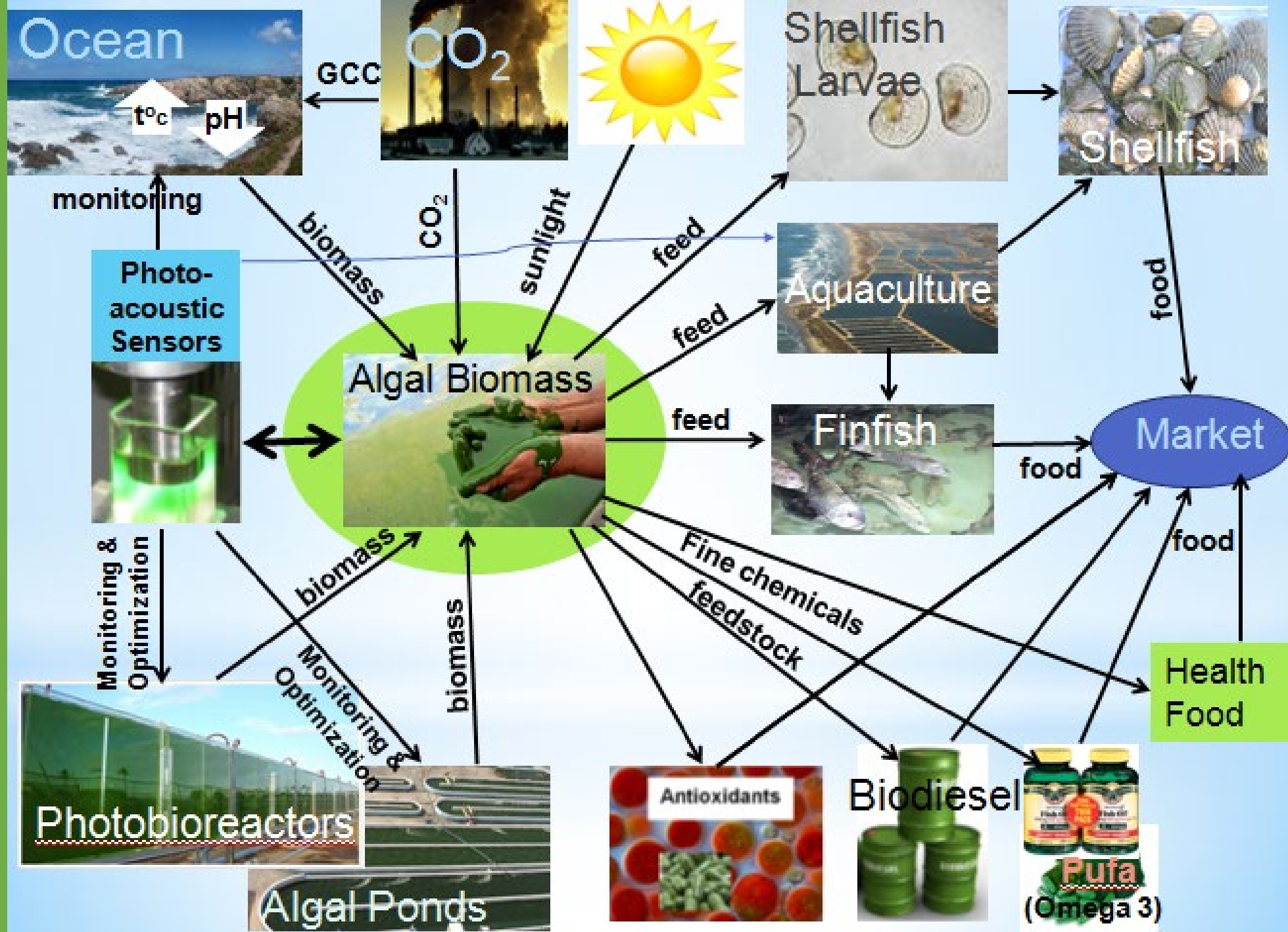
Ponds are fully computerized and optimized for the production of biodiesel and fine chemicals, using seawater and brackish water, fertilizer from sewage and CO₂ from industry.











a



b



Total Algae



Share of Macroalgae algae companies

- Macroalgae
- Microalgae
- No data/companies

Macroalgae



Share of Macroalgae production systems

- Aquaculture
- Farming
- Other
- No data/companies

Microalgae

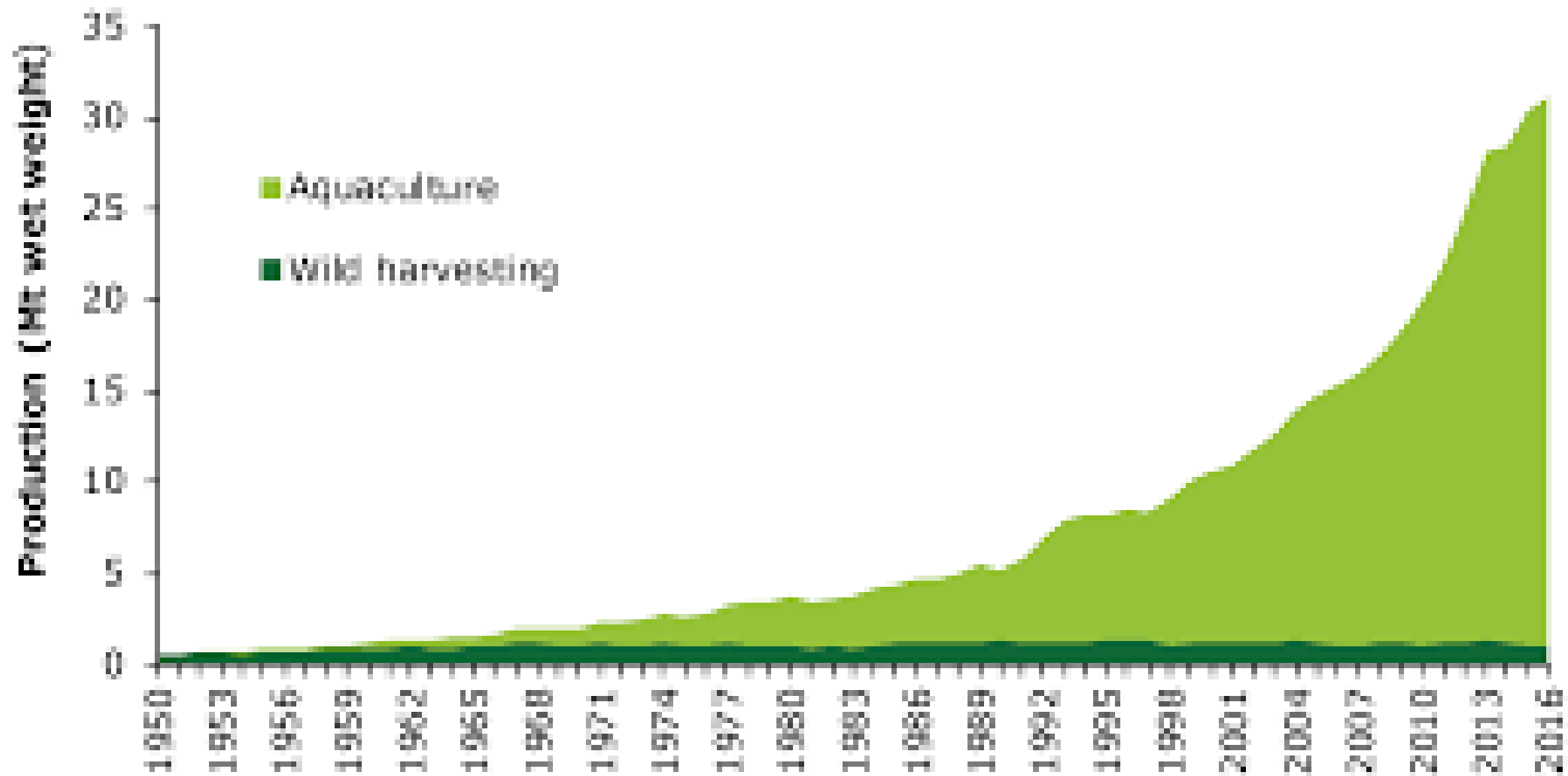


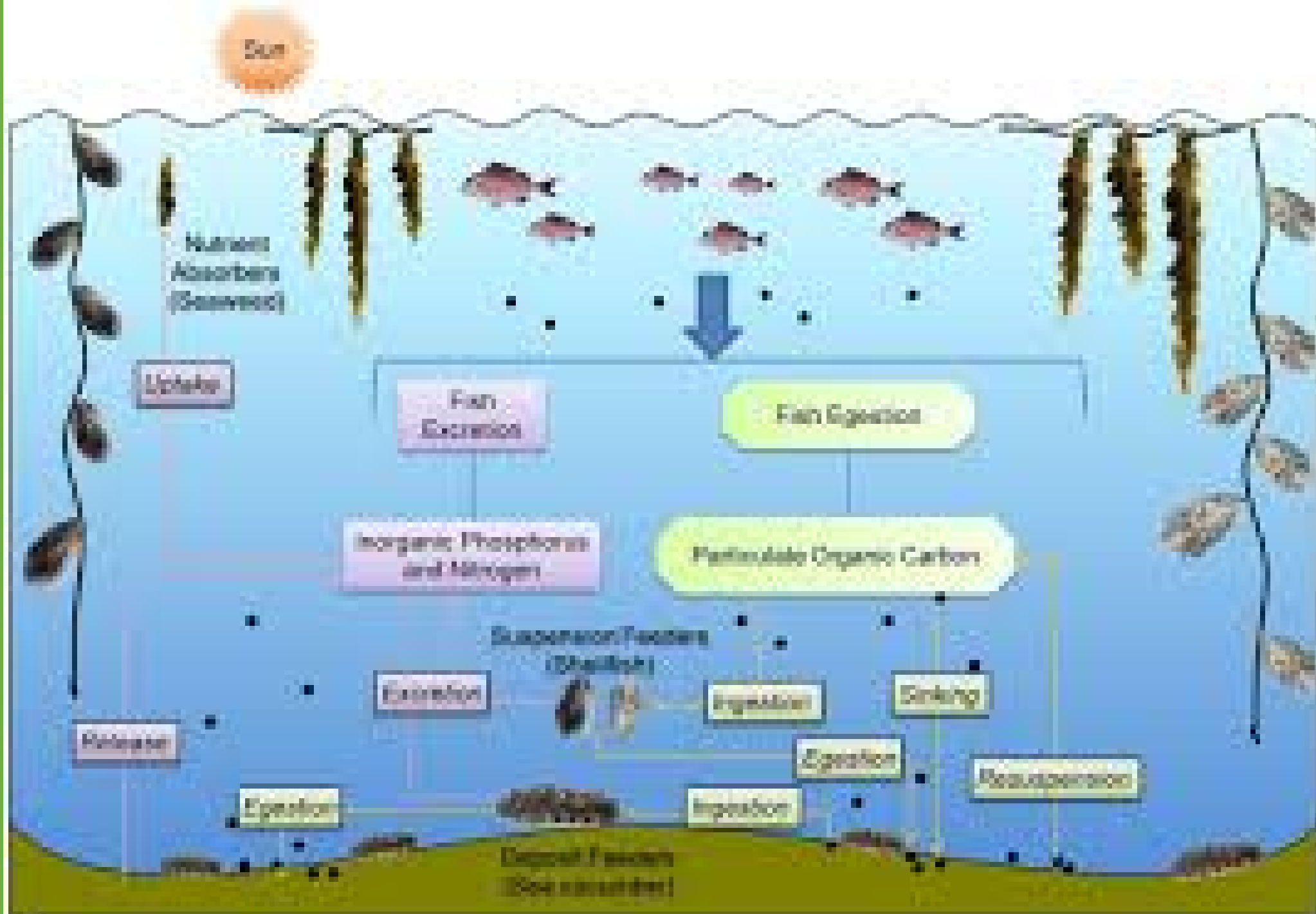
Share of Microalgae production systems

- Photobioreactors
- Open ponds
- Fermenters
- Other
- No data/companies

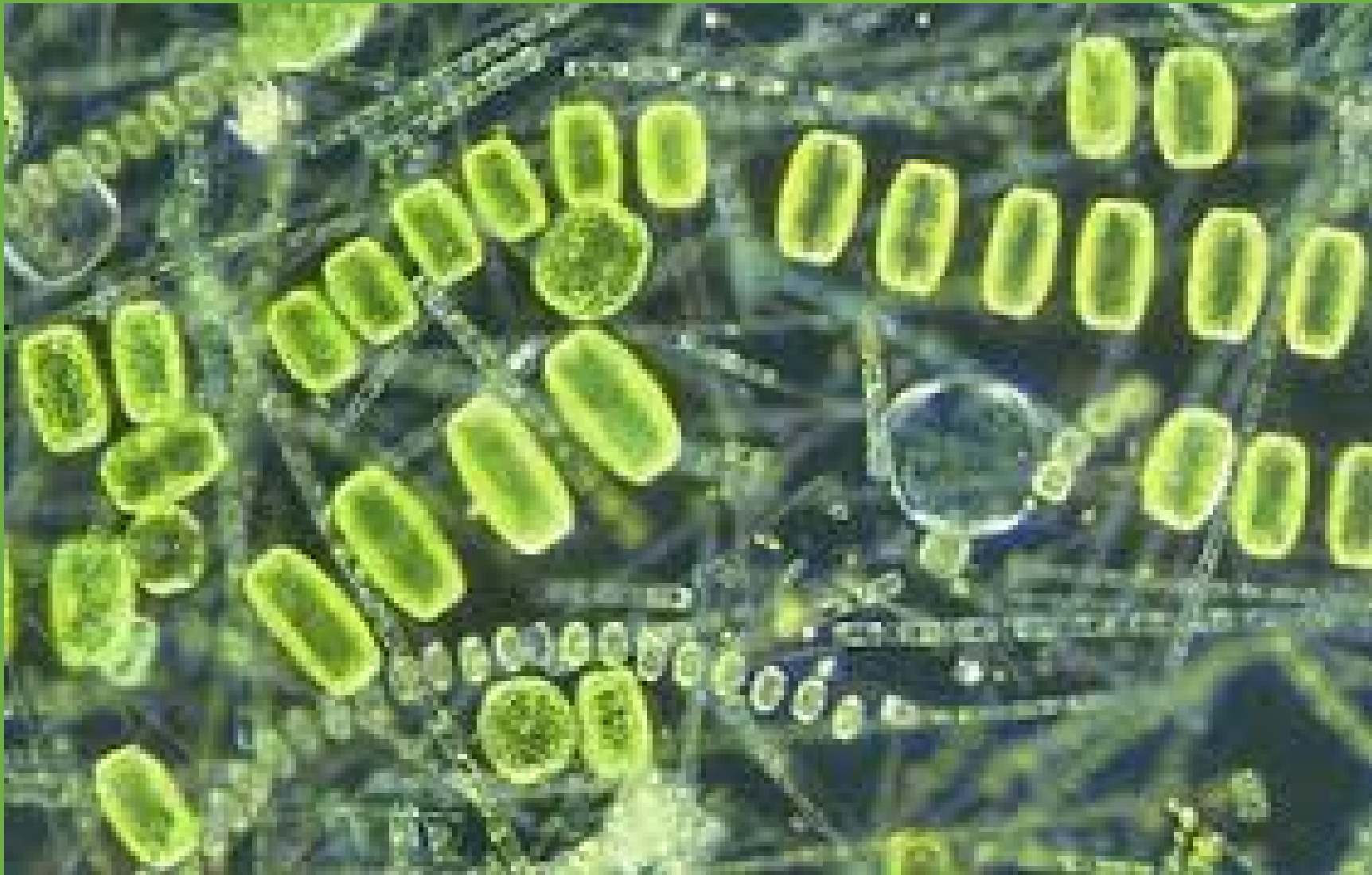
Source: Algae Europe
www.algae-europe.com

Global macroalgae biomass production



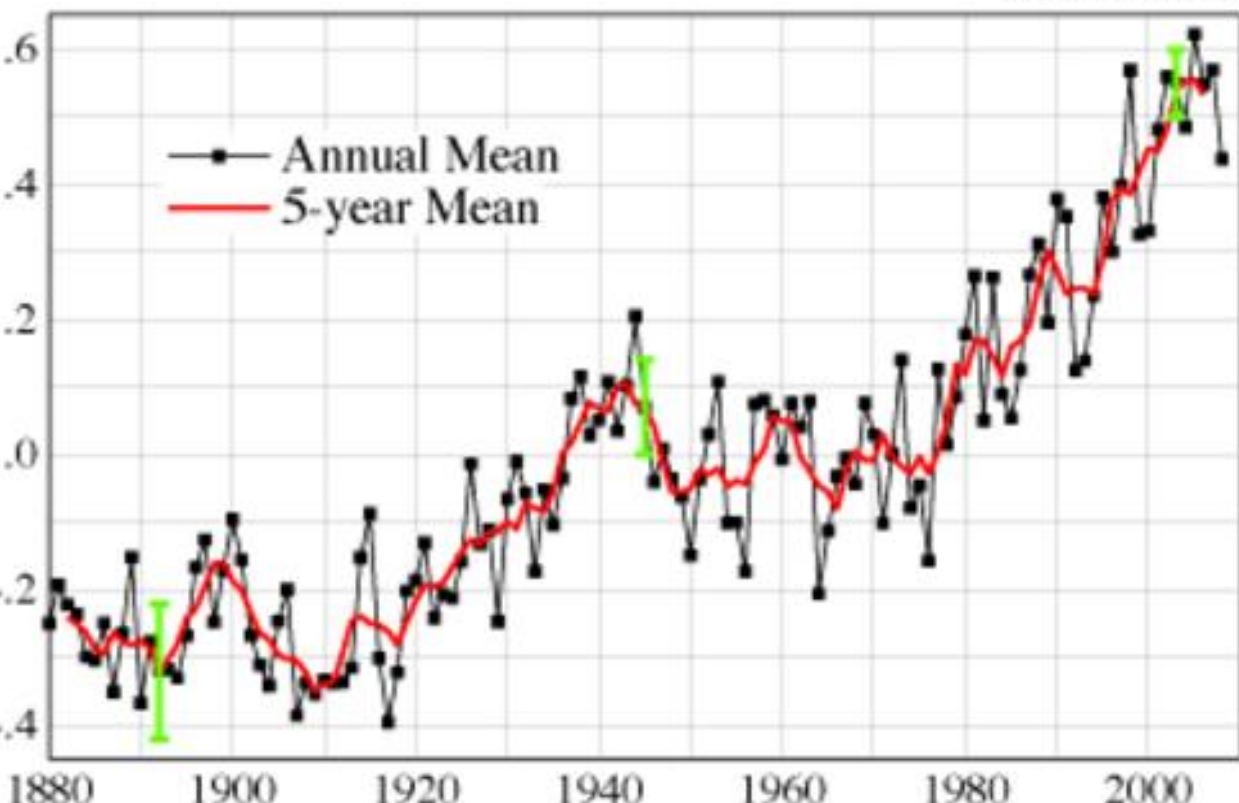






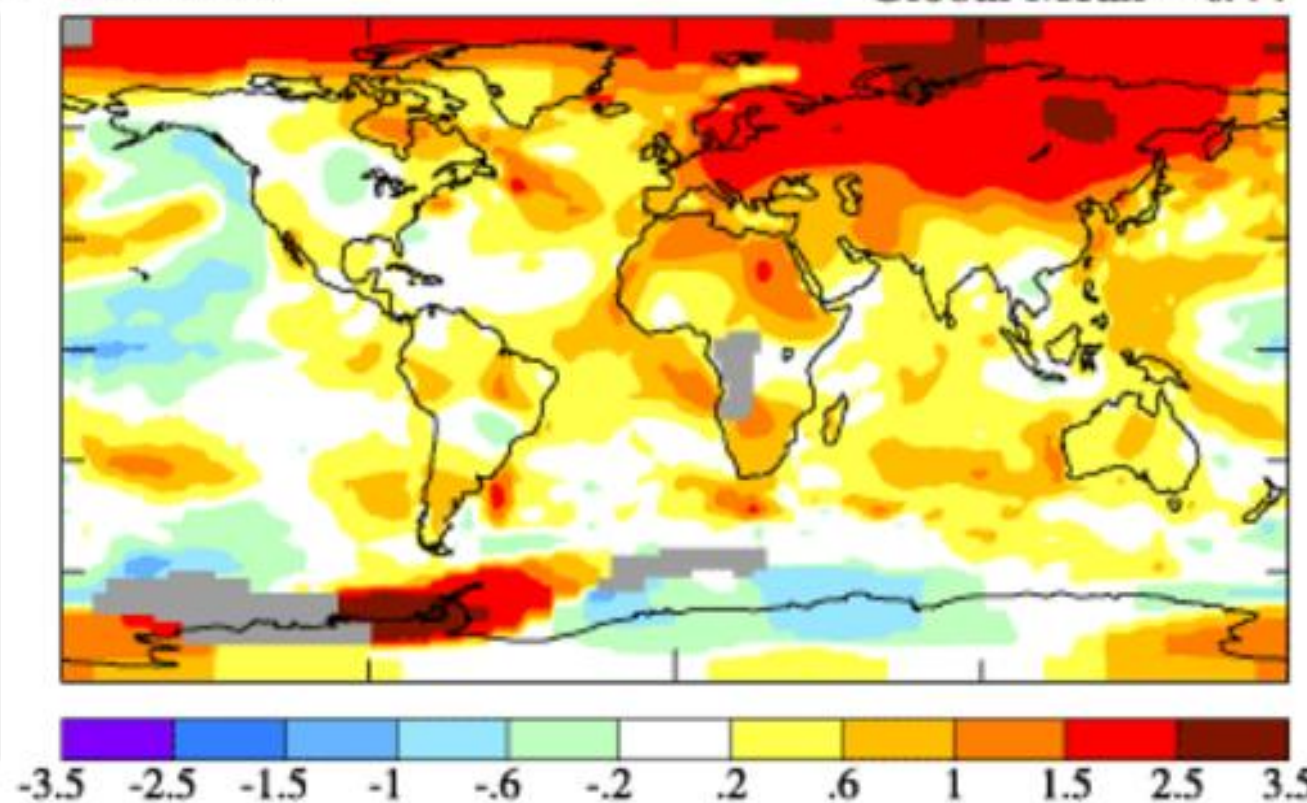
Global Land-Ocean Temperature Anomaly ($^{\circ}\text{C}$)

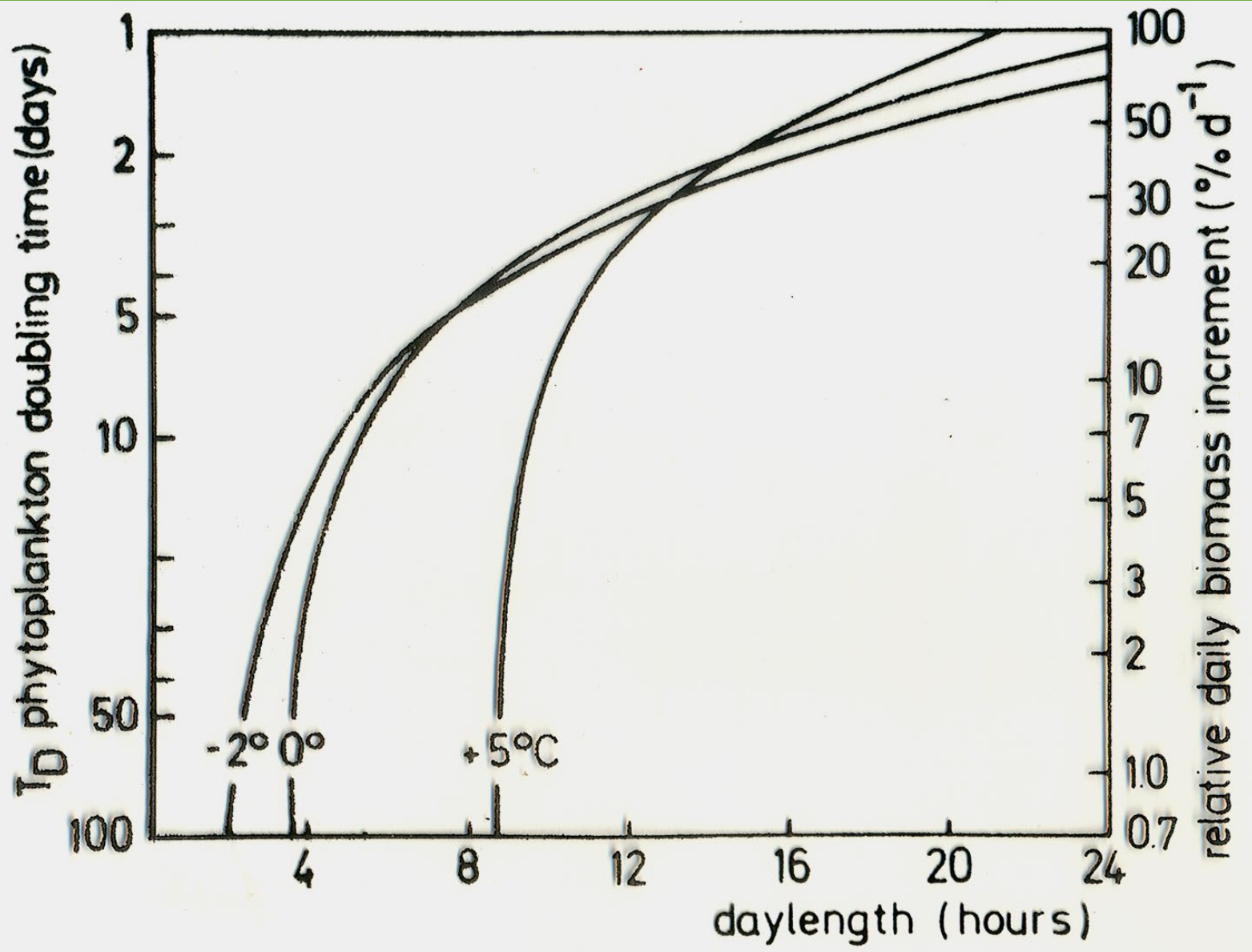
Base Period = 1951-1980



2008 Surface Temperature Anomaly ($^{\circ}\text{C}$)

Global Mean = 0.44





Aim: To realize the energy inputs of

the LCA to produce oil-rich algal

biomass for biodiesel.

We considered recent technologies

and main bottlenecks in algal biomass

production.

We calculated and compared the

energy requirements of the algal

cultivation systems.

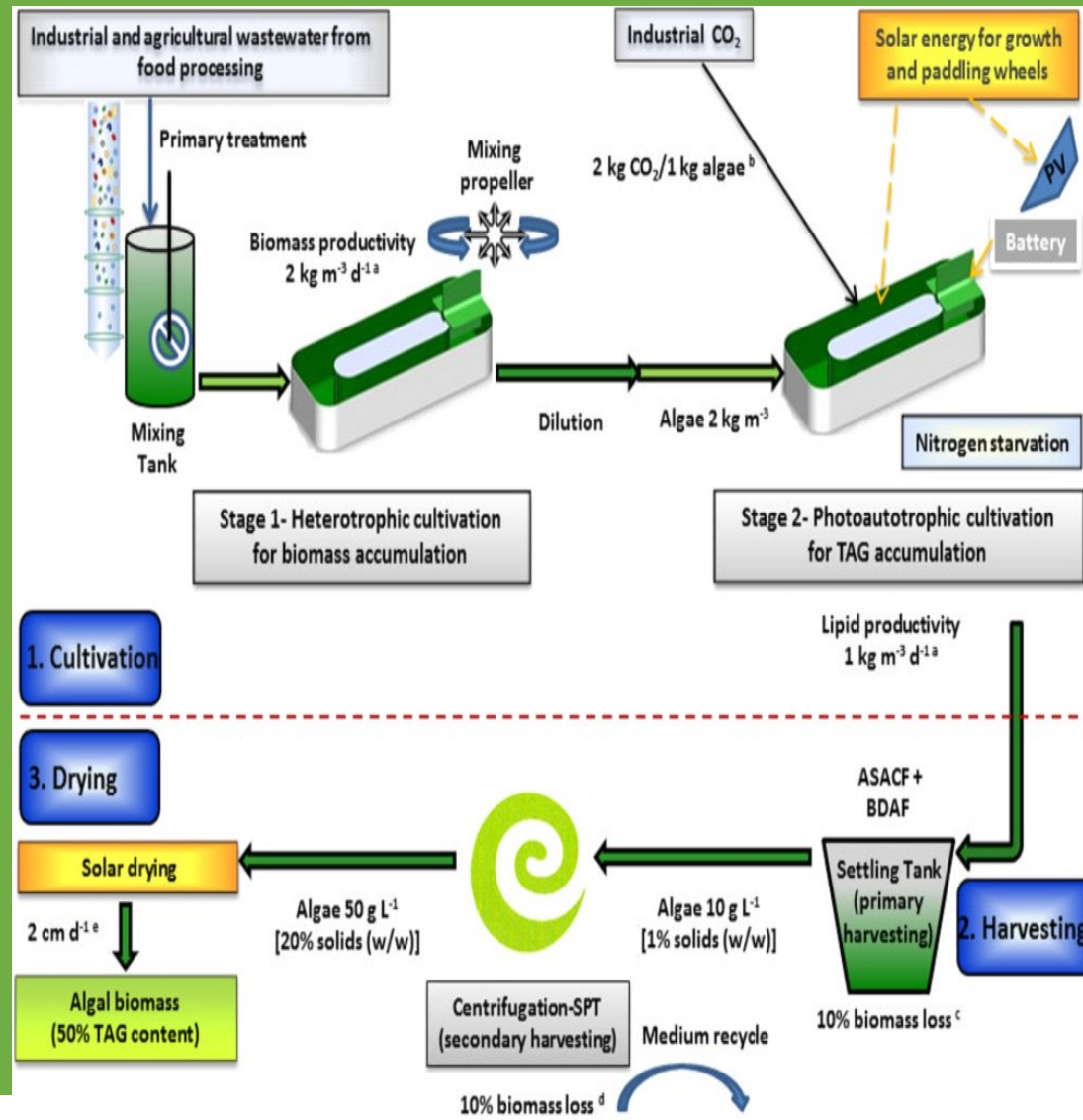
Lowest energy input for a complete

harvesting process requires

4 MJ per m³ algal suspension.

We propose the best scenario to grow

microalgae for biodiesel.



The end

